

EXTENZNÍ PÁNVE

(Pánve vzniklé natahováním litosféry)

intracratonic sags, rifty, pasivní okraje

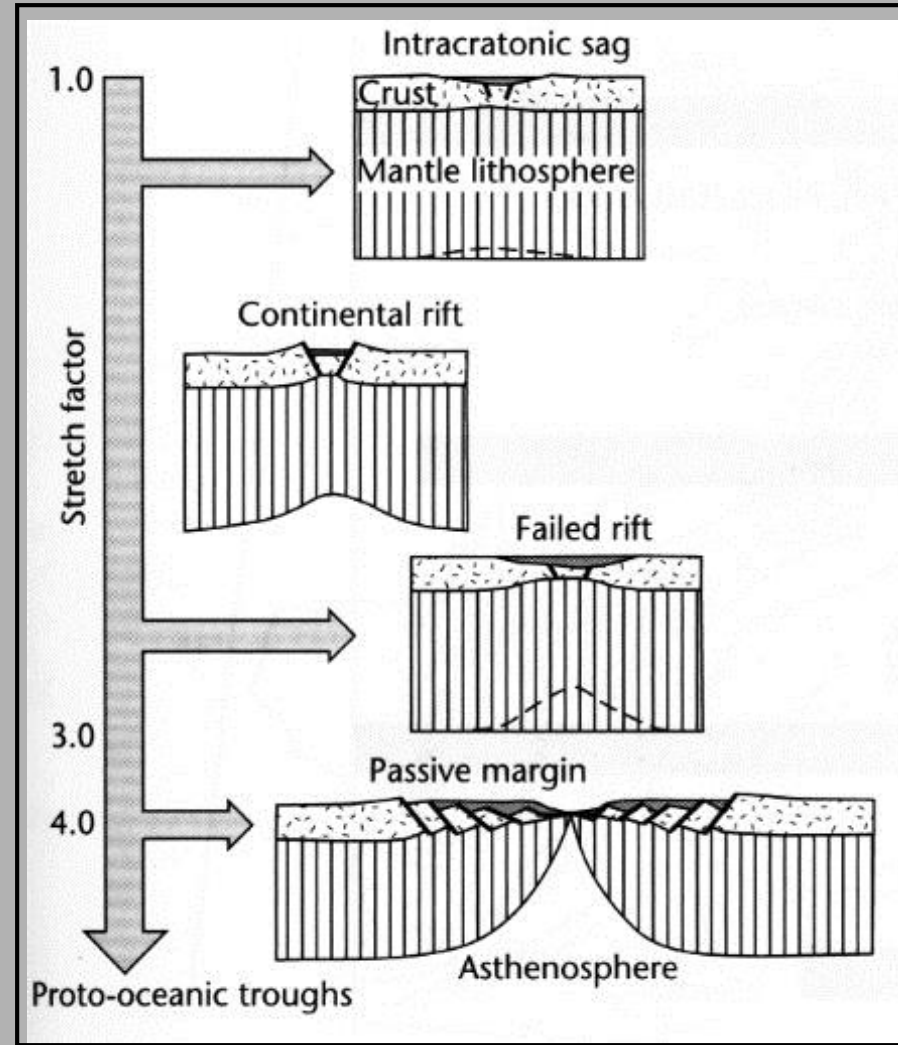
-všechny tyto pánve vznikají v extenzním režimu

Projevy extenze

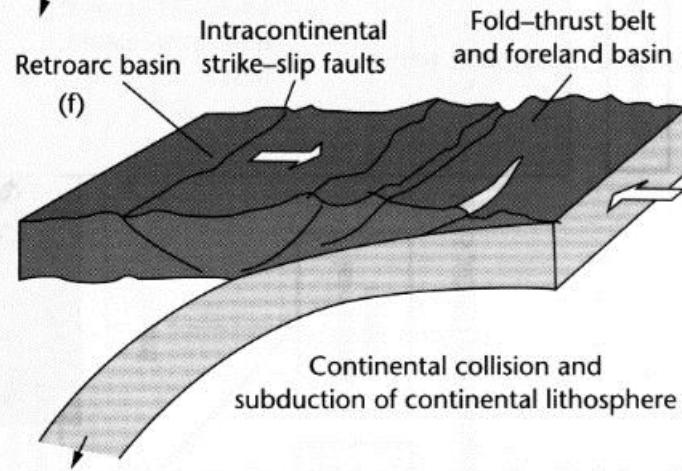
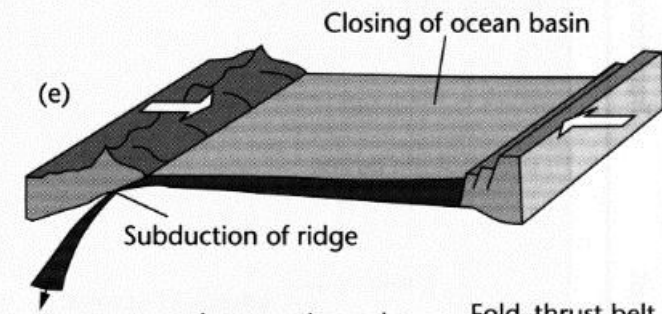
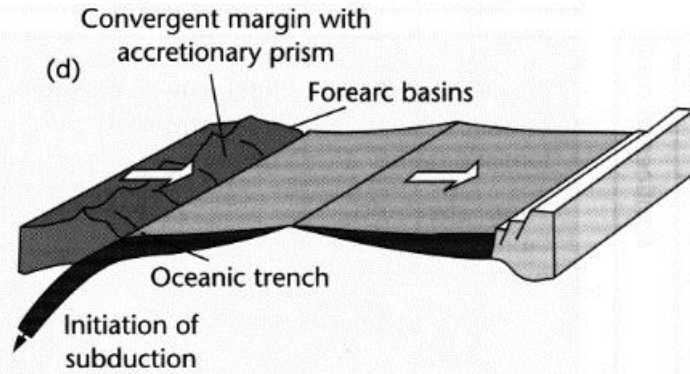
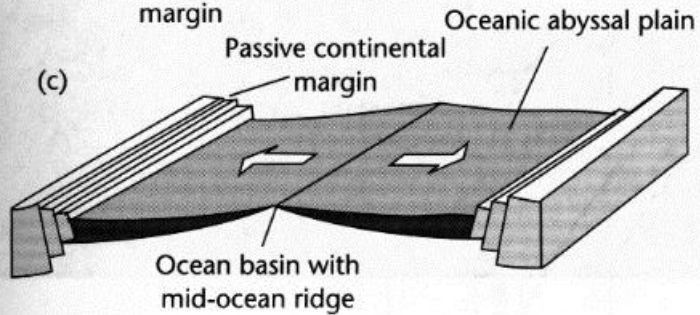
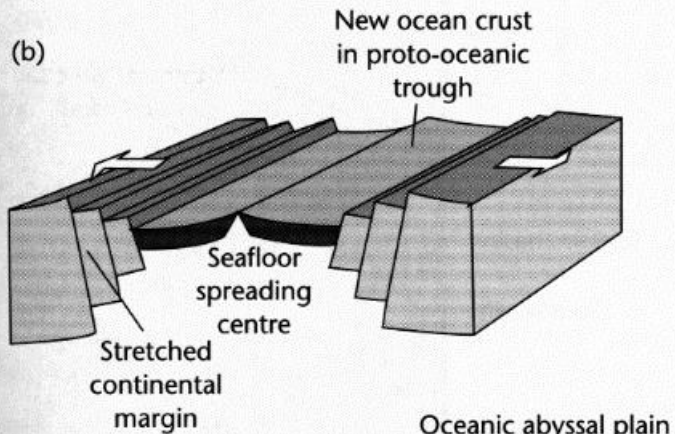
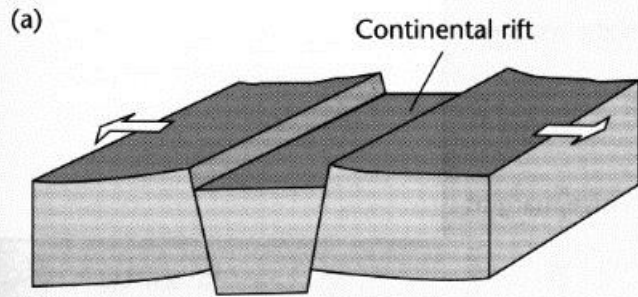
➤ **křehké** ⇒ rifty

➤ **duktilní** ⇒ intracratonic sags

(není překročena pevnost hornin)



THE WILSON CYCLE



RIFTOVÉ PÁNVE



RIFT

-příkopová struktura nebo systém příkopových struktur vzniklý v oblasti **ztenčení kůry** a vymezený soustavou poklesových zlomů.

- mělká hloubka MOHO
- vysoký tepelný tok při povrchu
- vulkanická aktivita
- negativní Bouguerovy tíhové anomálie
- často vyzdvižené okraje riftové struktury

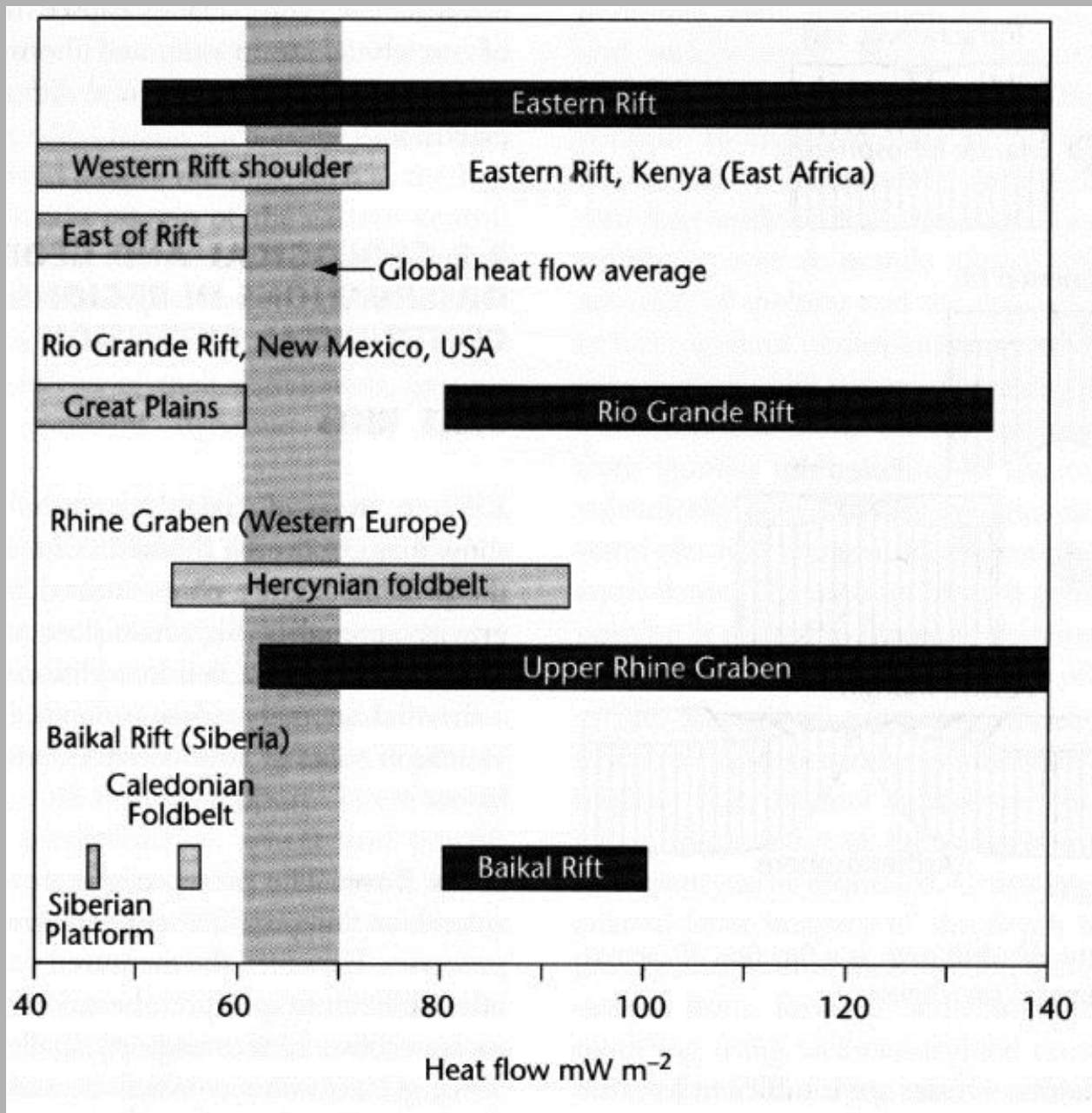
TEPELNÝ TOK

Vulkanická činnost a zvýšený tepelný tok dokládají zvýšenou aktivitu termálních procesů.

Průmerný tepelný tok v riftových oblastech **90-110 mWm⁻²**

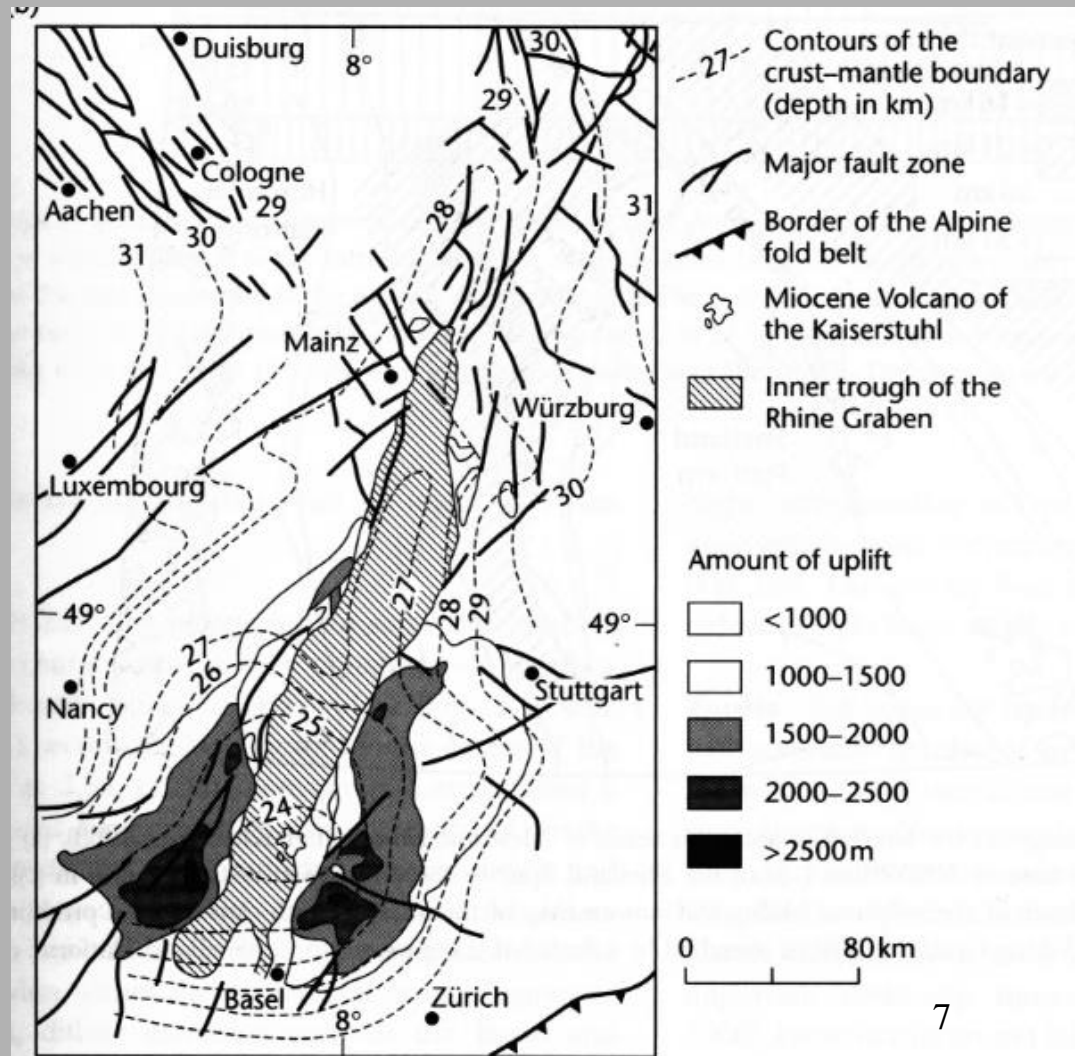
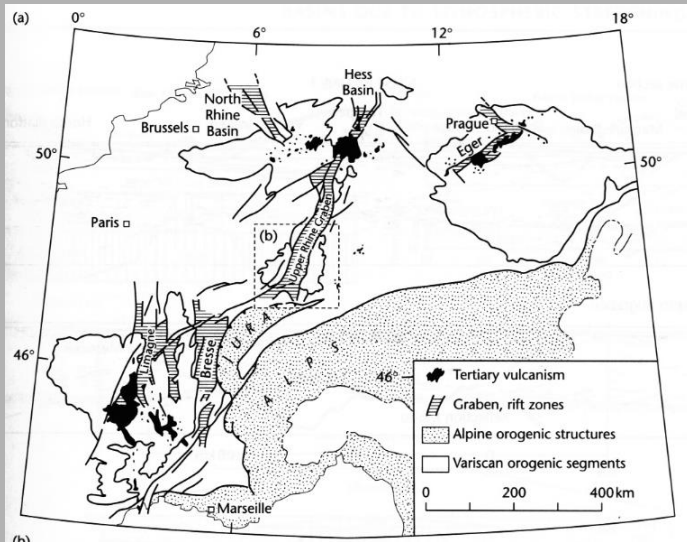
⇒ to je 2x více než v "neriftovaném" okolí

Hodnota tepelného toku závisí rovněž na horninovém složení (granity mají vyšší tok než bazika nebo sedimenty).

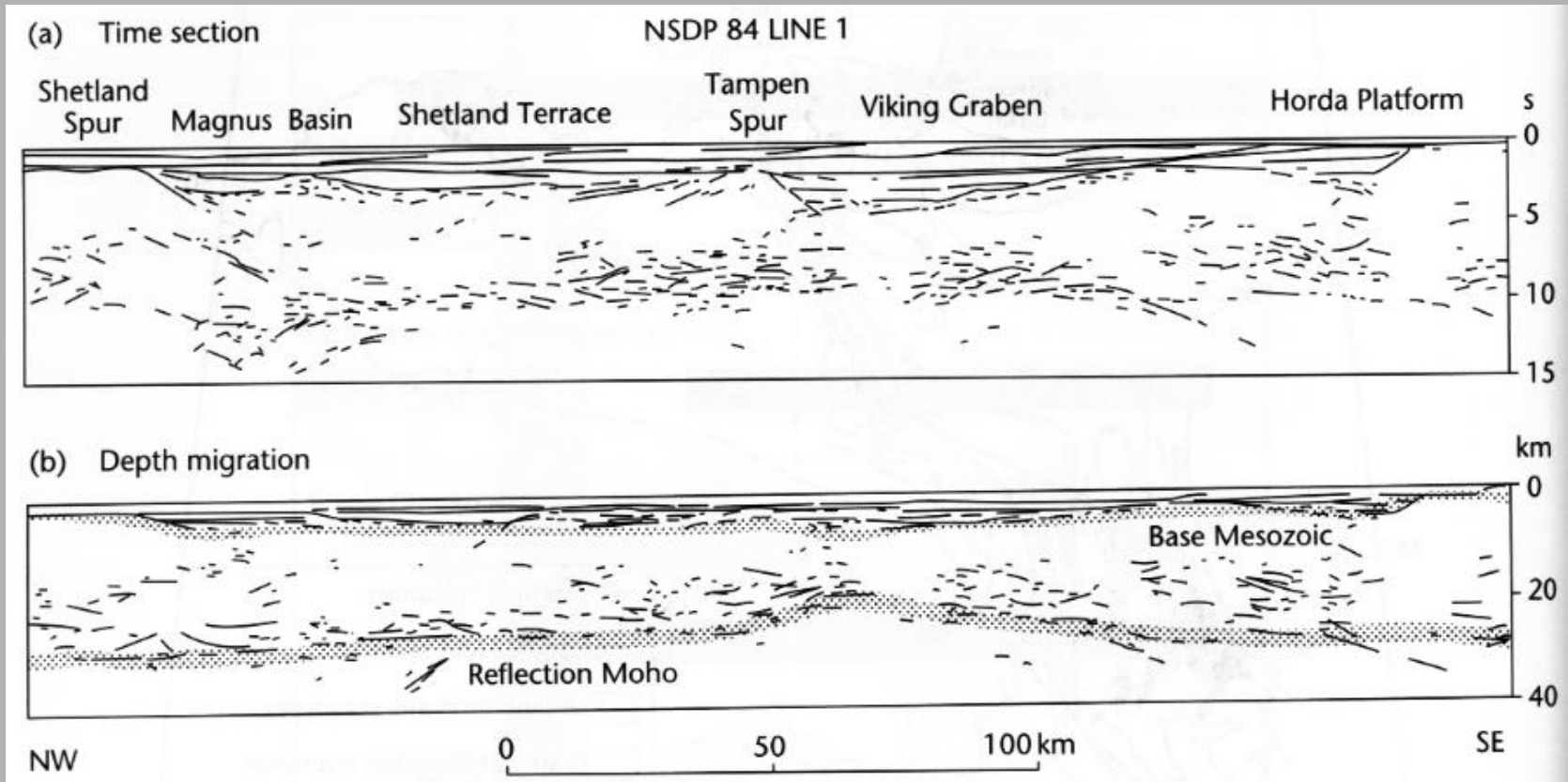


MOCNOST KŮRY

Vyklenování MOHO v místě riftingu \Rightarrow ztenčení kůry

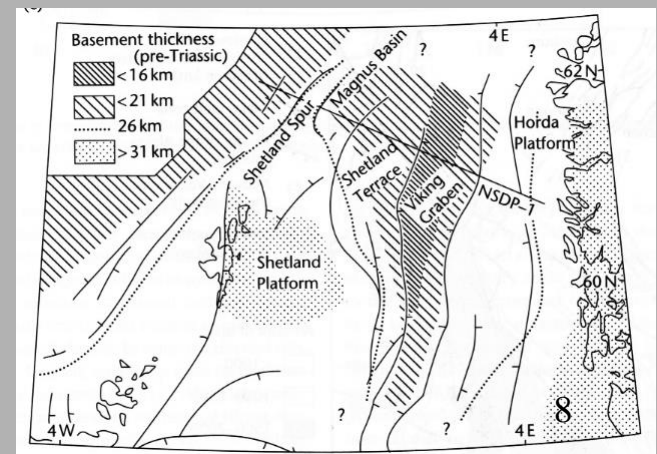


Allen and Allen, 2005



Oblast severního moře – rifting v mesozoiku

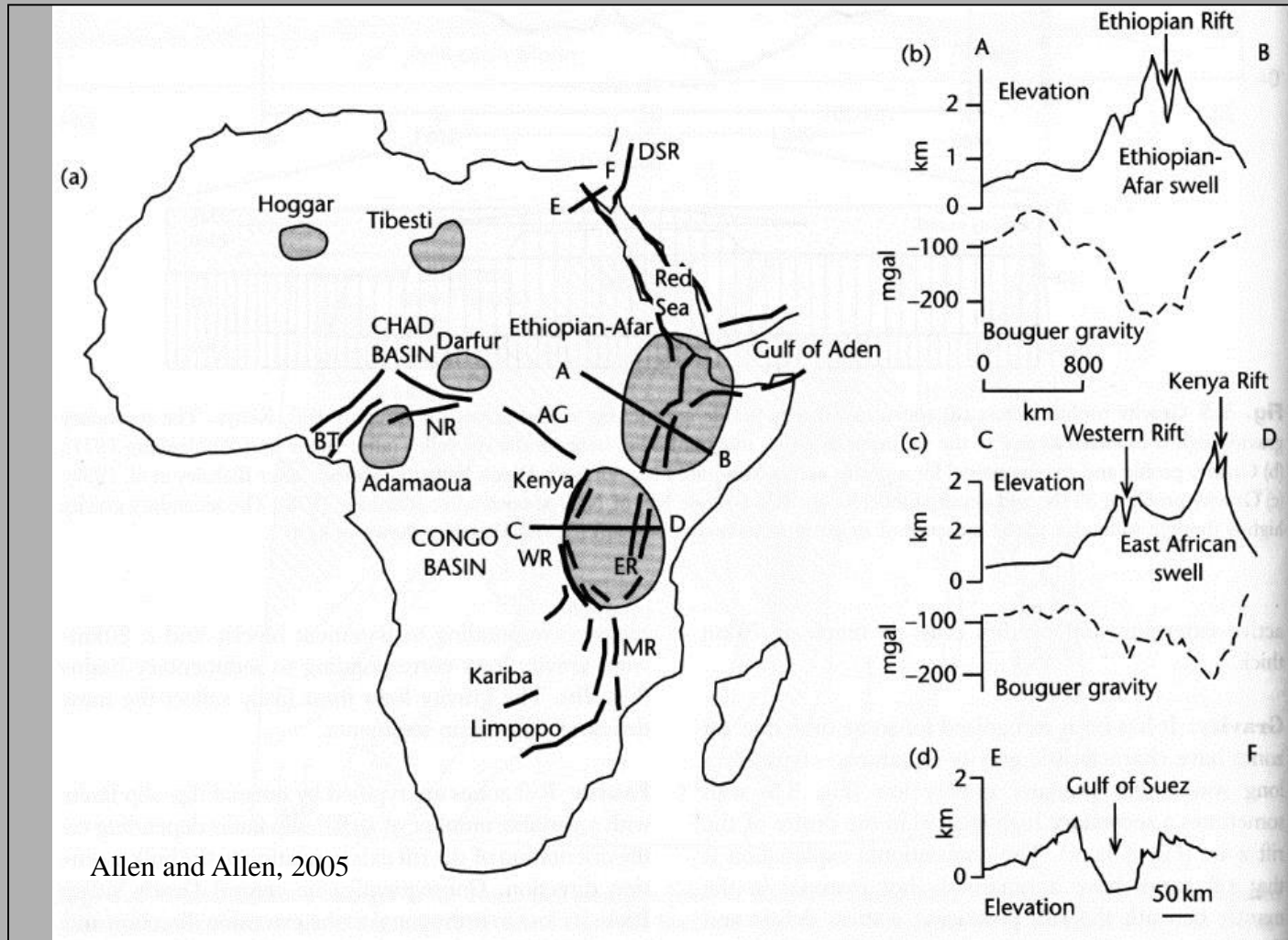
Allen and Allen, 2005

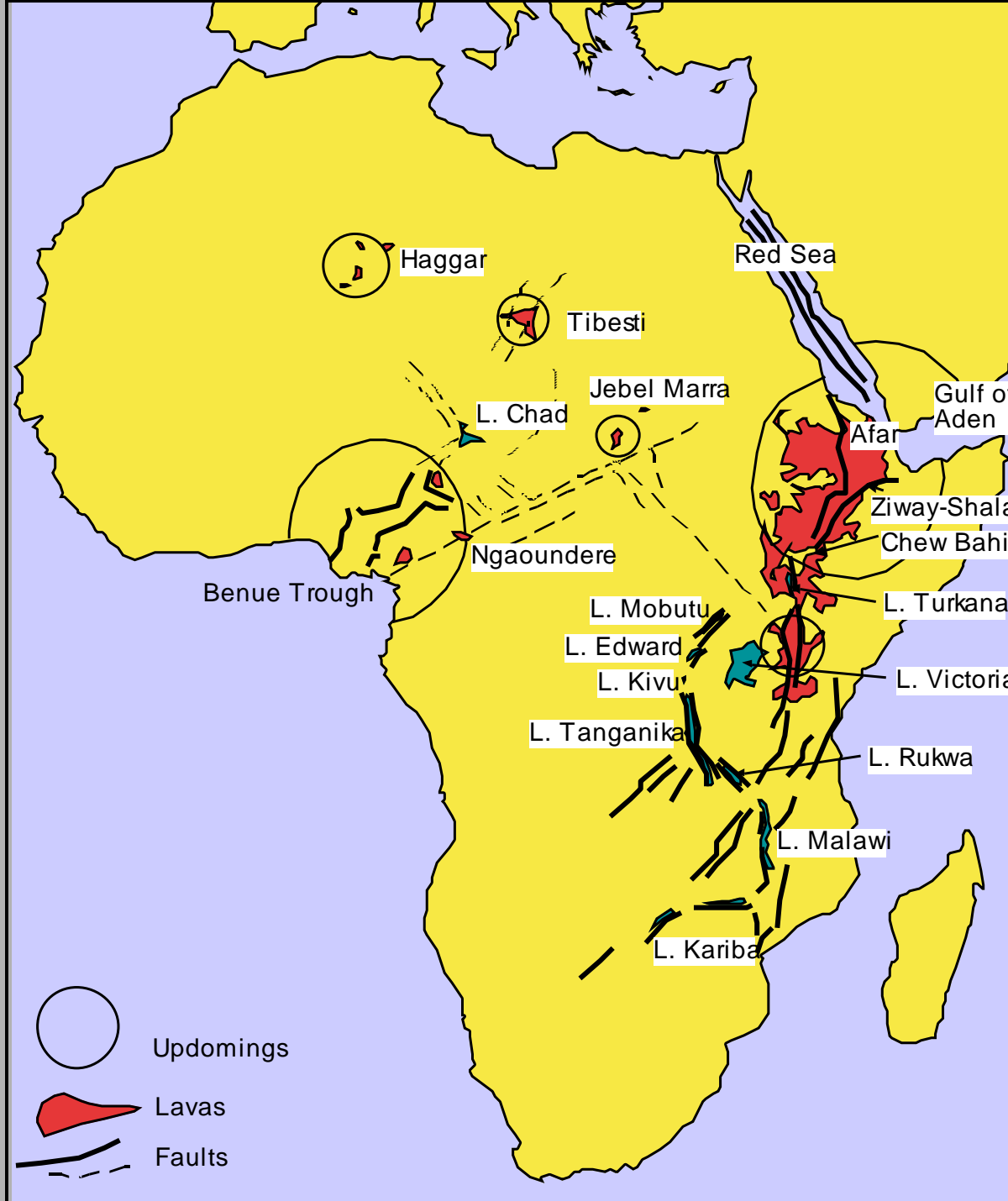


TOPOGRAFIE

- současné rifty jsou typické **výzdvíženými** okraji

⇒ **eroze** (např. obnažení Herc. krystalinika v oblasti Vosges a Černého lesa v případě výzdvihu okrajů riftu Rhine Graben)



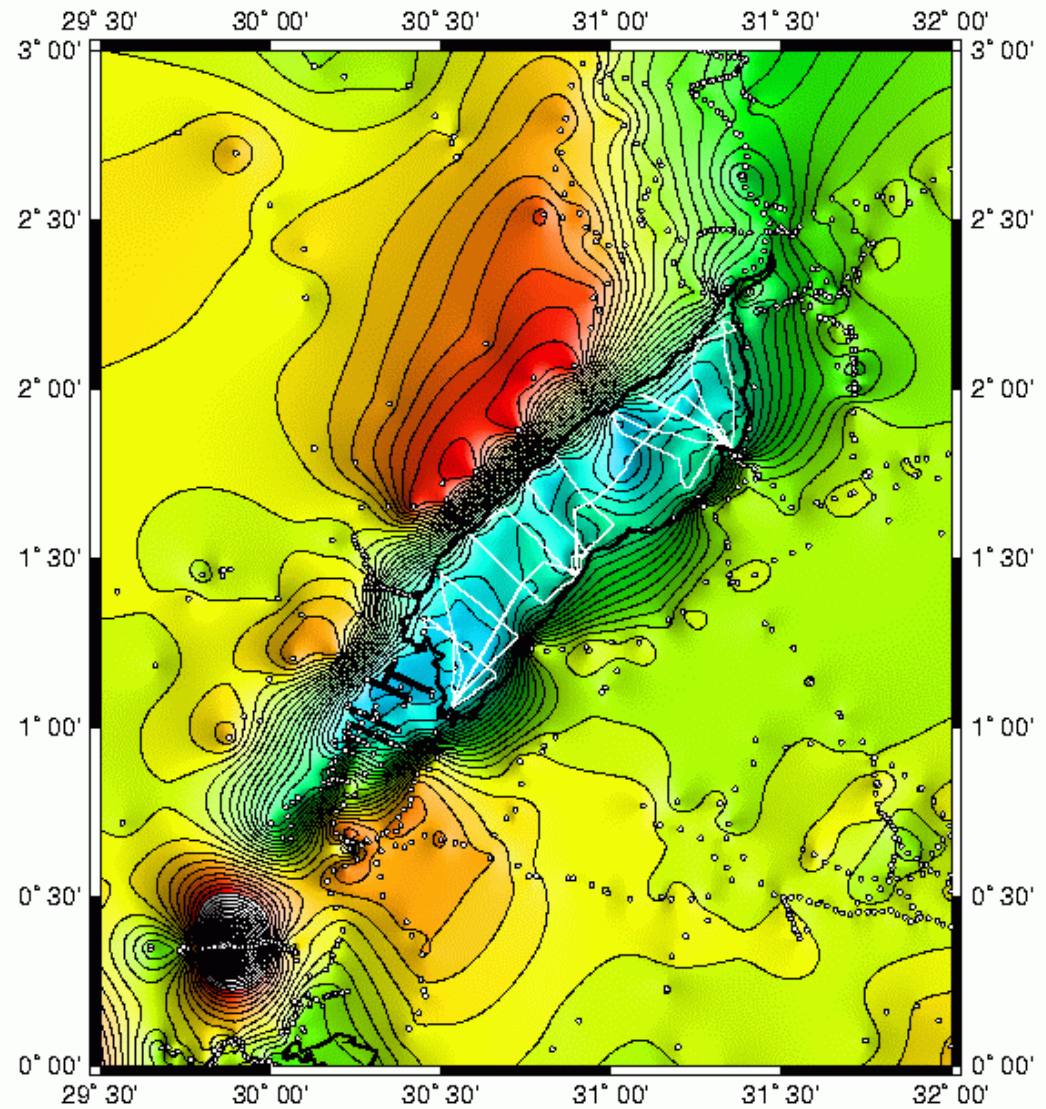


GRAVIMETRIE

Zóna riftingu - pod povrchem

prohřátý materiál

⇒ **negativní tíhové anomálie**



Free-air gravity (mgals)



Free-air gravity contour map of the Albertine rift system

ČASOVÉ MĚŘÍTKO a VELIKOST EXTENZE

- rychlost rozpínání < 1 mm/rok
- doba rozpínání 10 - 30 mil. let s celkovou extenzí do 10 km
(platí pro rifty na normálně mocné kůře; např. Rhine Graben, Bajkal, Rio Grande)

MODELY KONTINENTÁLNÍ EXTENZE

AKTIVNÍ a PASIVNÍ rifting - idealizované koncové modely

□ AKTIVNÍ RIFTING

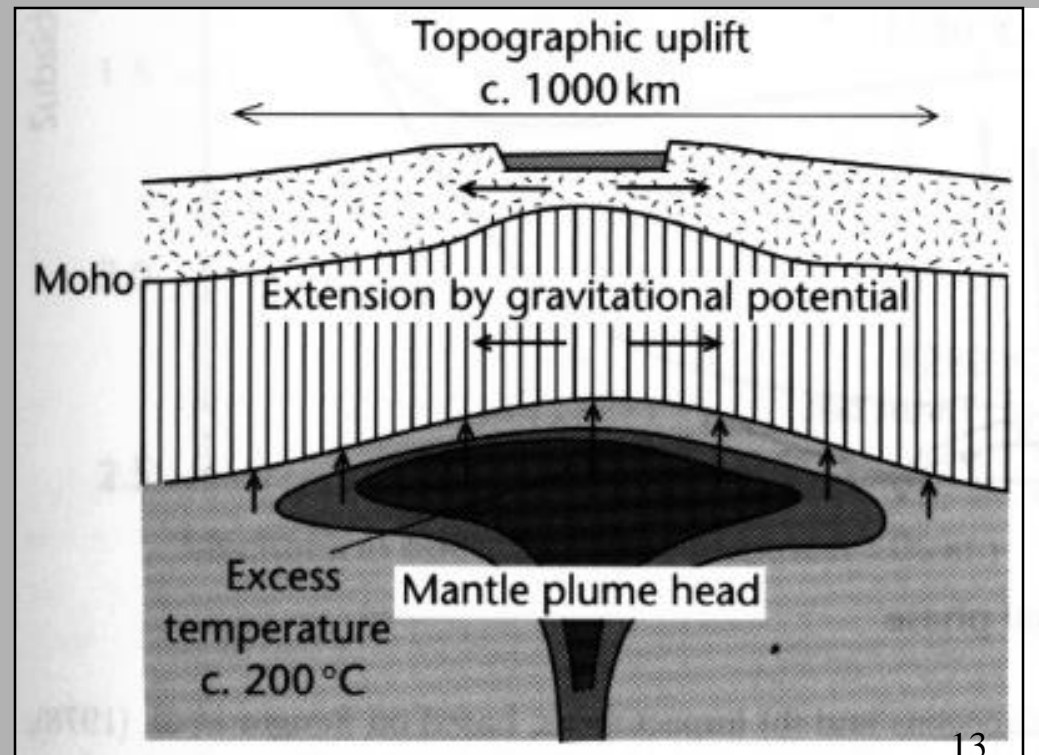
prohřátí a následné ztenčení litosféry nad plášt'ovým plumem

⇒ vyklenování (v důsledku zvýšeného teplotního toku a prohřátí)

⇒ tenzní napětí

⇒ rifting

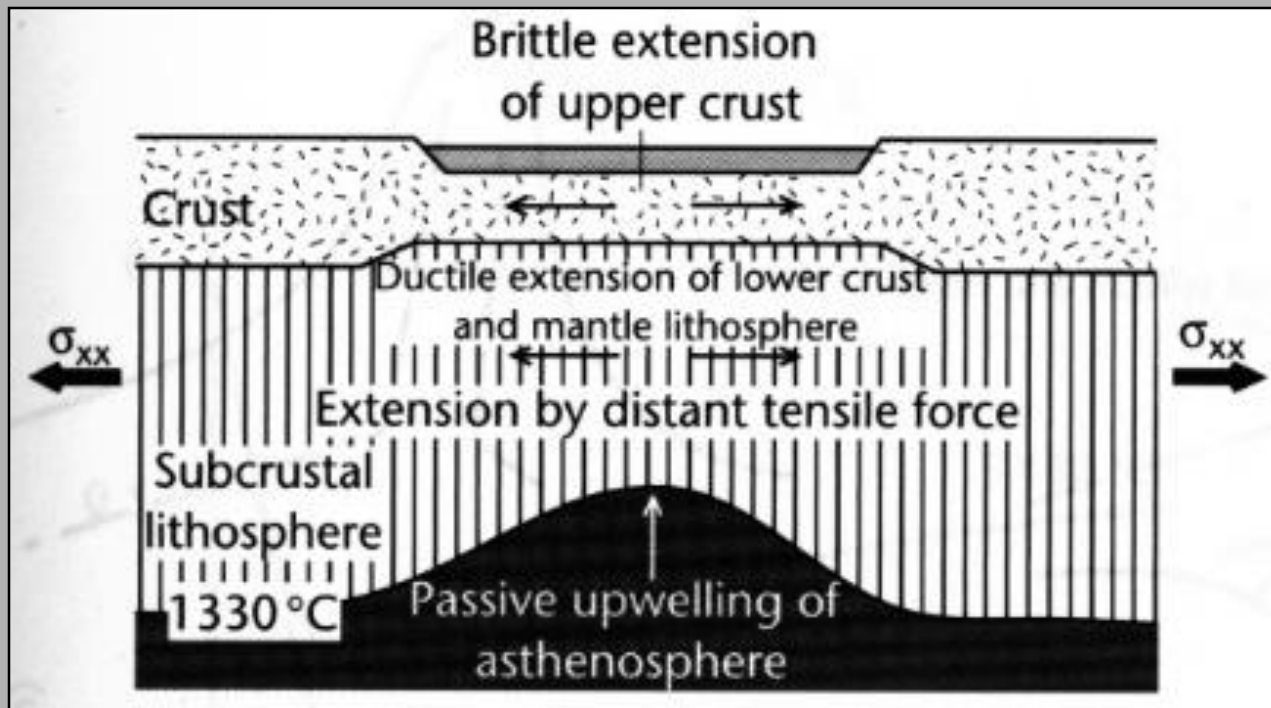
(např. East Africa Rift)



□ PASIVNÍ RIFTING

- ztenčení a porušení kontinentální litosféry v důsledku **tenze**
⇒ vyklenování a vulkanická činnost (druhotné procesy)

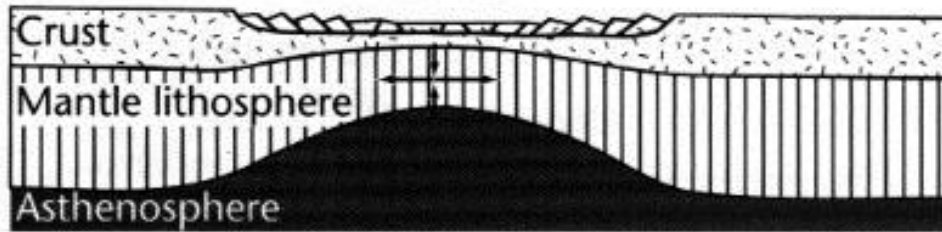
(např. Rift Rio Grande)



Allen and Allen, 2005



(a) **PURE SHEAR**



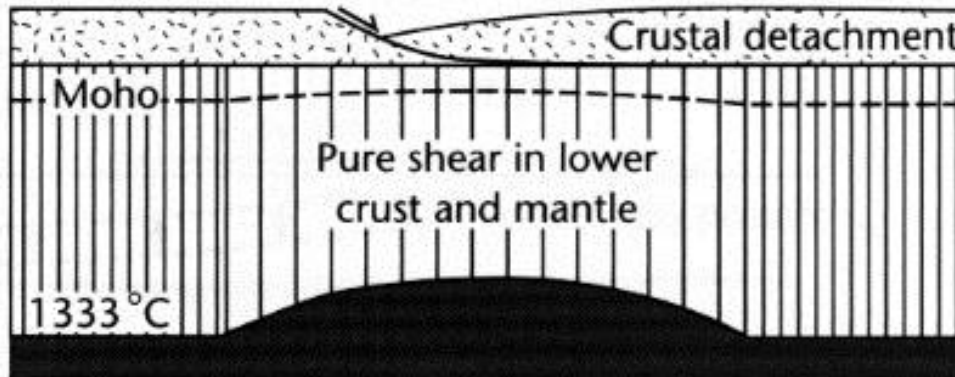
(b) **SIMPLE SHEAR**

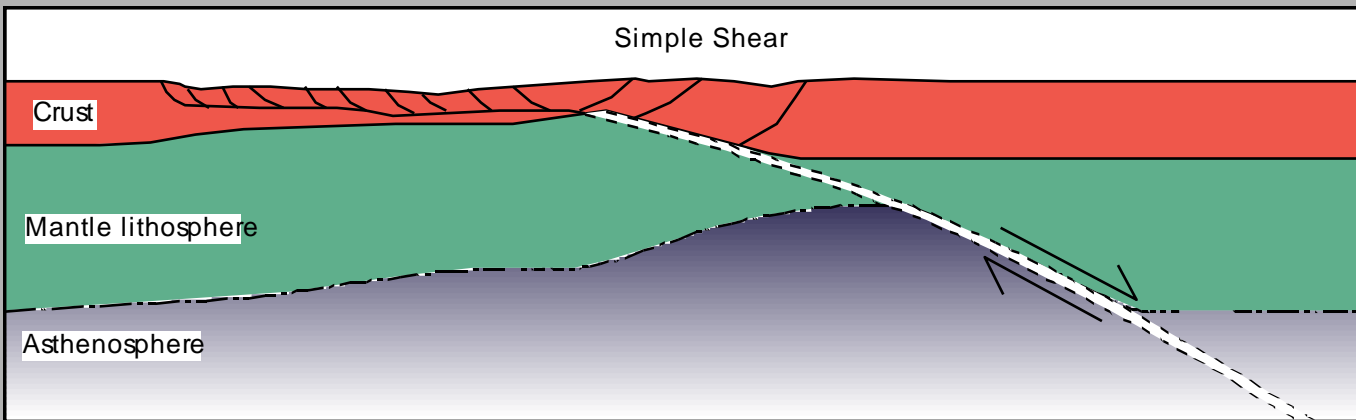
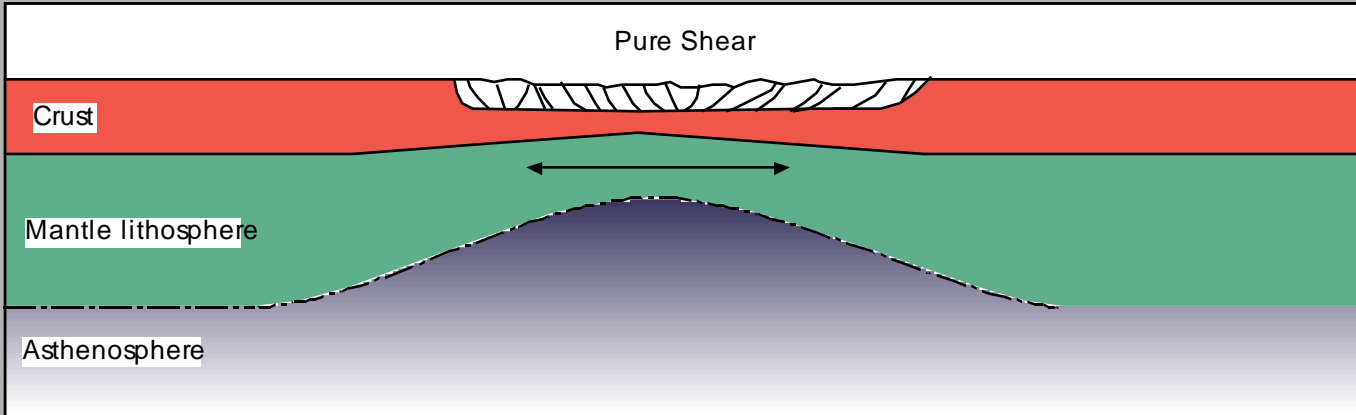
———— Subcrustal stretching —————
—— Crustal stretching ——

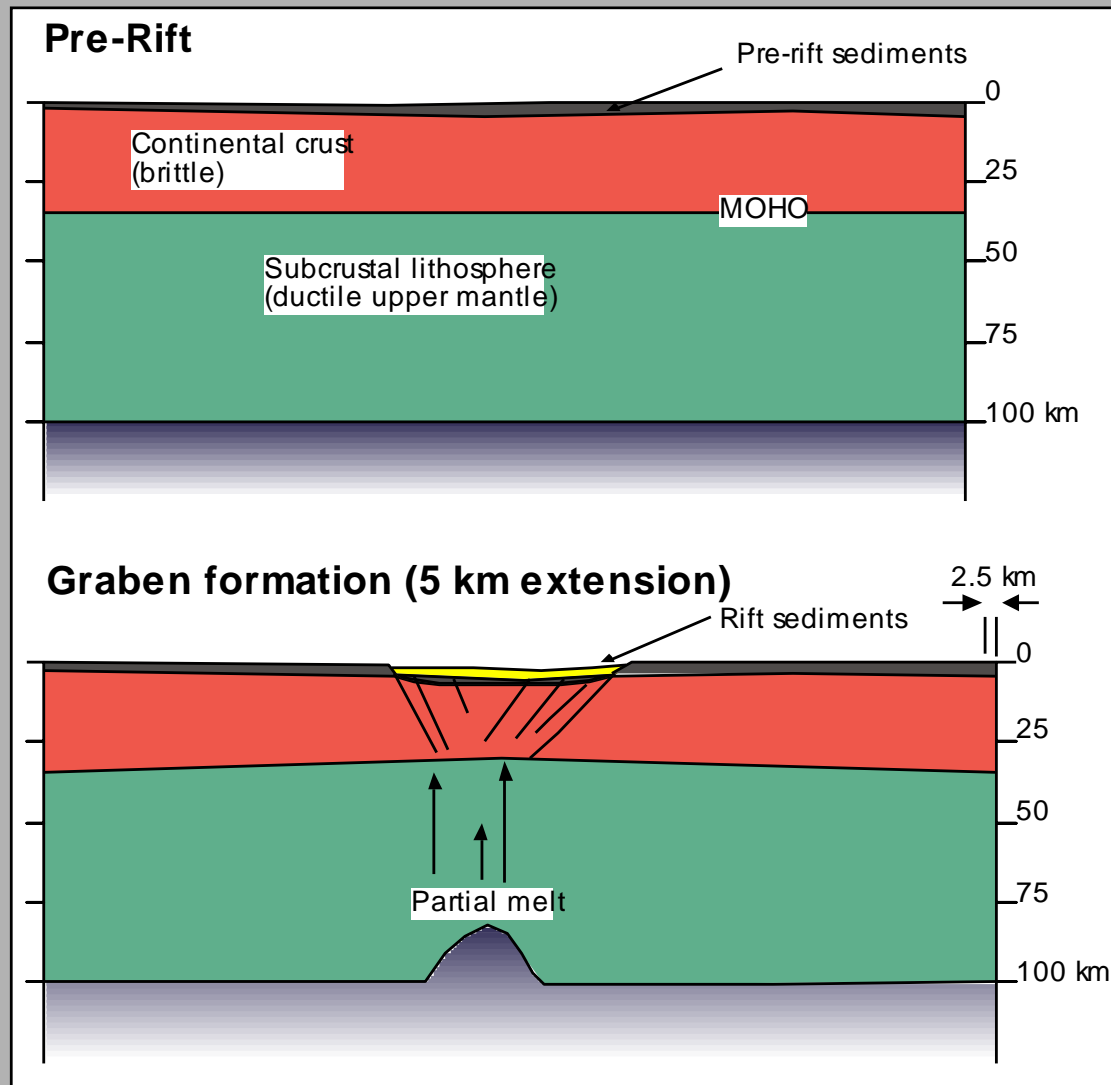


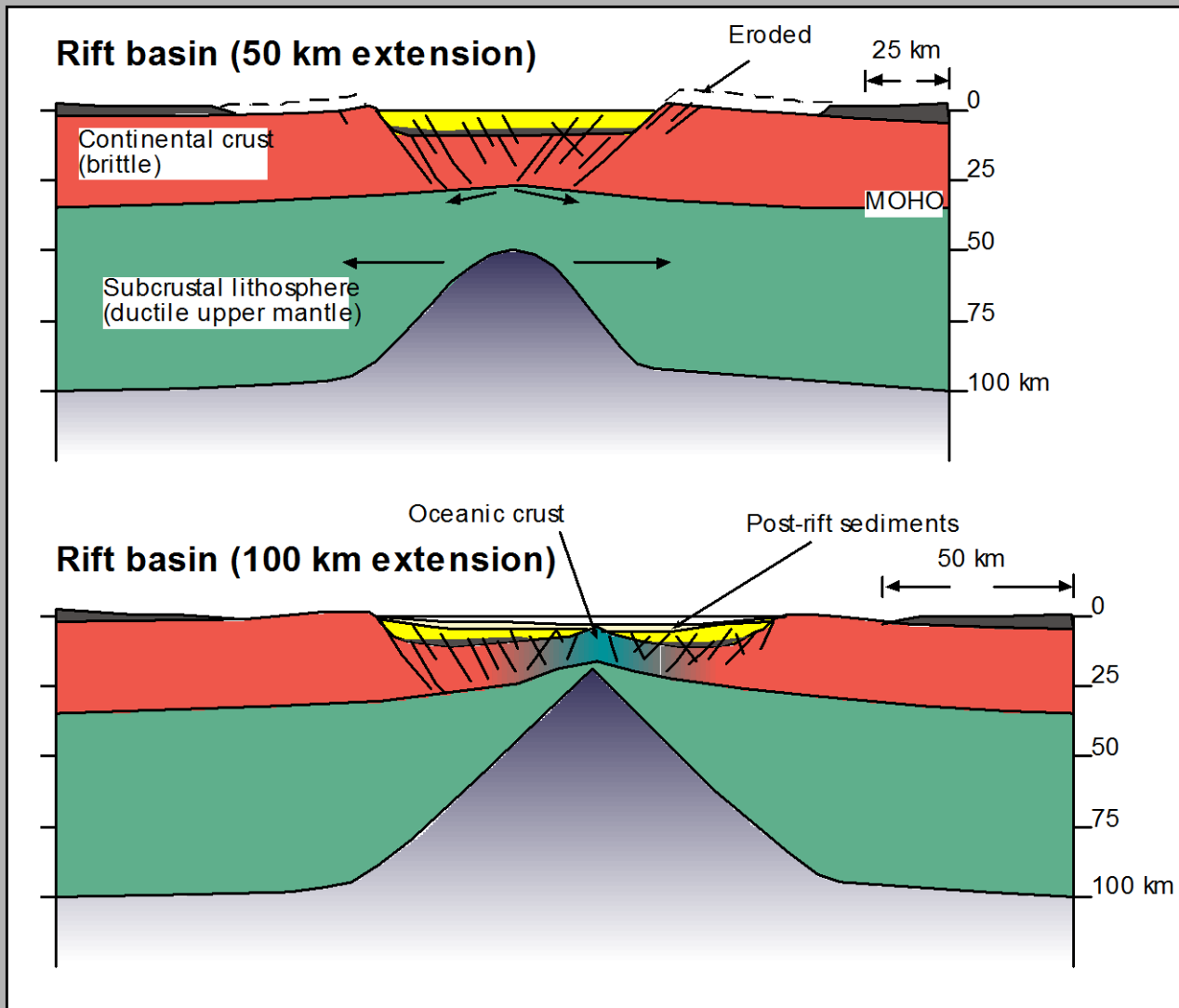
(c) **SIMPLE SHEAR–PURE SHEAR**

Simple shear in upper crust

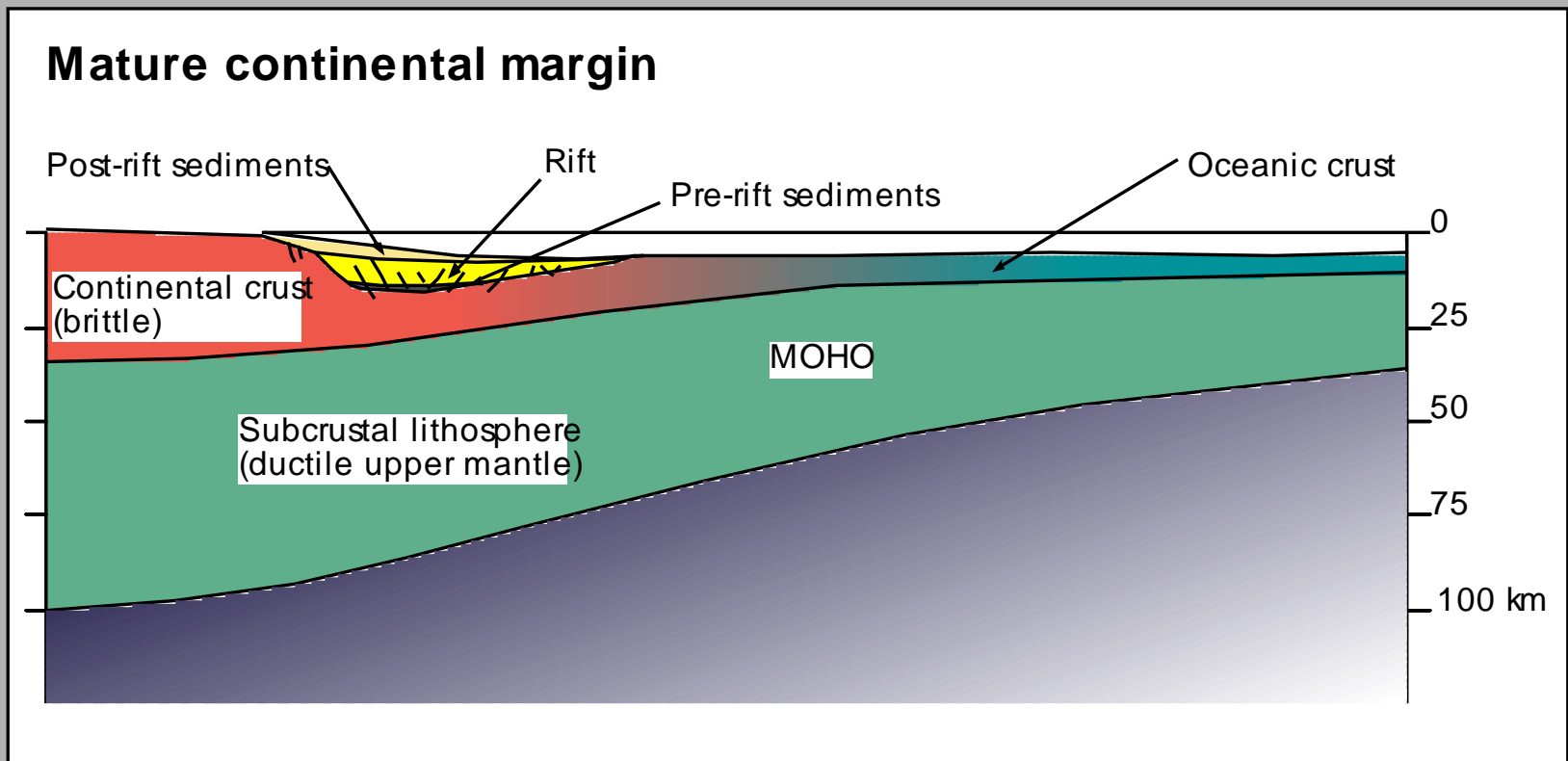








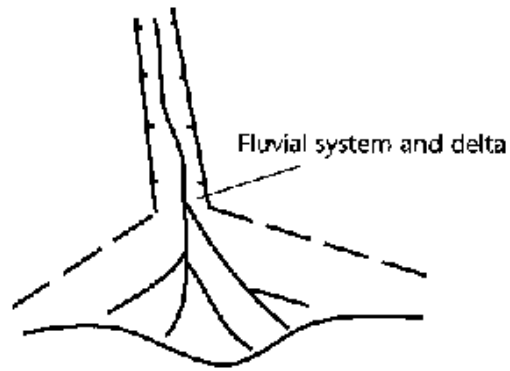
pasivní kontinentální okraj



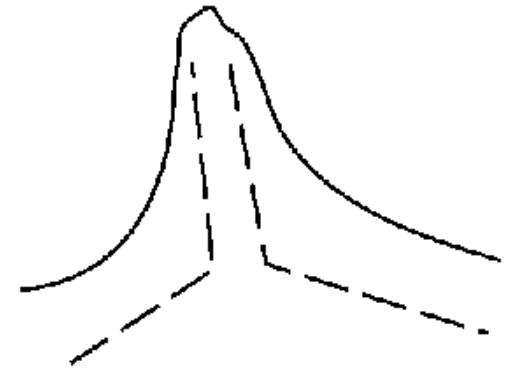
(a) **AULACOGENS**



RRR Triple junction

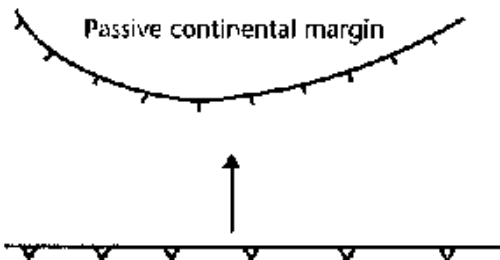


Ocean opening along two successful rift arms

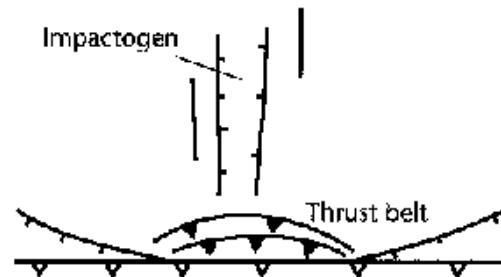


Postrift marine phase

(b) **IMPACTOGENS**



Passive continental margin
Convergent margin
Precollision phase



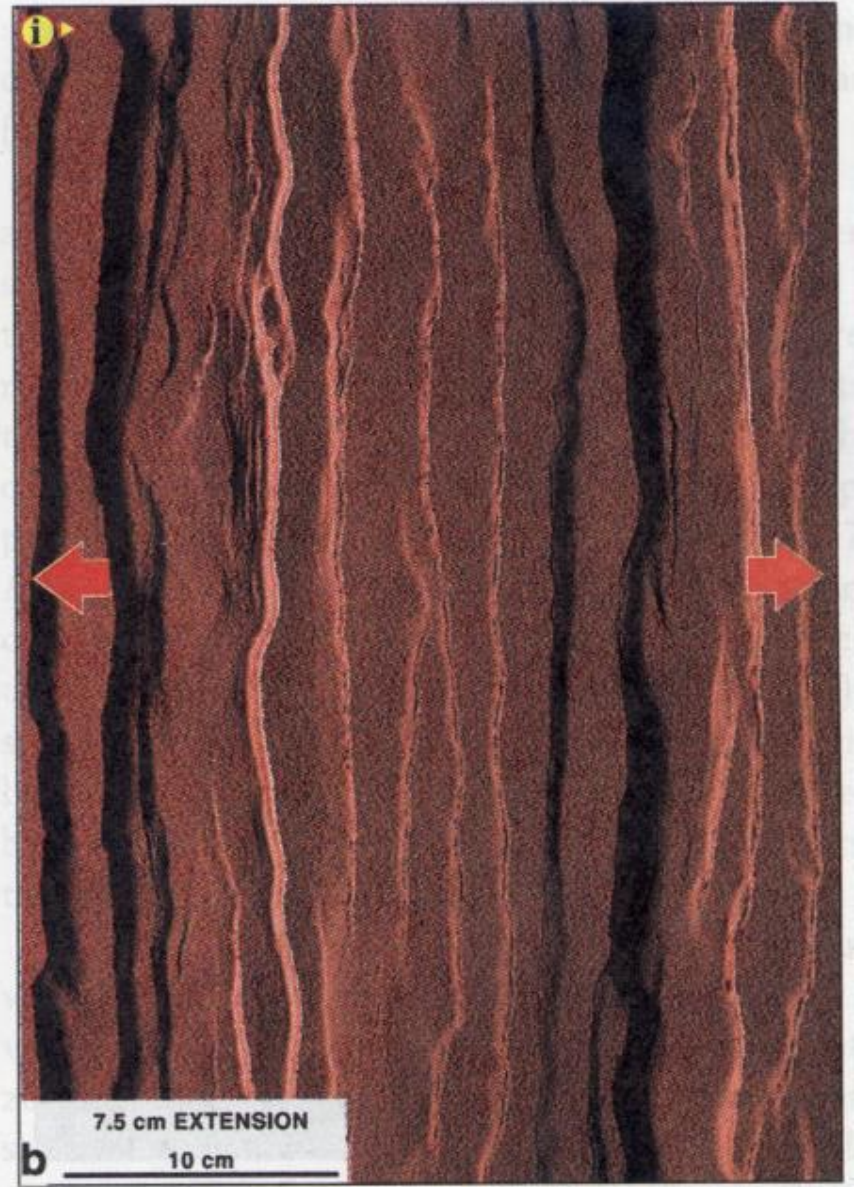
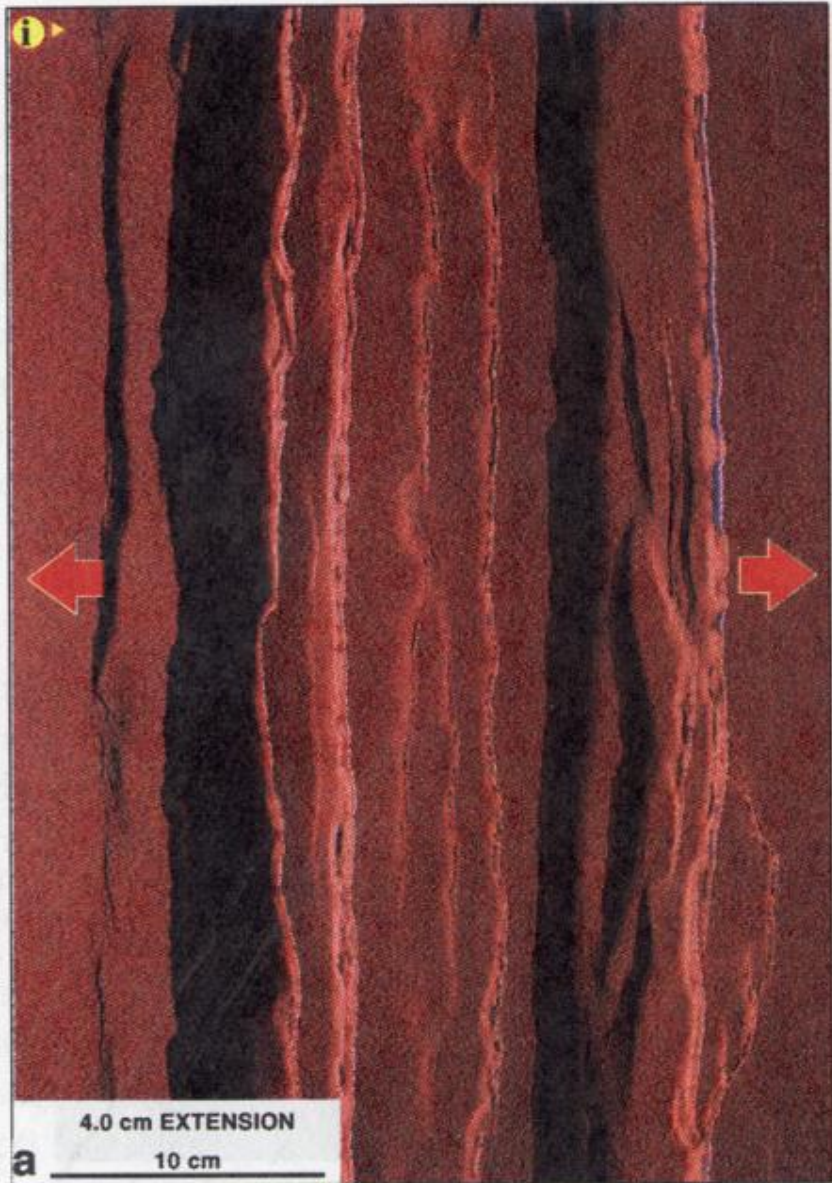
Impactogen
Thrust belt
Collisional phase

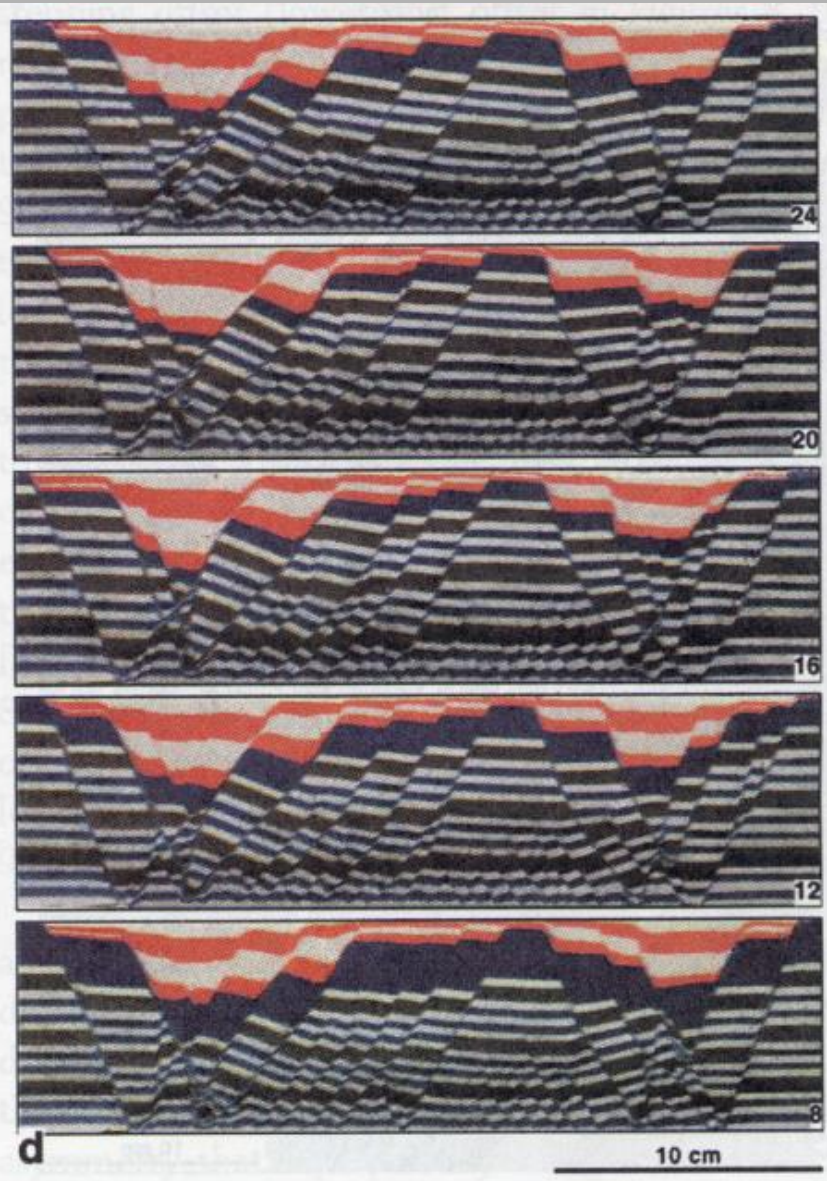
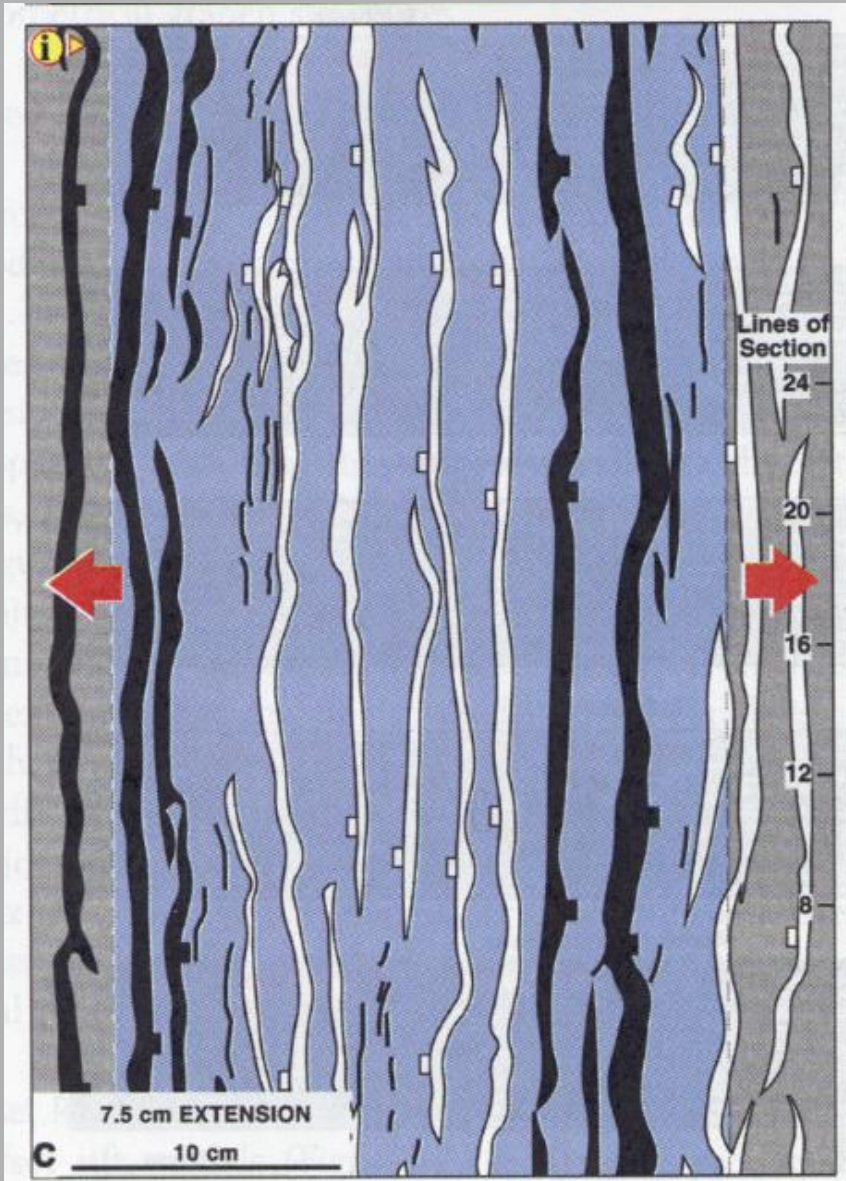
Allen and Allen, 2005

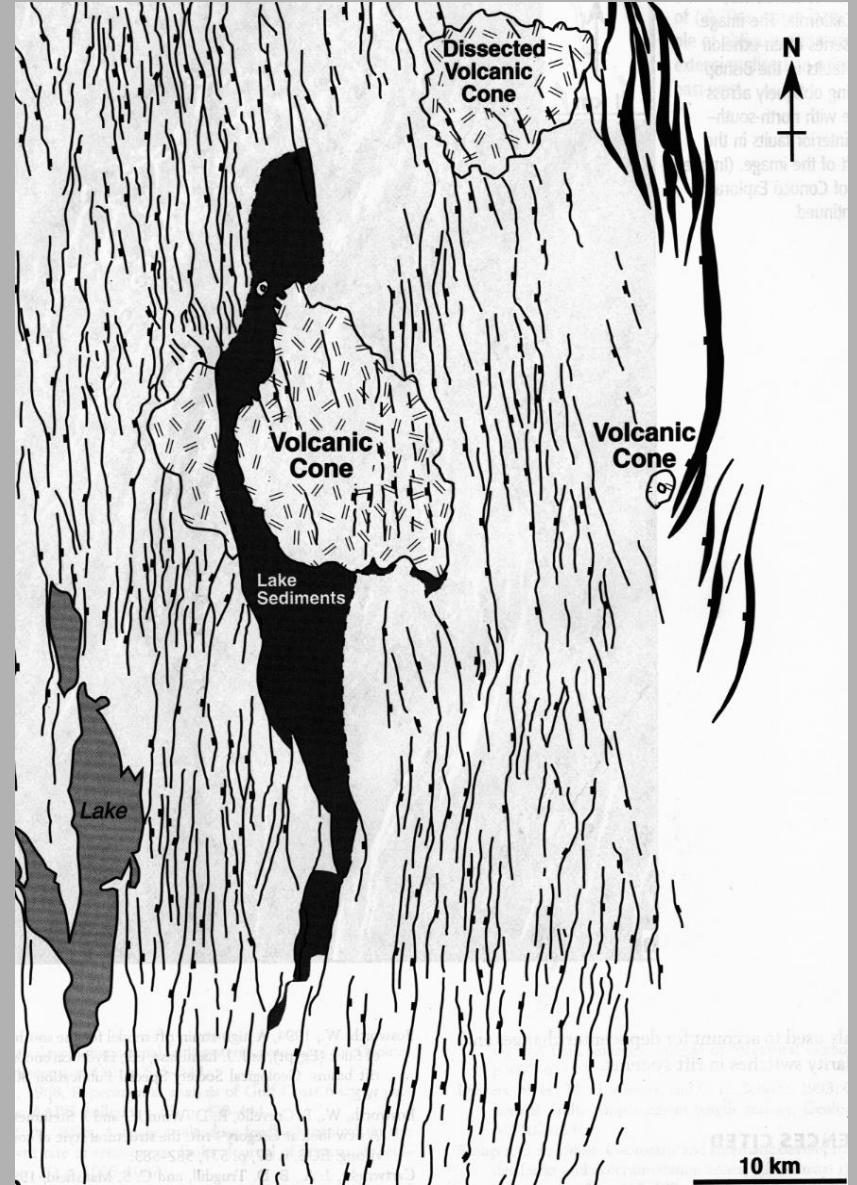
ARCHITEKTURA RIFTOVÝCH PÁNVÍ: ZLOMOVÉ SYSTÉMY A GEOMETRIE DEPOCENTER

Riftové pánve jsou tvořeny jednou nebo více příkopovými strukturami (depocentry) vymezenými poklesovými zlomy (sklon $45 - 70^\circ$) a často doprovázenými množstvím strike-slipových zlomů.

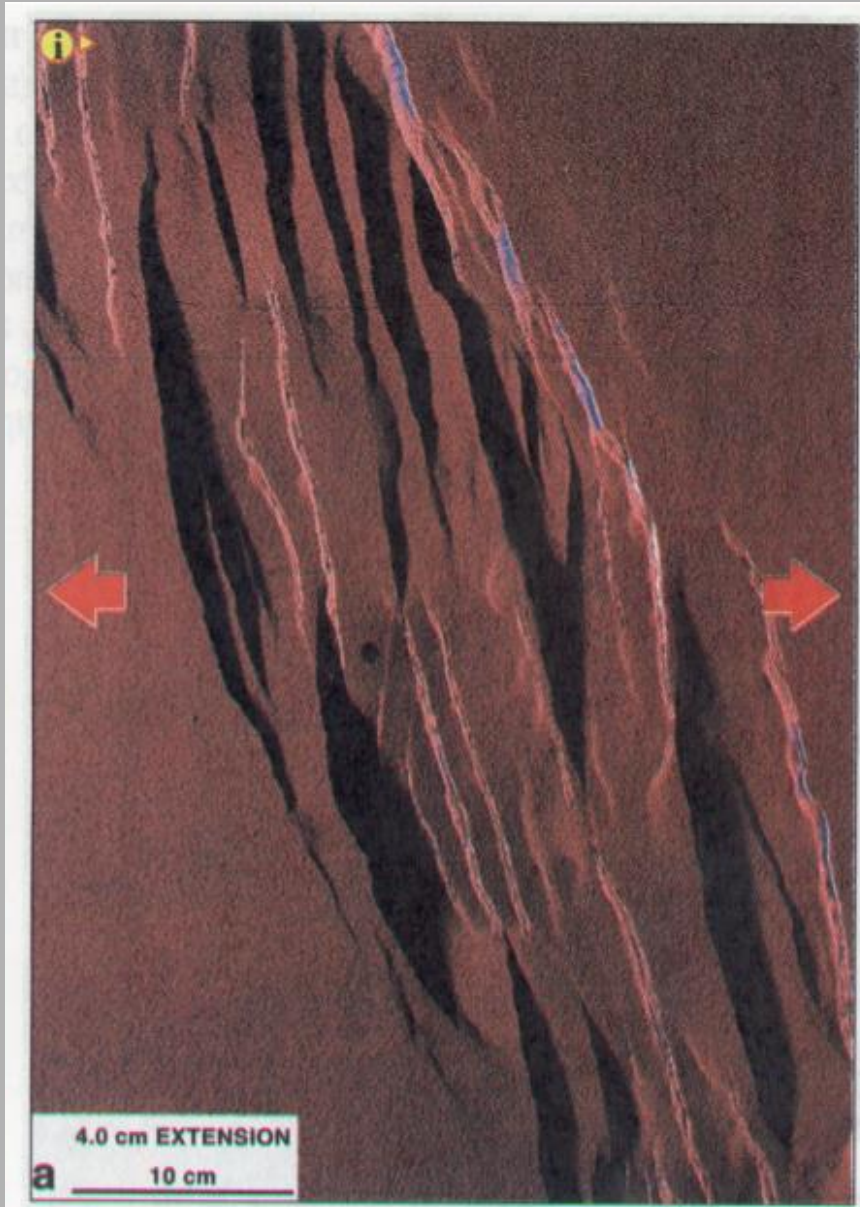
Geometrie depocenter, povaha (pokles/strike-slip) a uspořádání zlomů závisí na orientaci vektoru extenze a osy riftu.

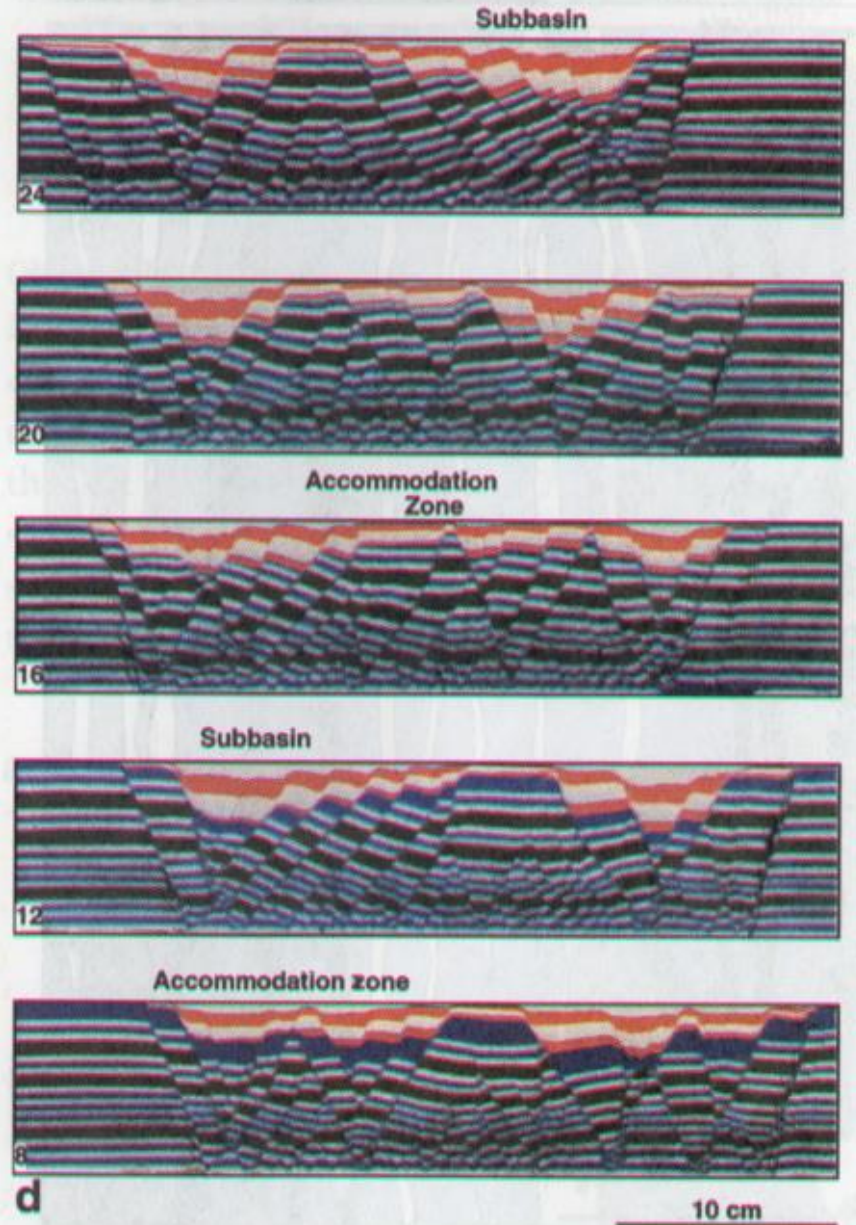
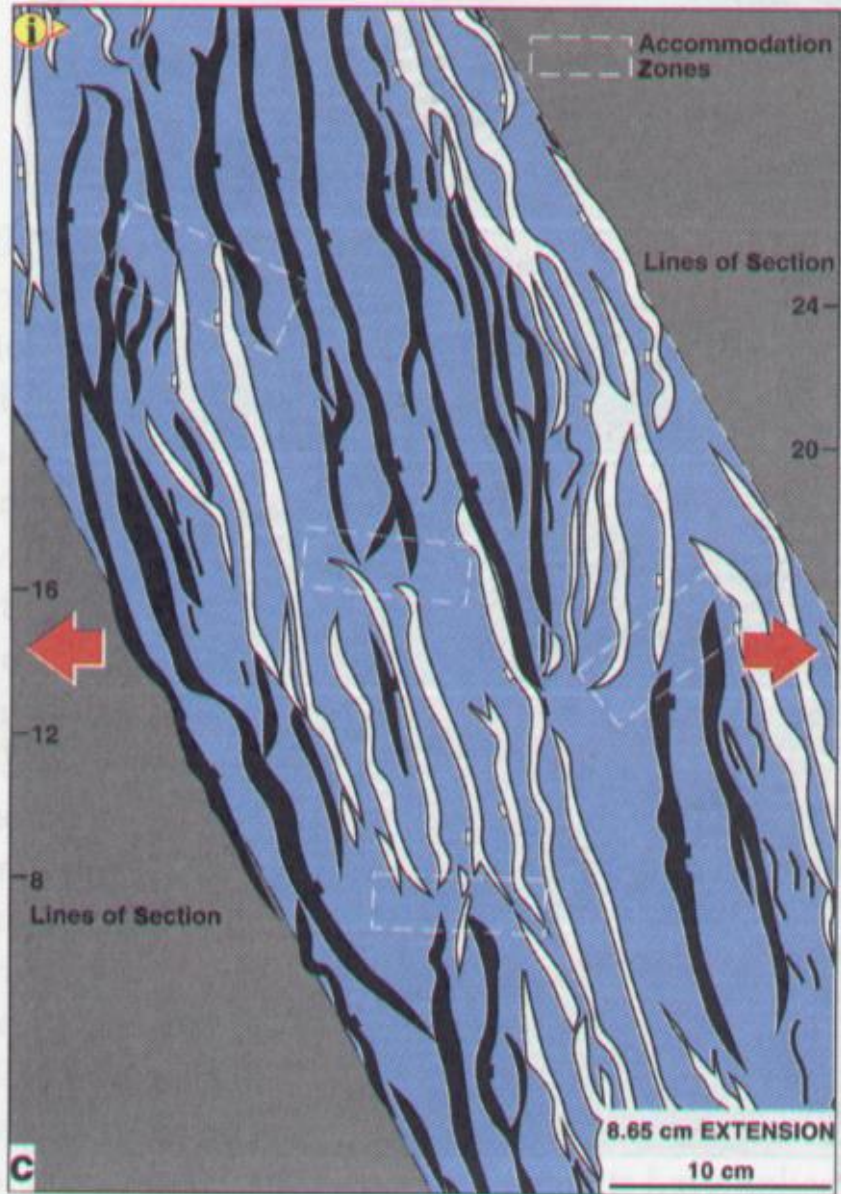


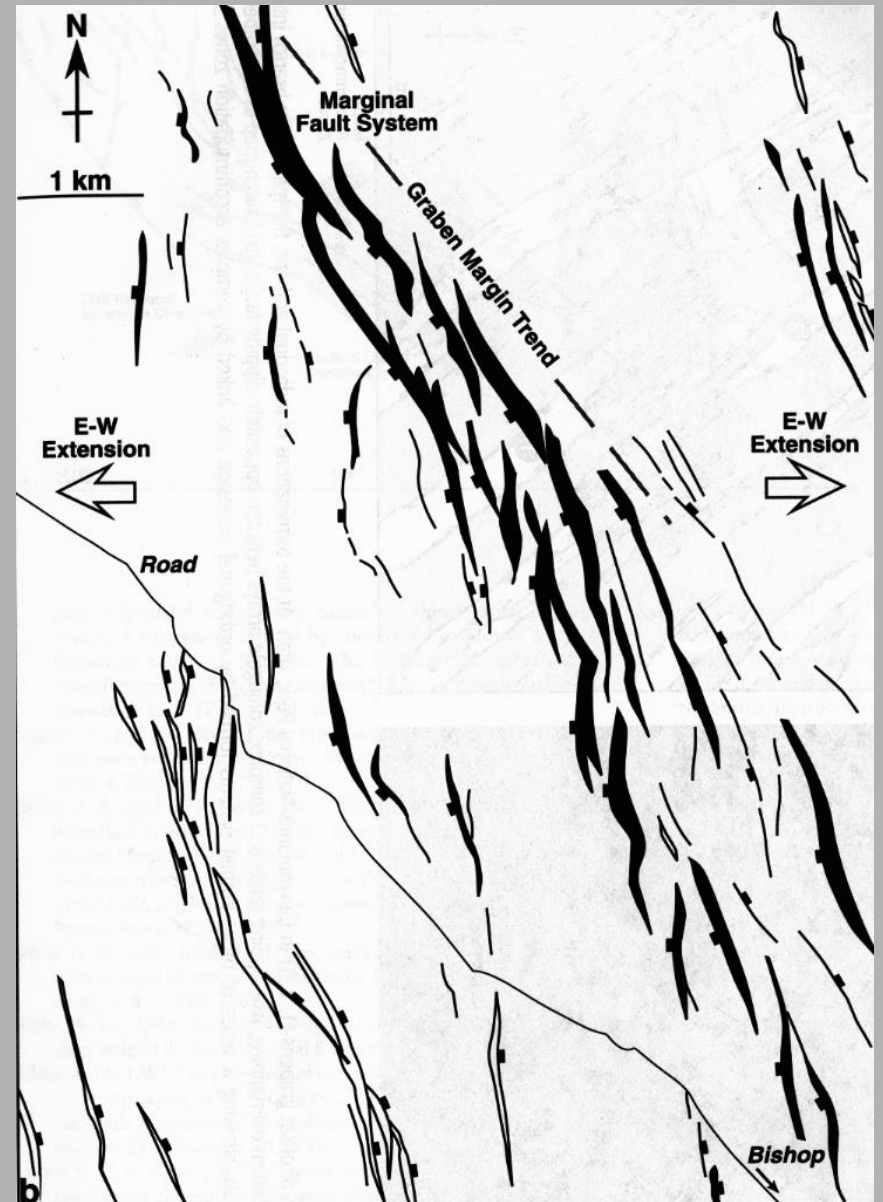
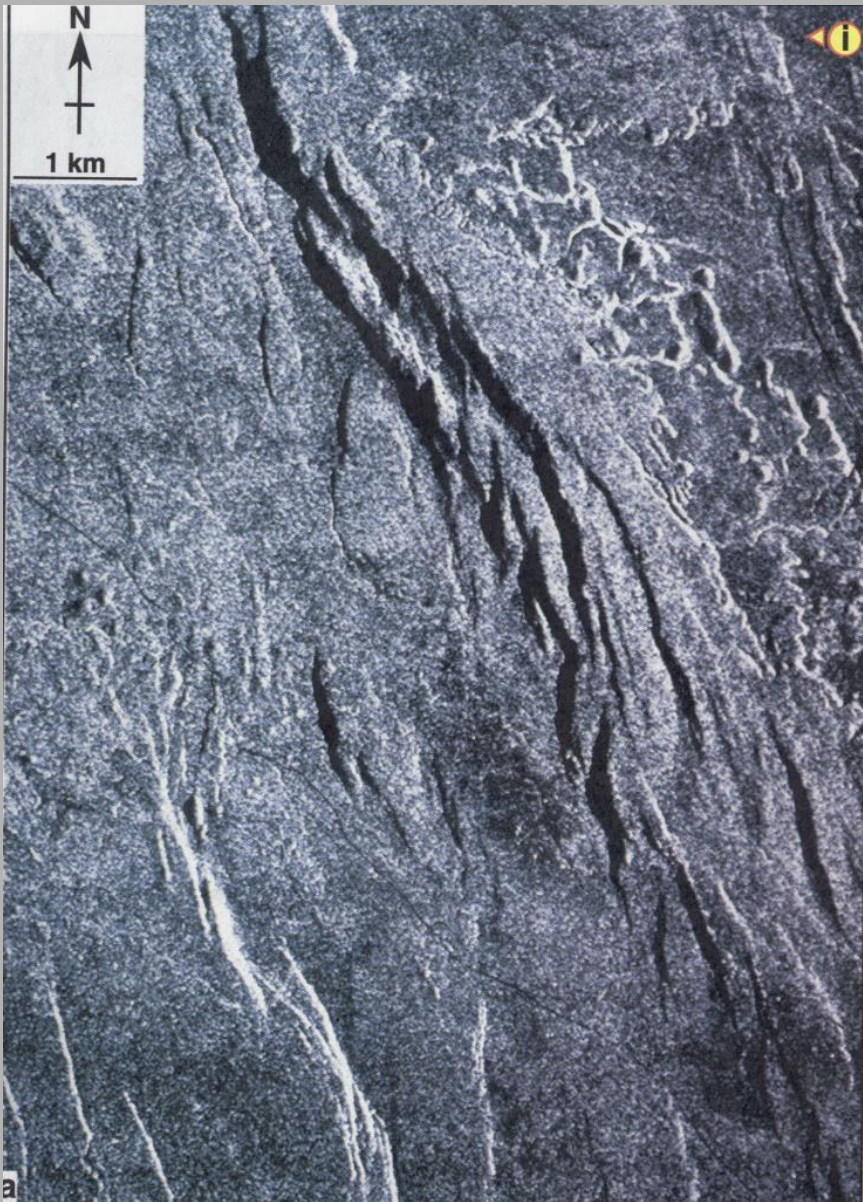


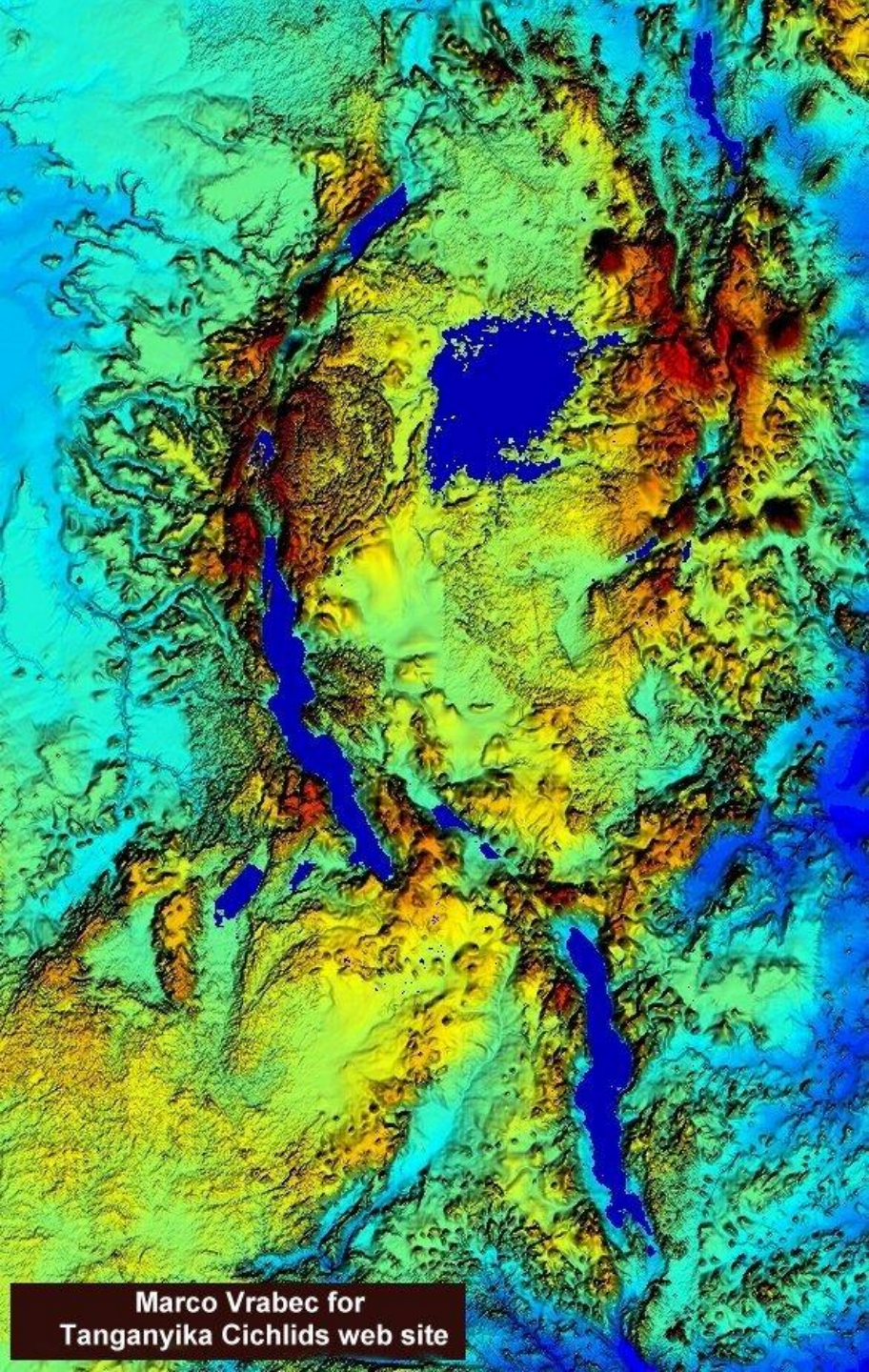


Gregory Rift, west of Nairobi, Kenya; McClay et al., 2002

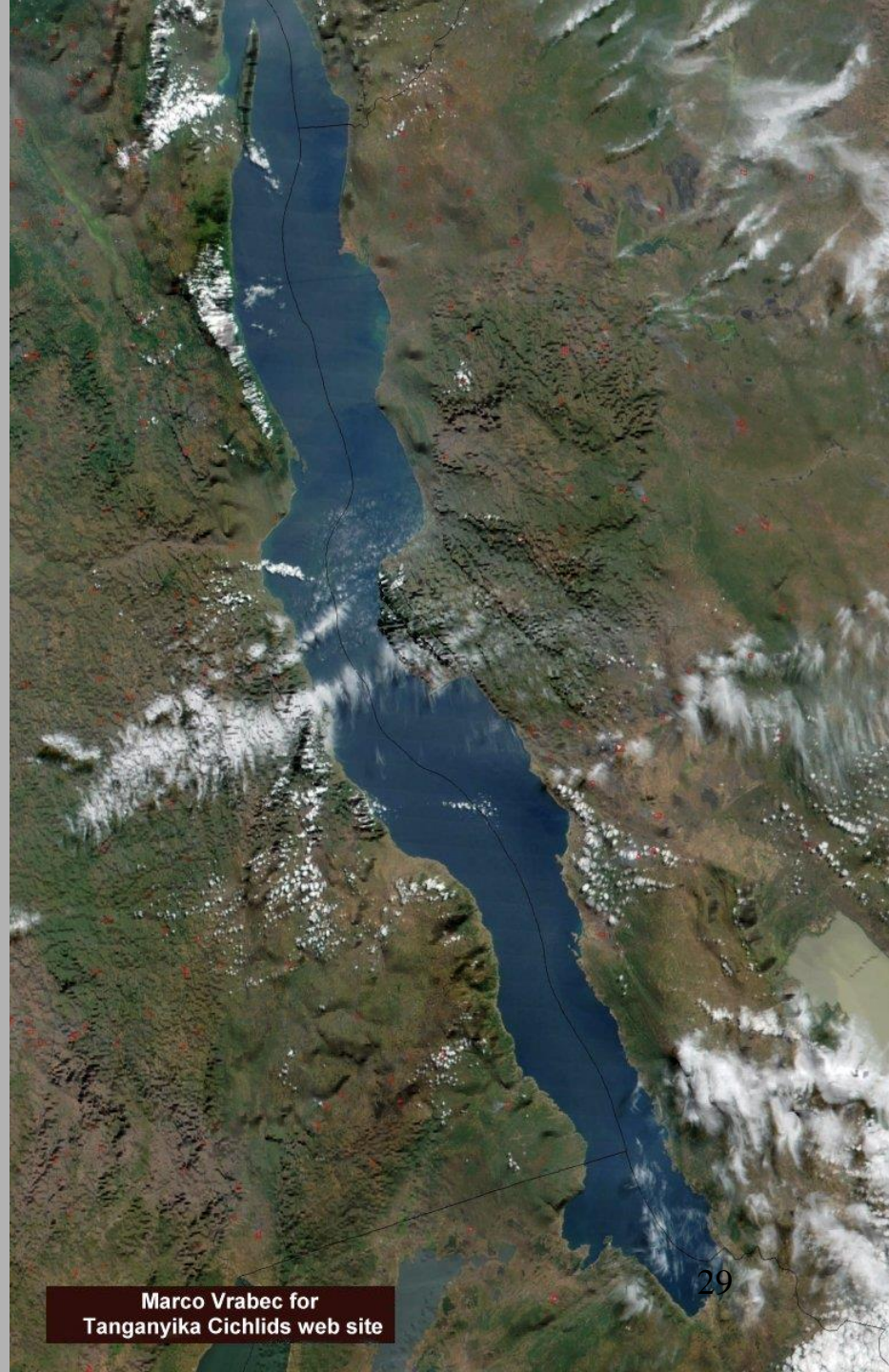




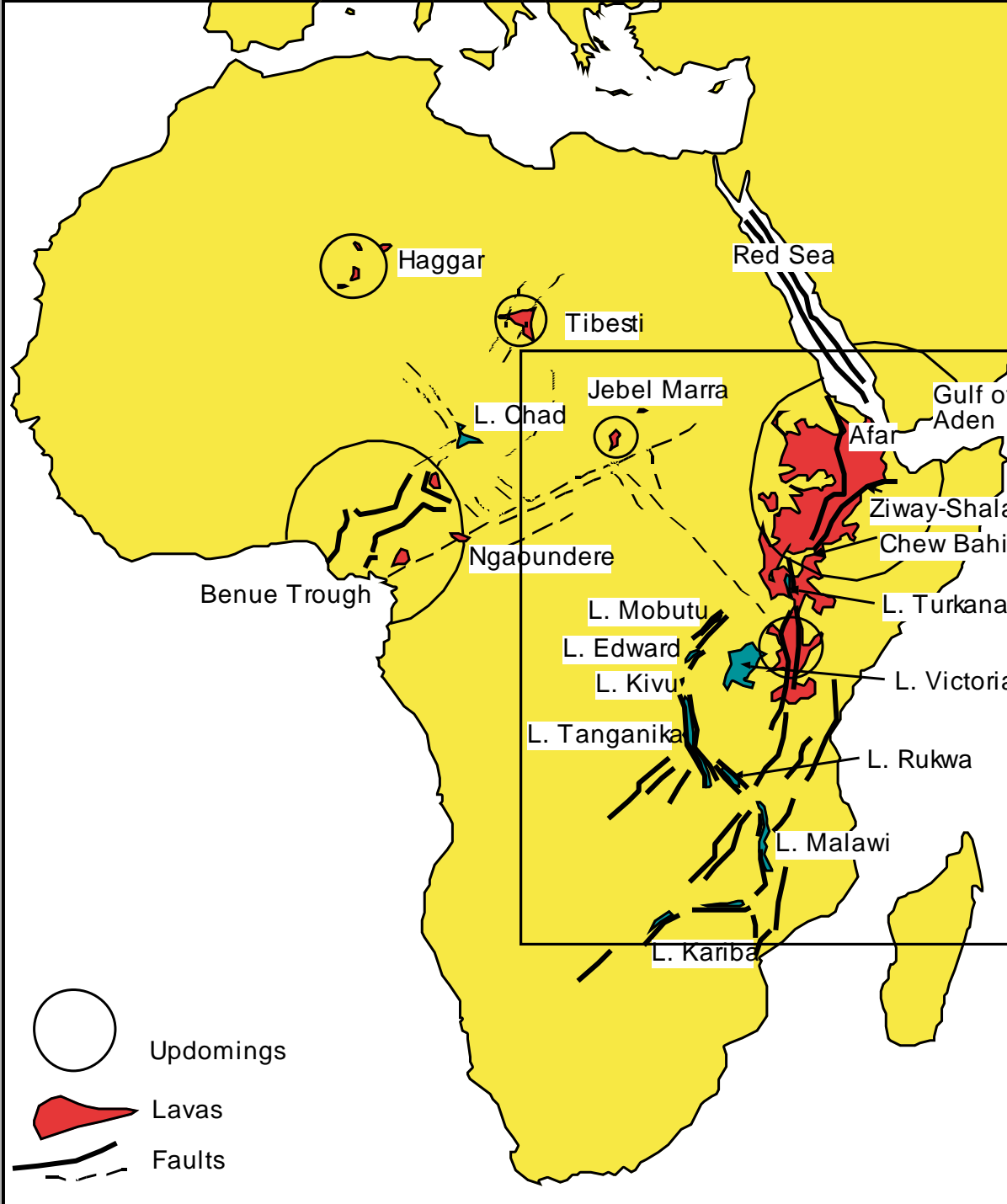


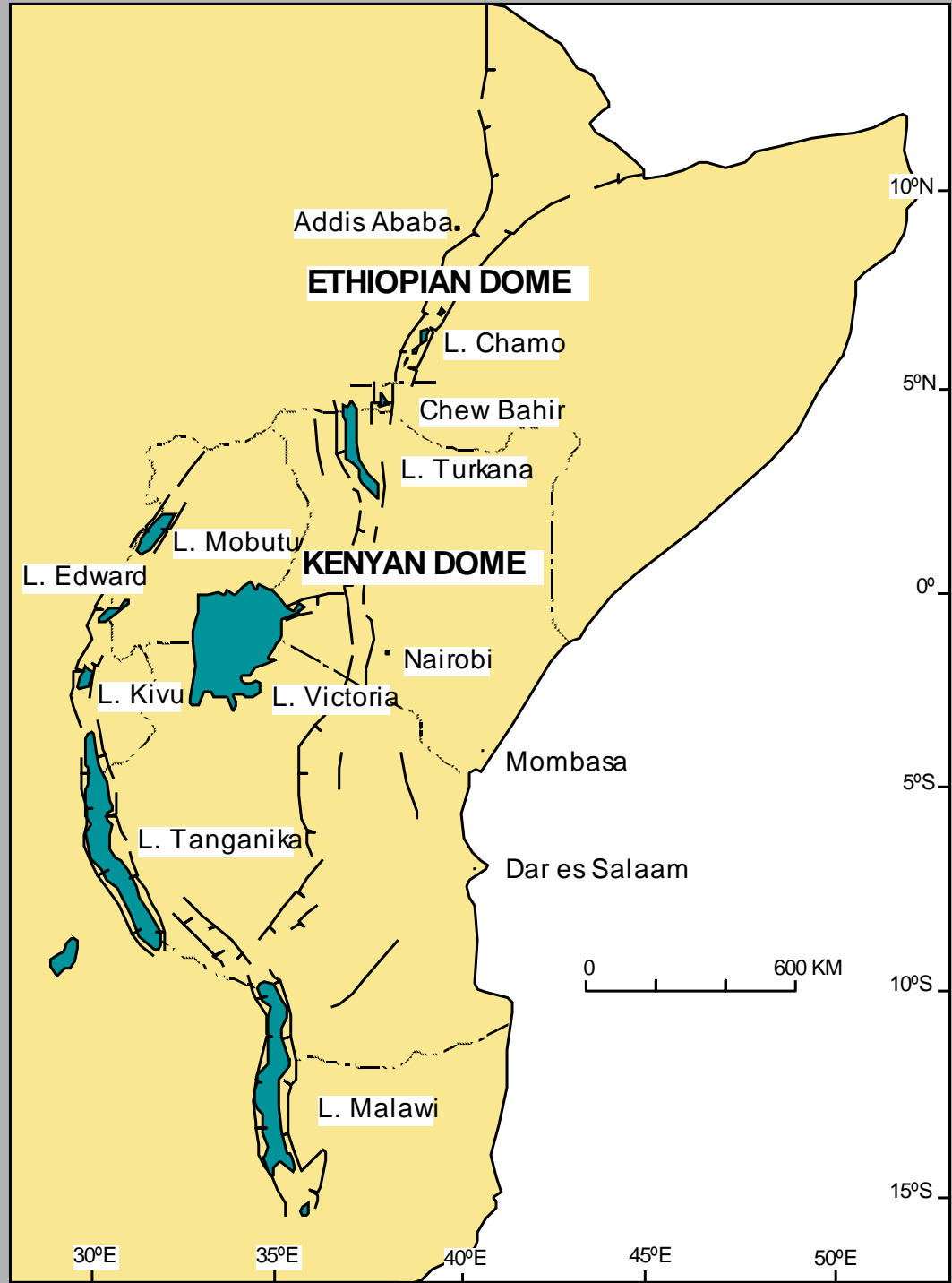


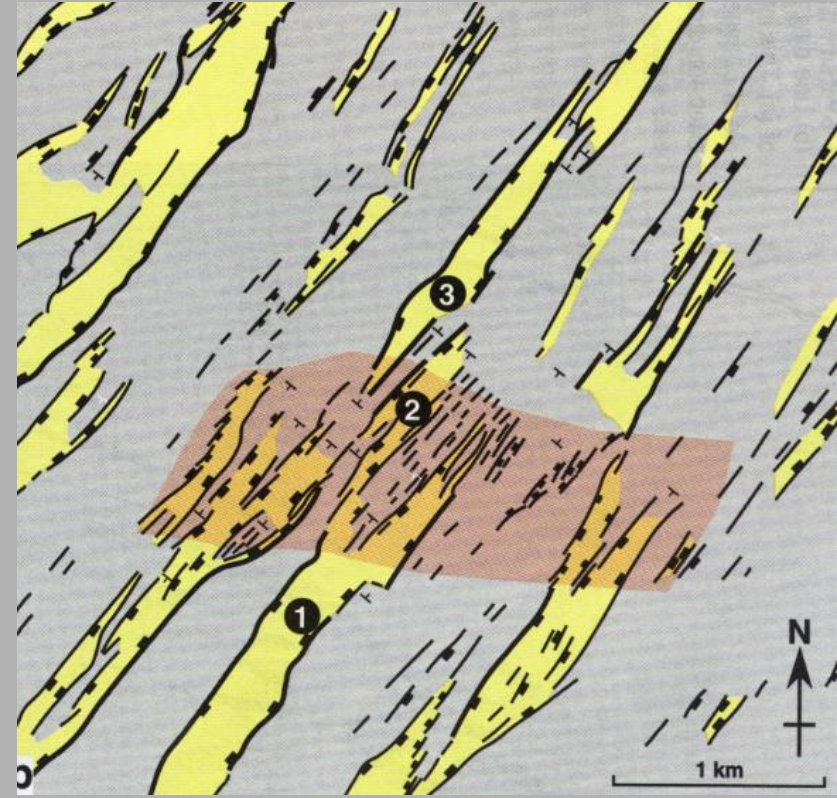
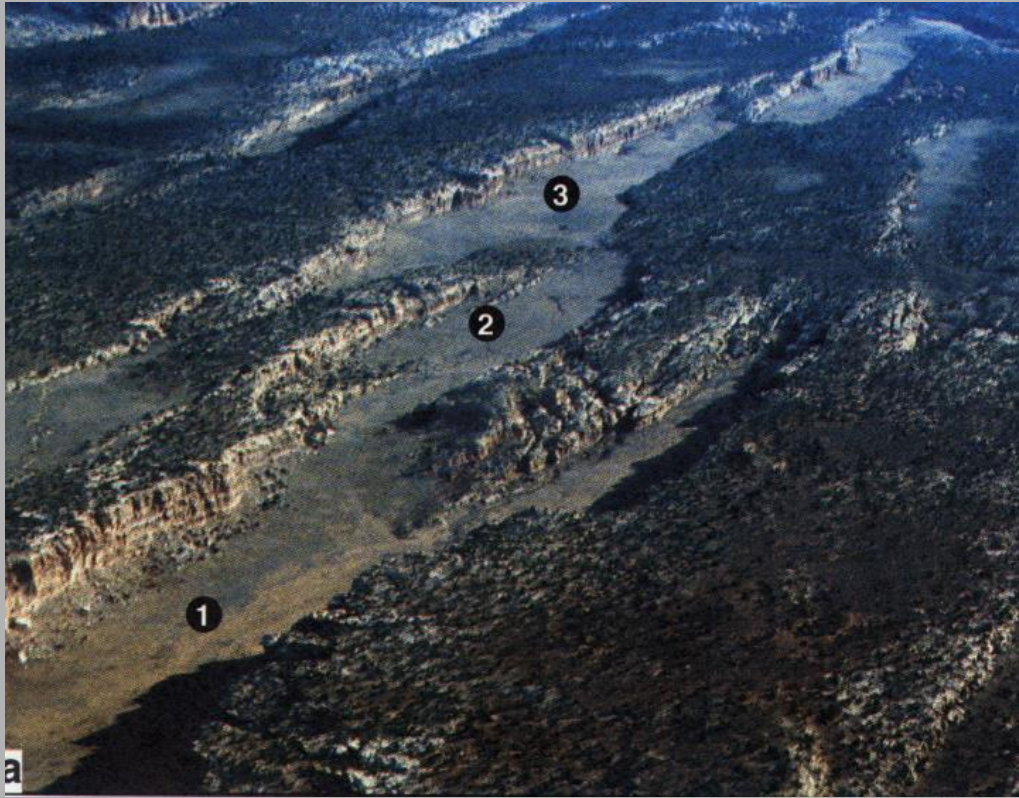
Marco Vrabec for
Tanganyika Cichlids web site



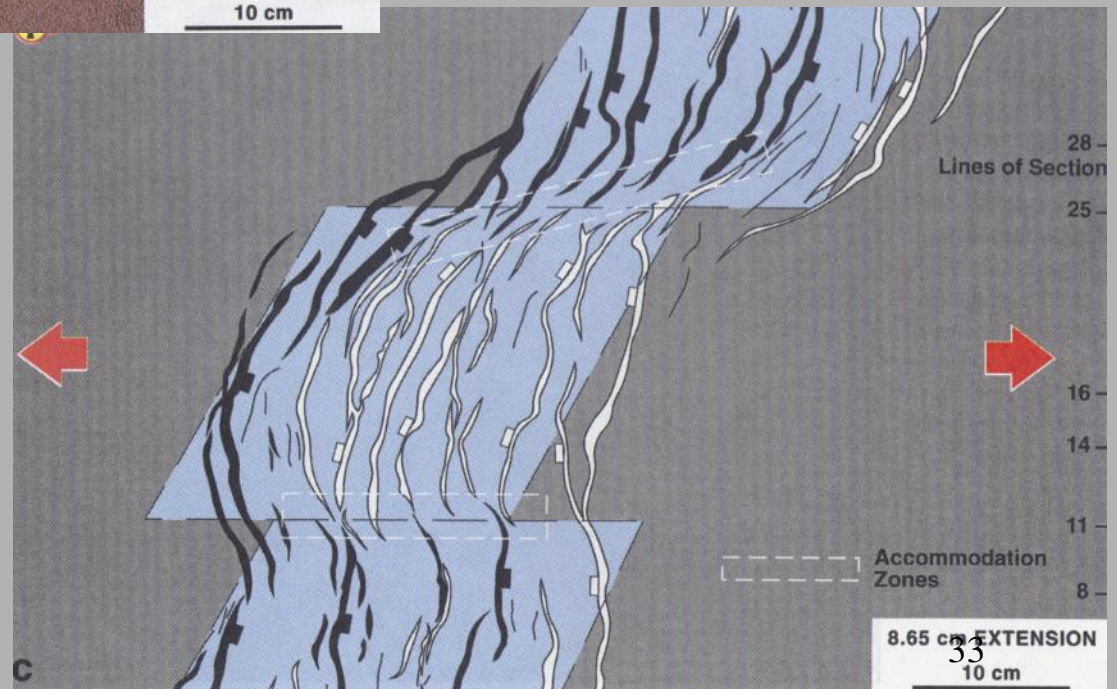
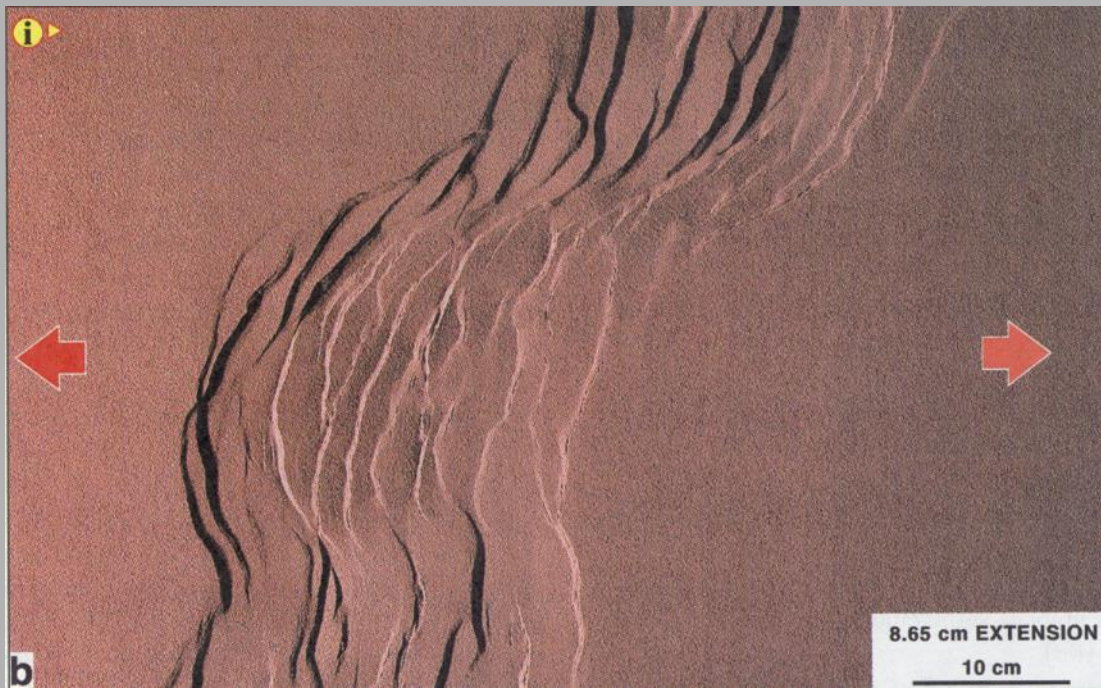
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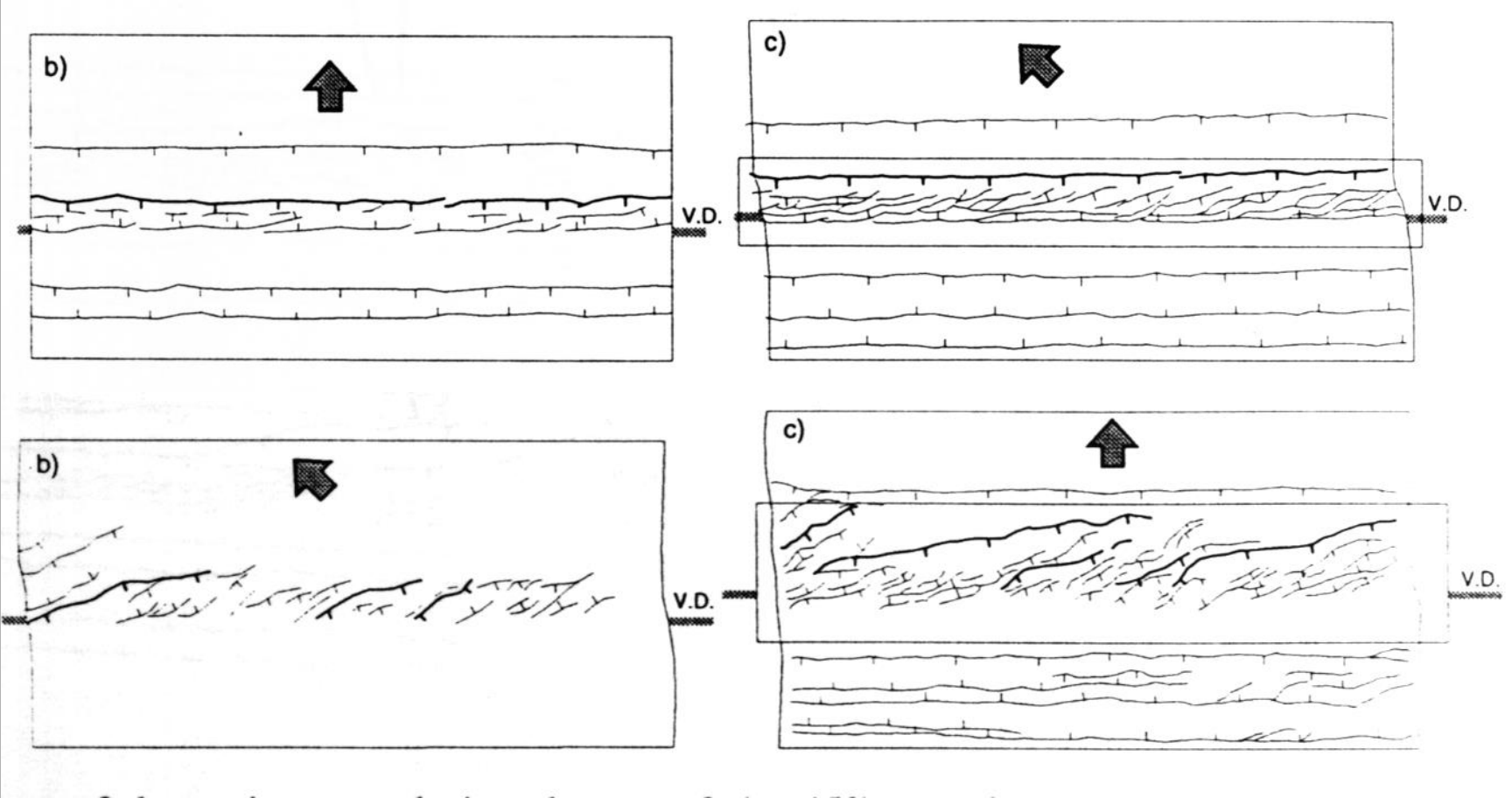






Canyonlands National Park, Utah; McClay et al., 2002





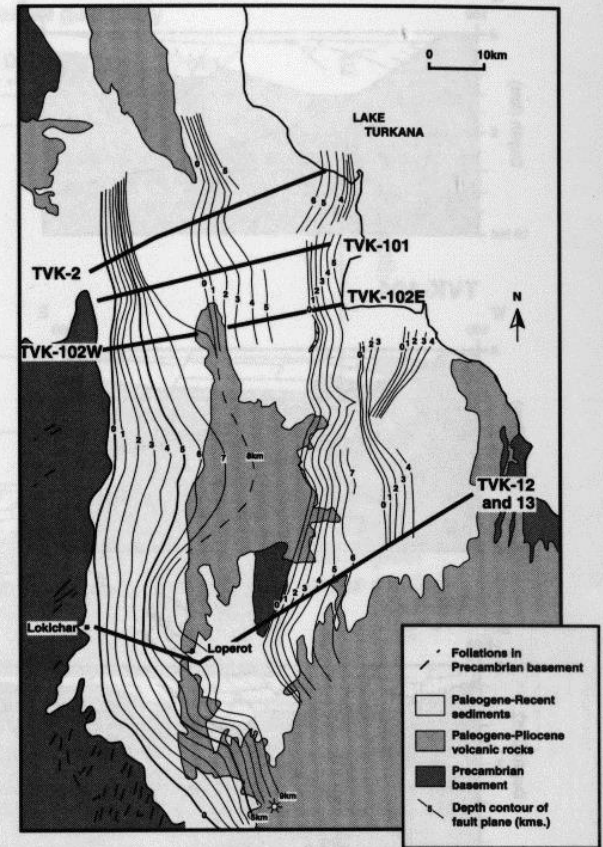
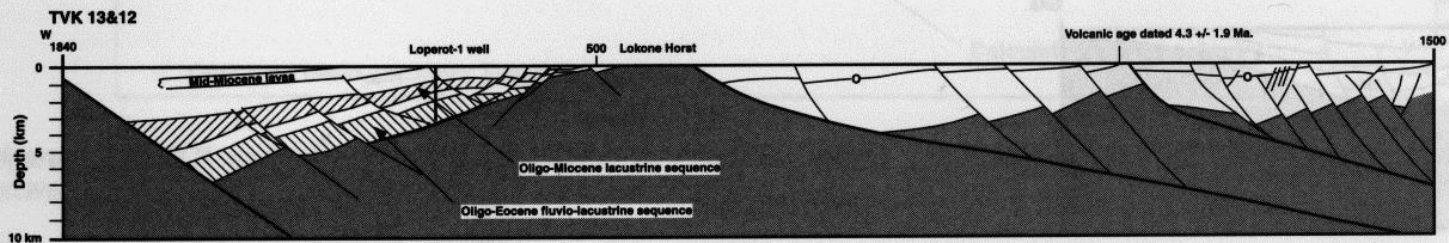
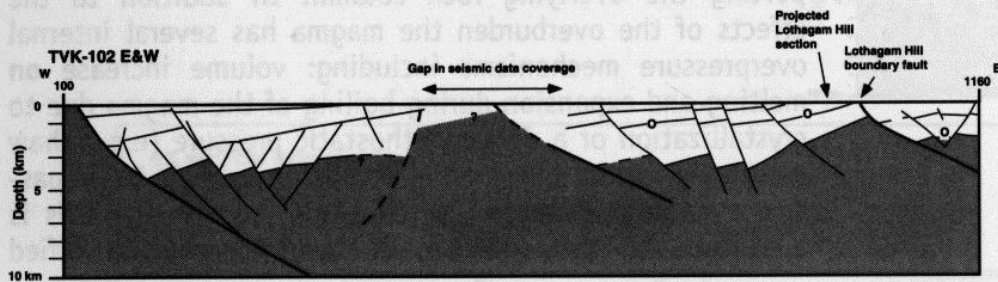
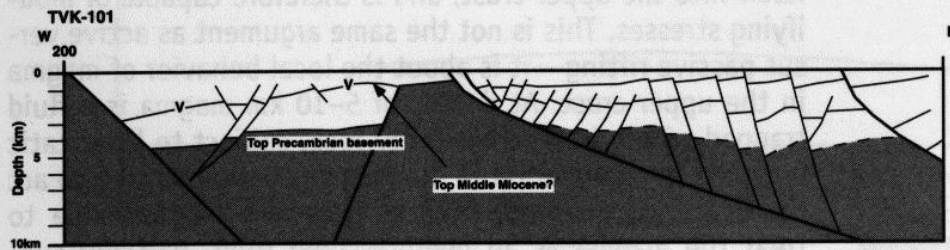
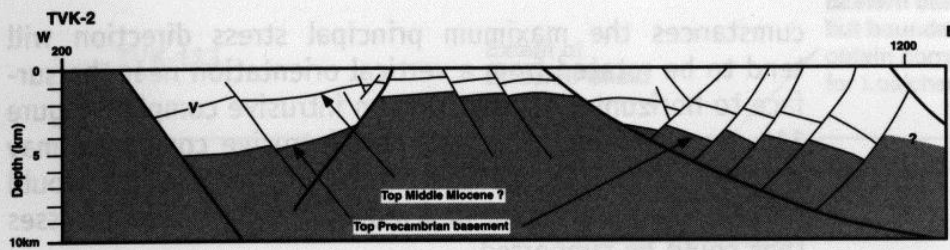
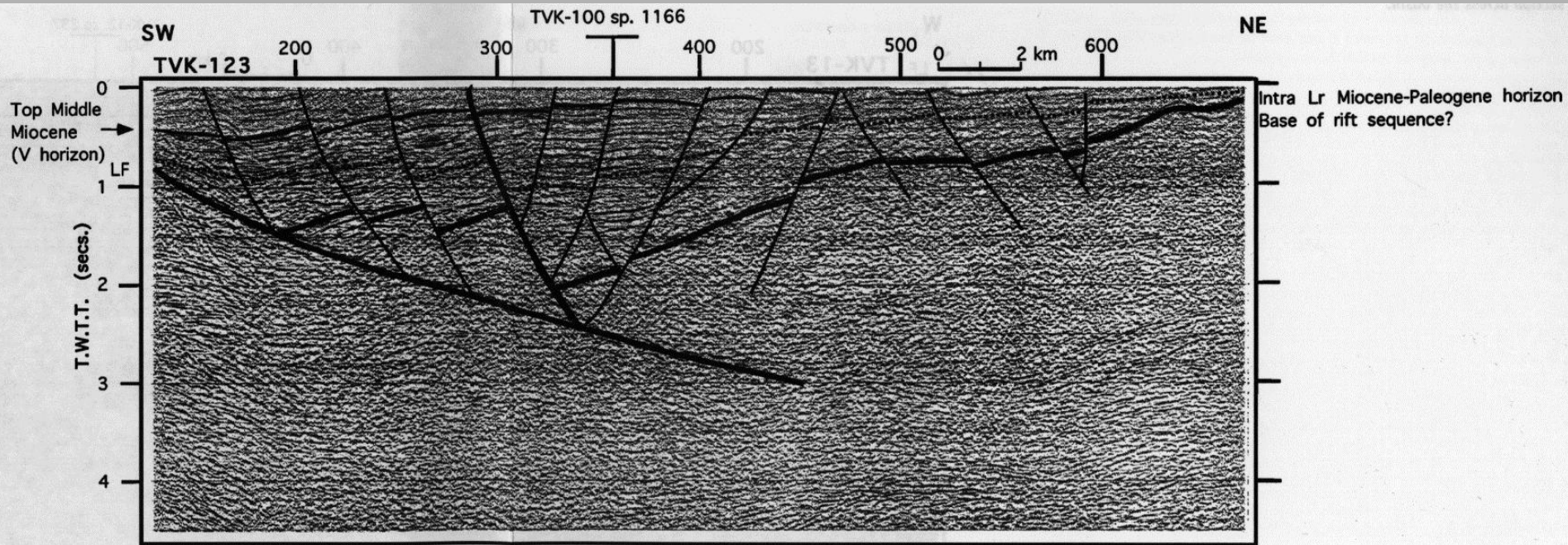
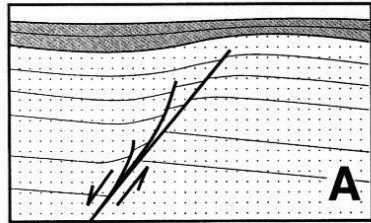


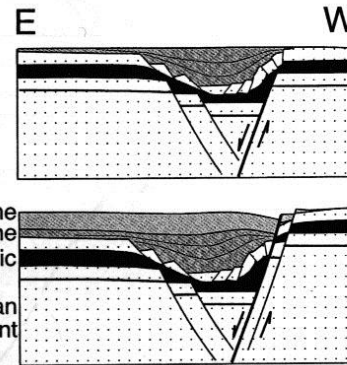
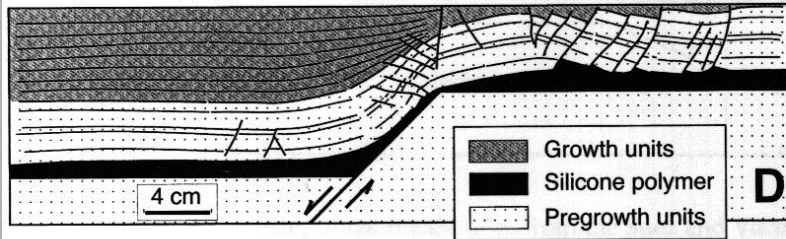
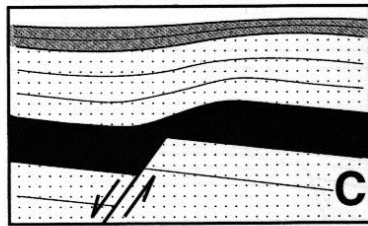
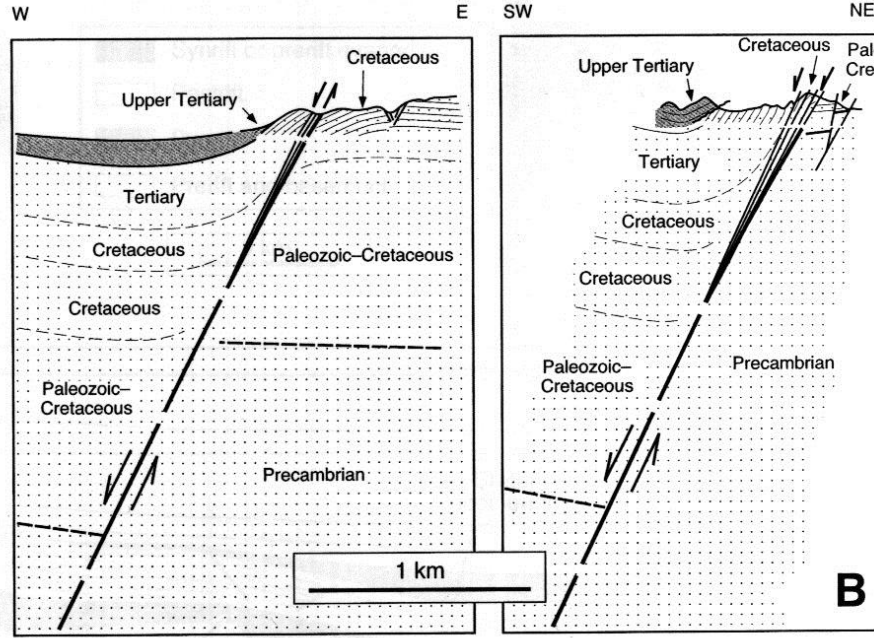
Figure 10. Regional cross sections through the Turkana area based on seismic reflection data.



růstové zlomy

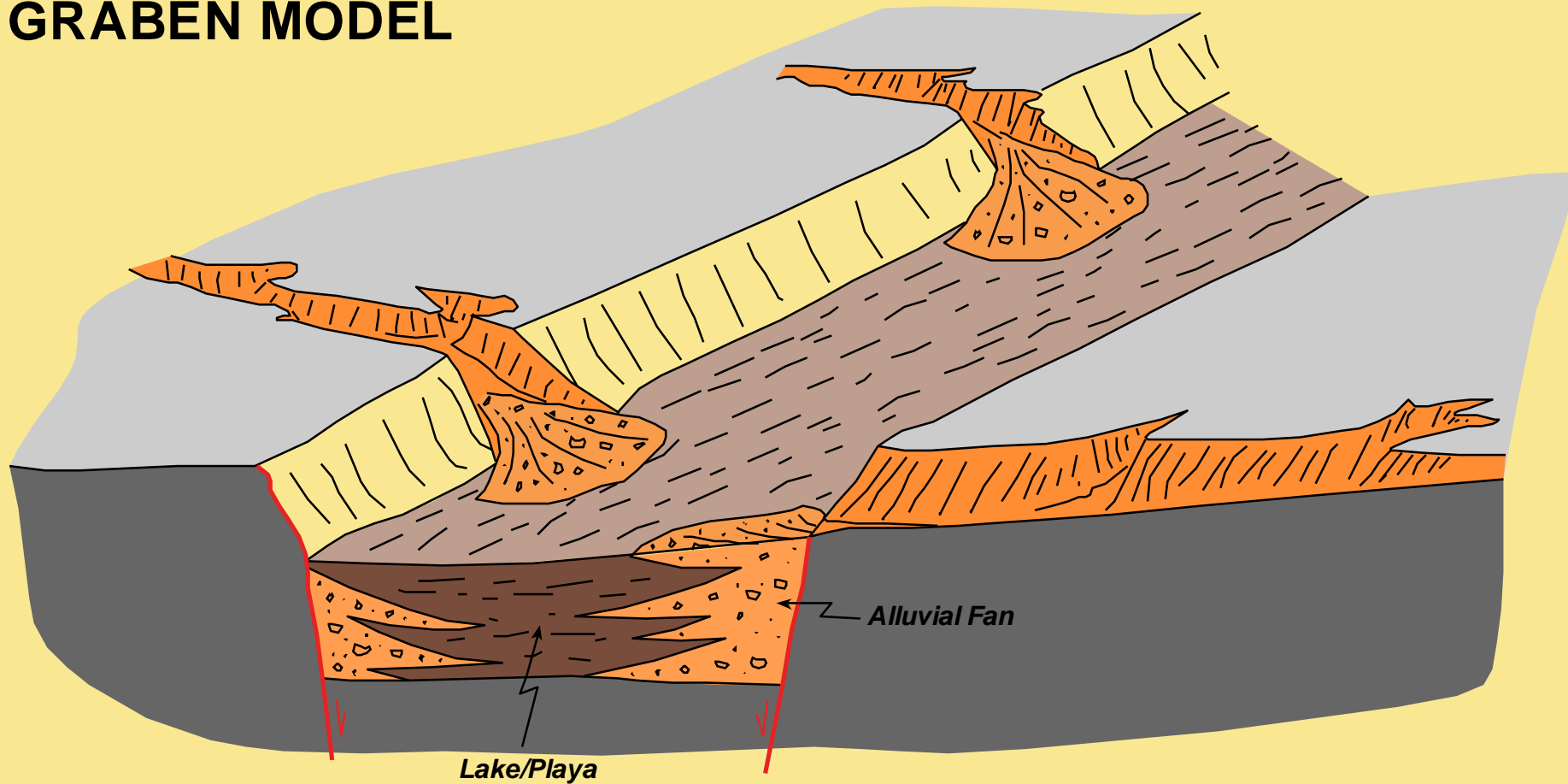


- Synrift or prerift evaporites
- Postrift
- Synrift
- Prerift and basement

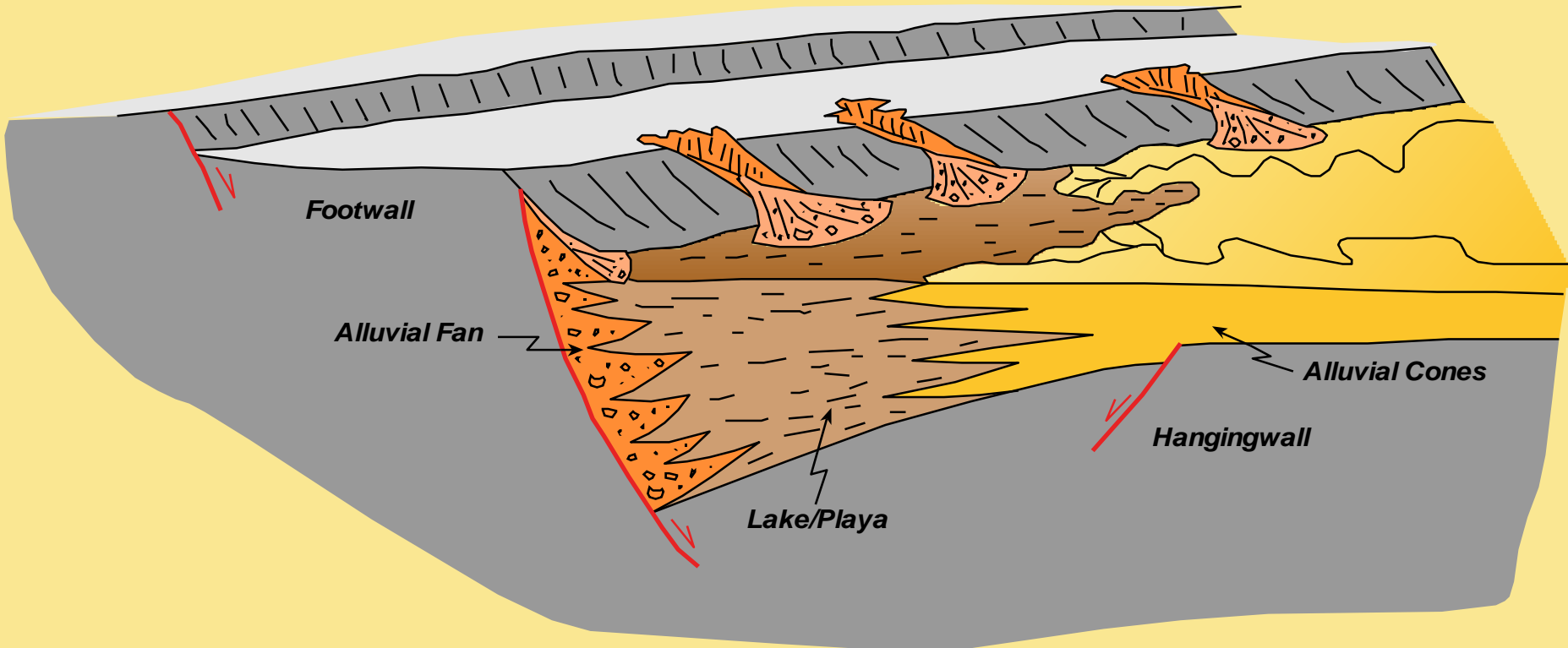


- Growth units
 - Silicone polymer
 - Pregrowth units
- ~ 5 km
No vertical exaggeration

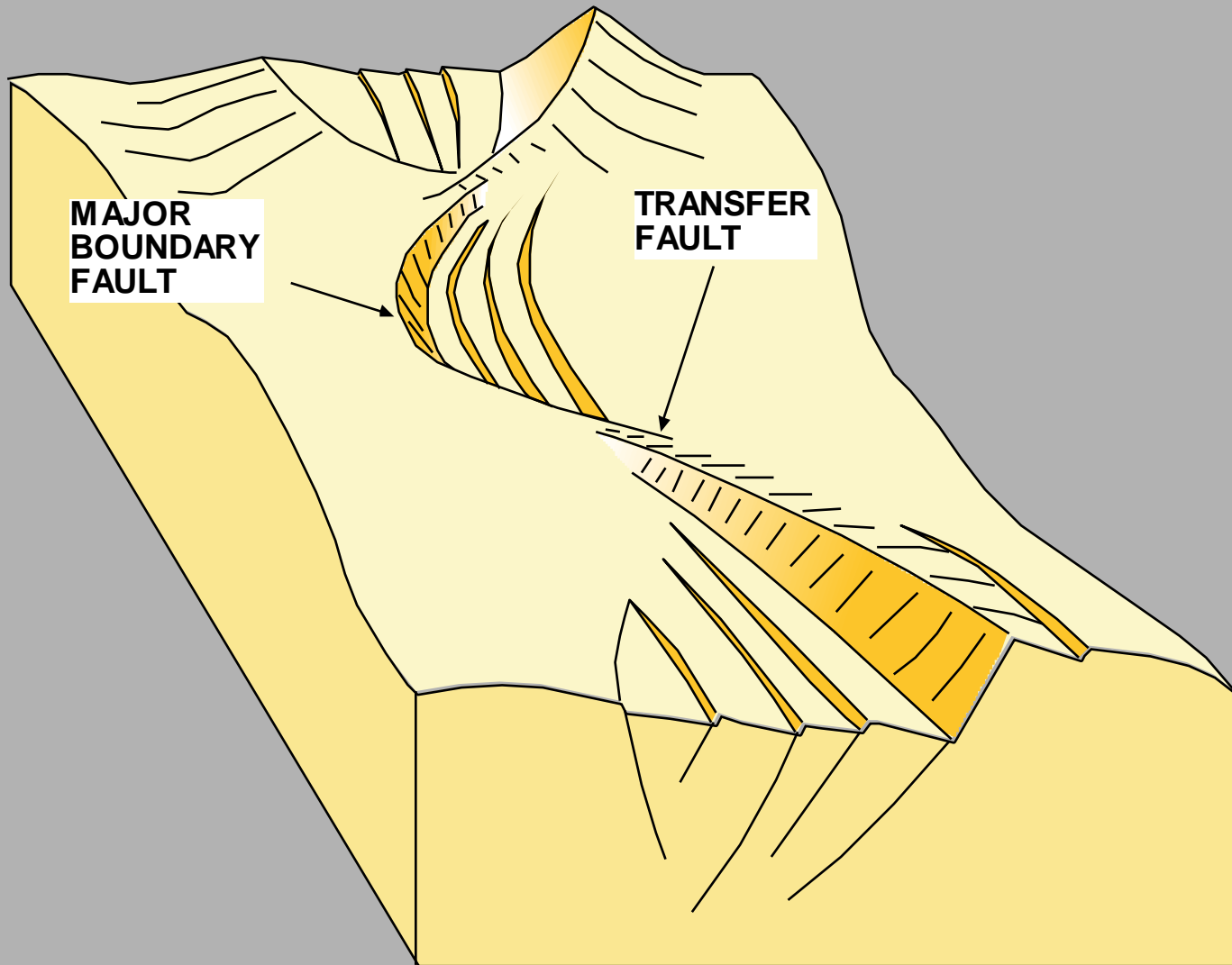
GRABEN MODEL



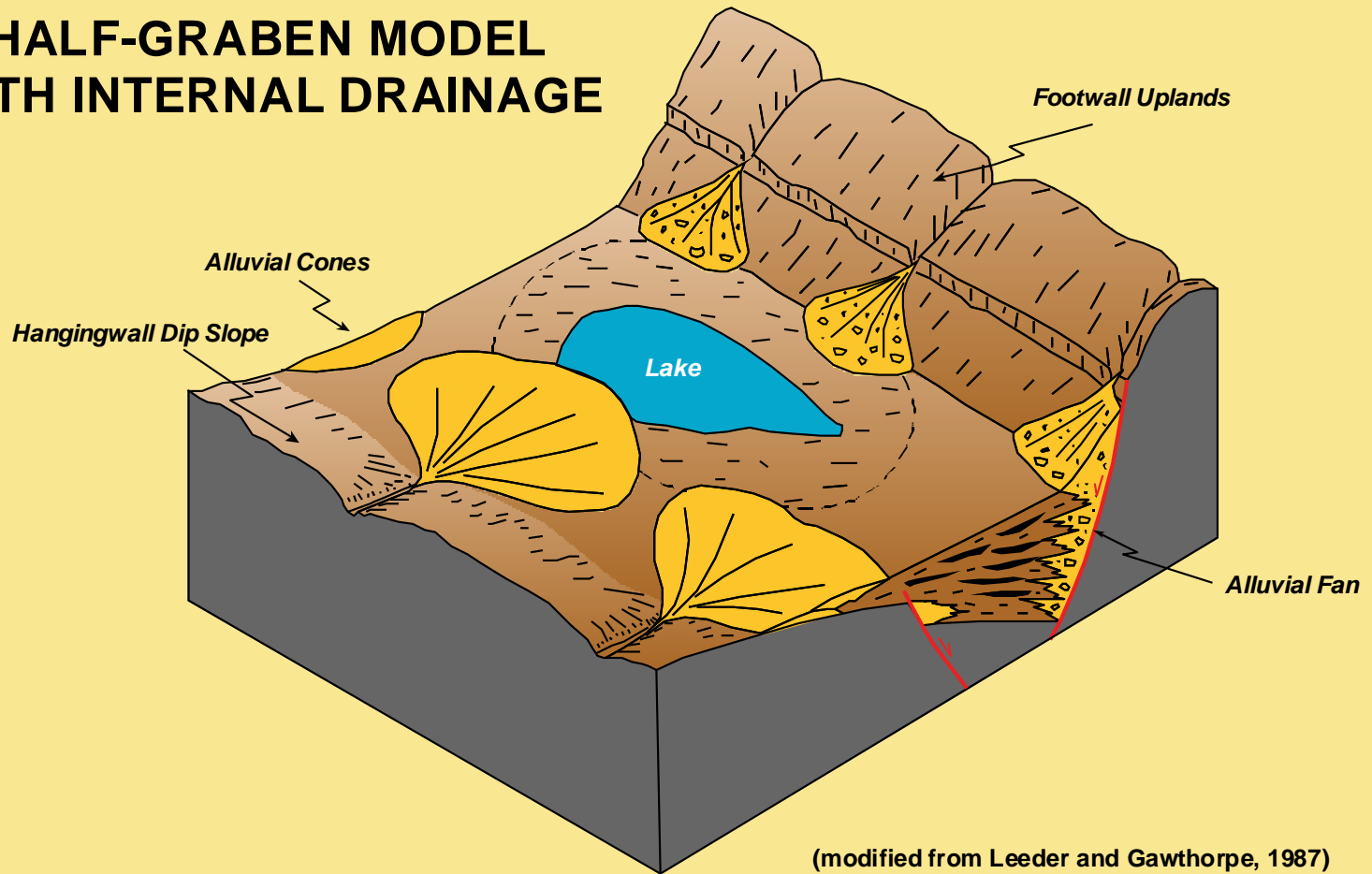
HALF GRABEN MODEL

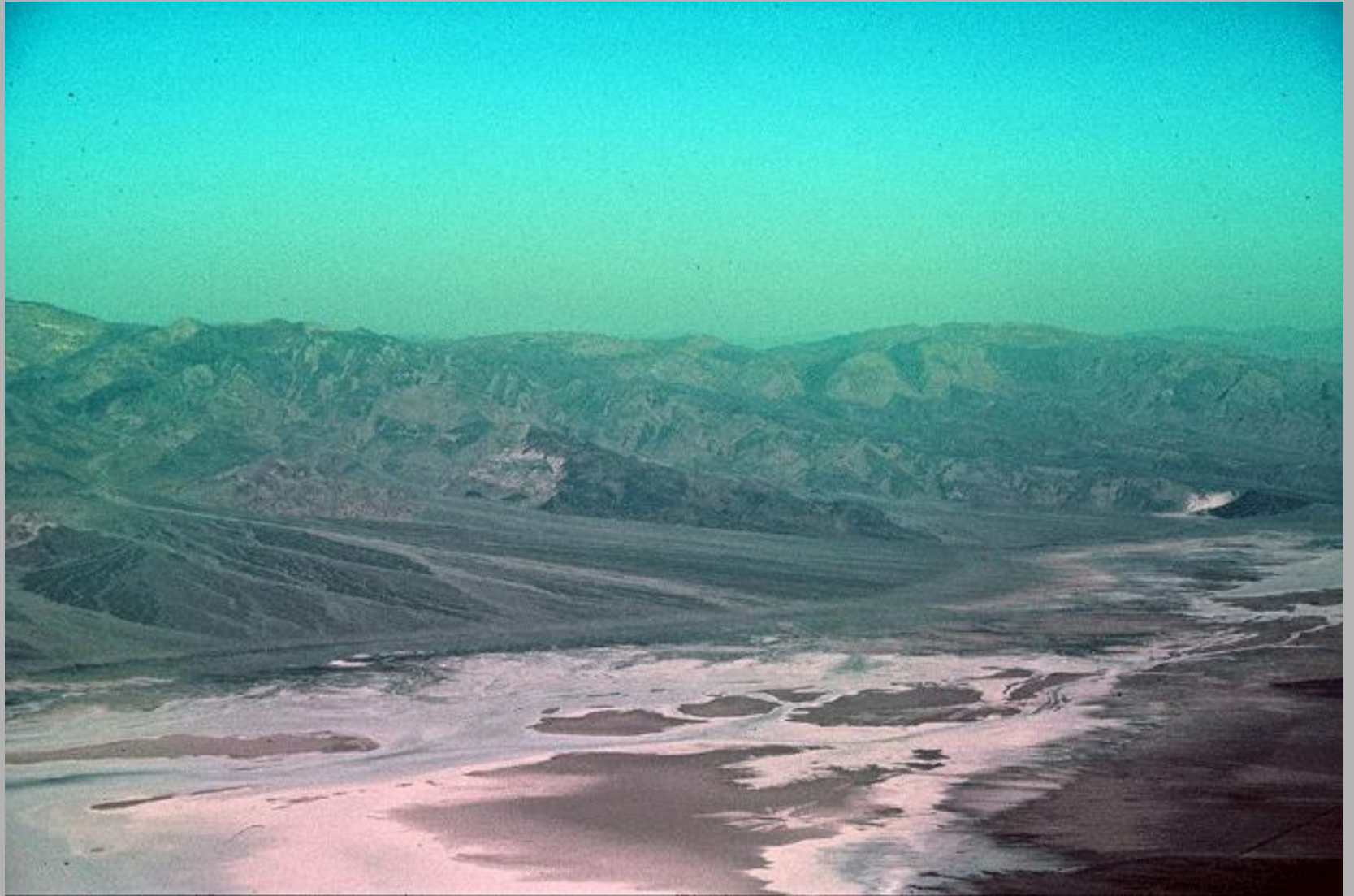


(modified from Forstick and Reid, 1981)



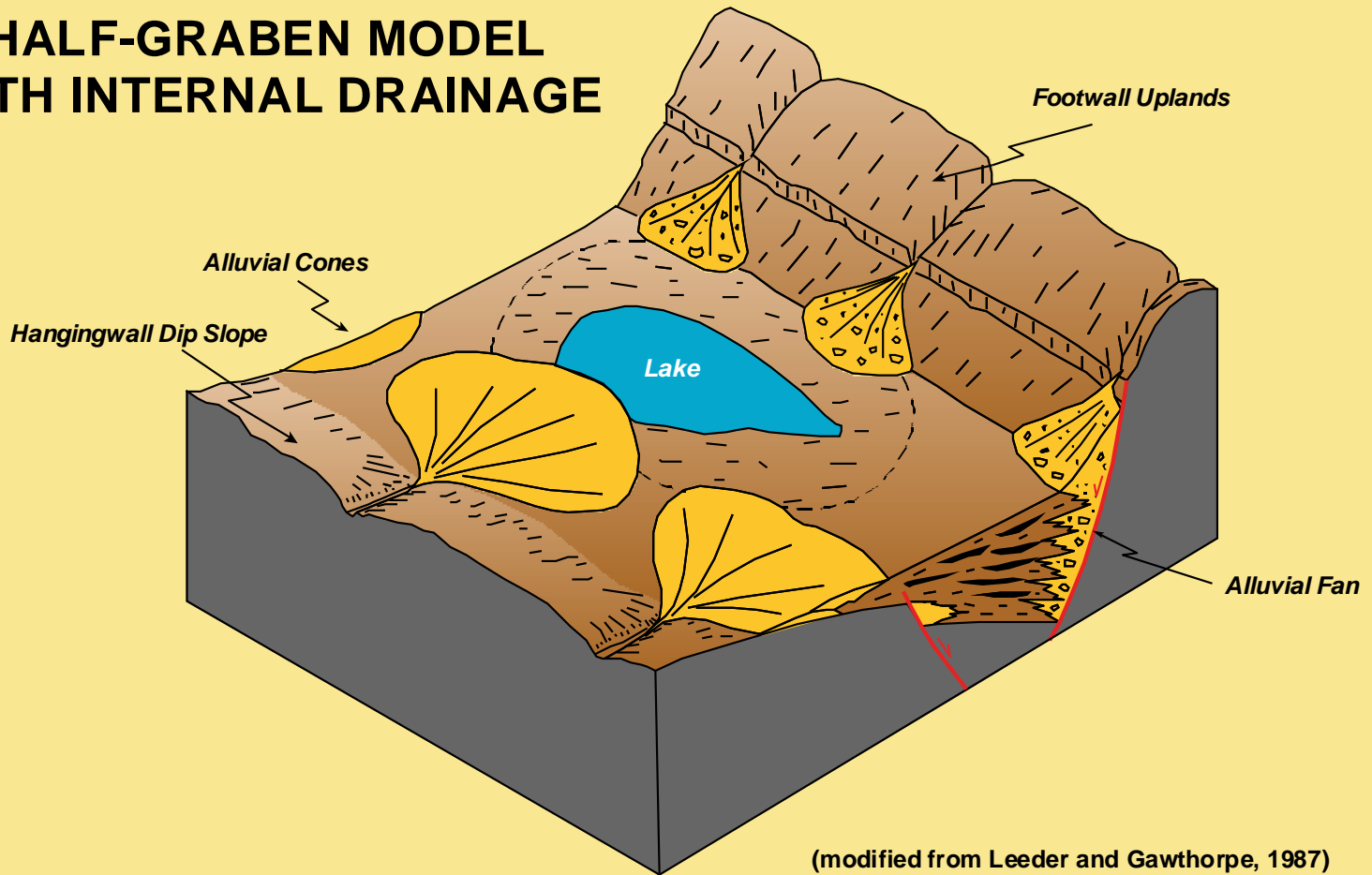
CONTINENTAL HALF-GRABEN MODEL WITH INTERNAL DRAINAGE



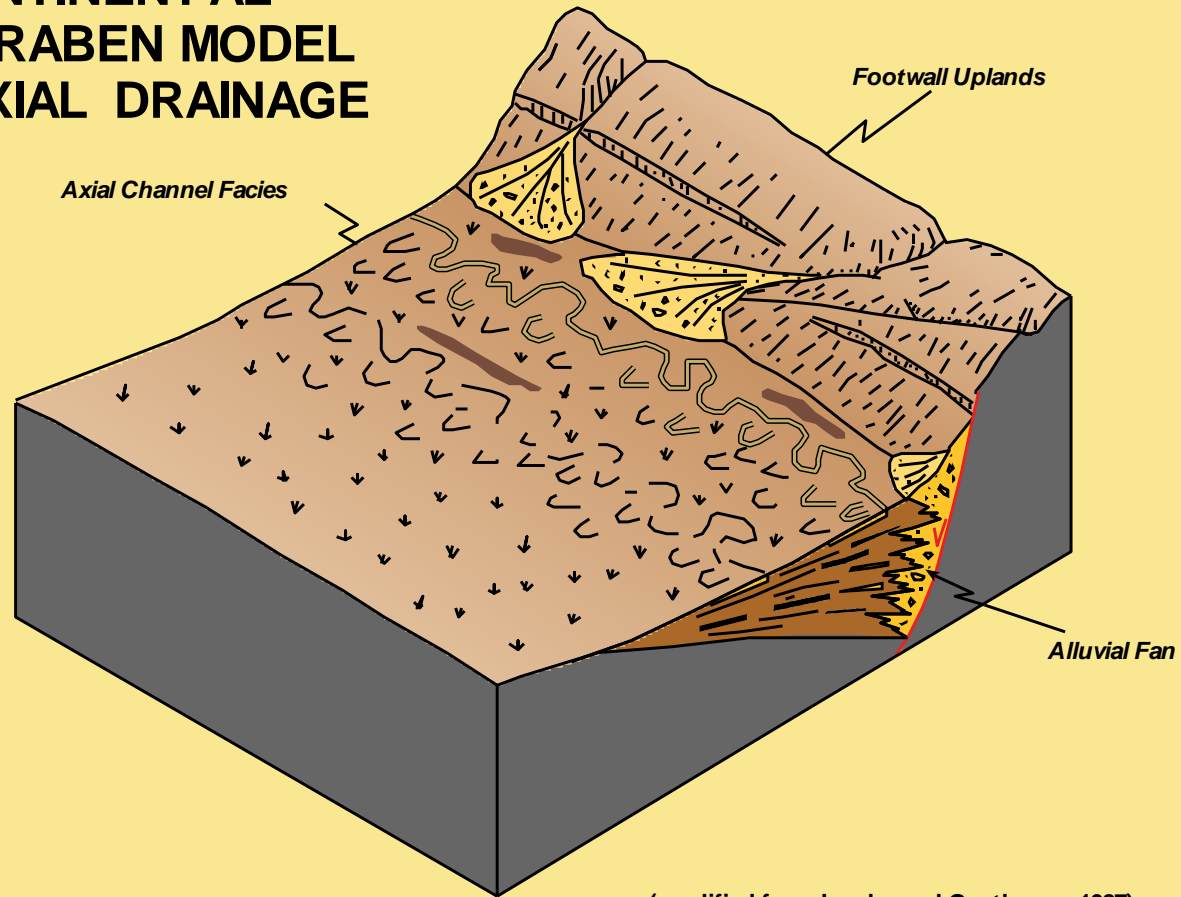




CONTINENTAL HALF-GRABEN MODEL WITH INTERNAL DRAINAGE



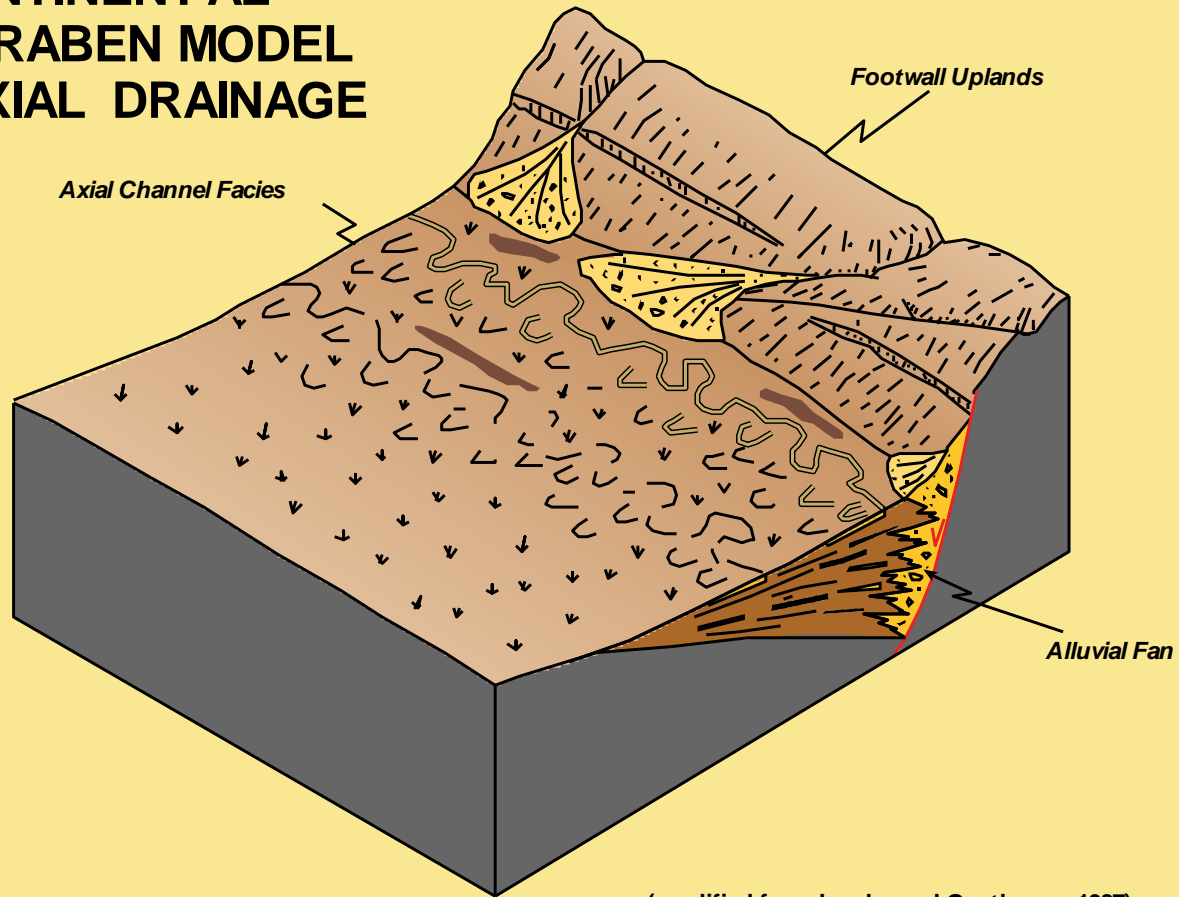
CONTINENTAL HALF-GRABEN MODEL WITH AXIAL DRAINAGE



(modified from Leeder and Gawthorpe, 1987)

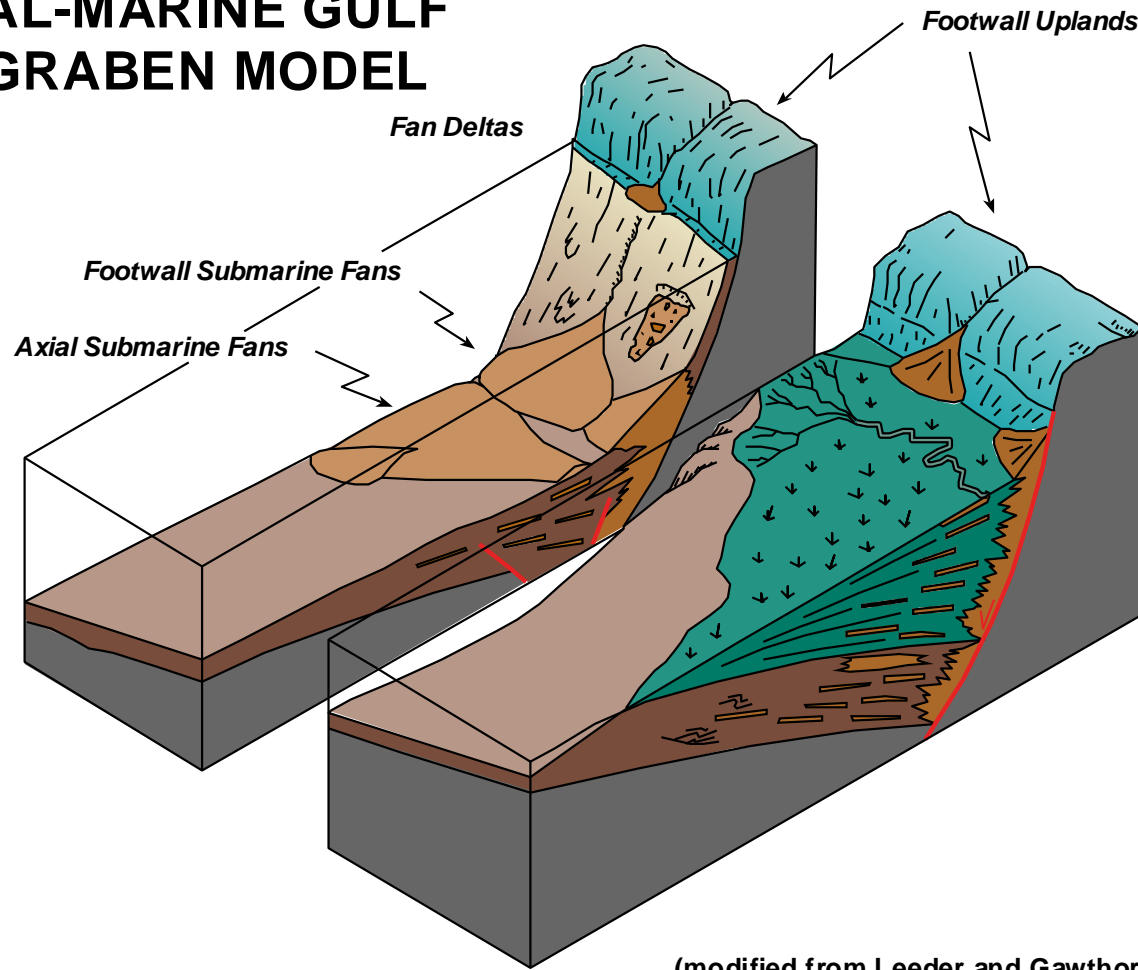


CONTINENTAL HALF-GRABEN MODEL WITH AXIAL DRAINAGE



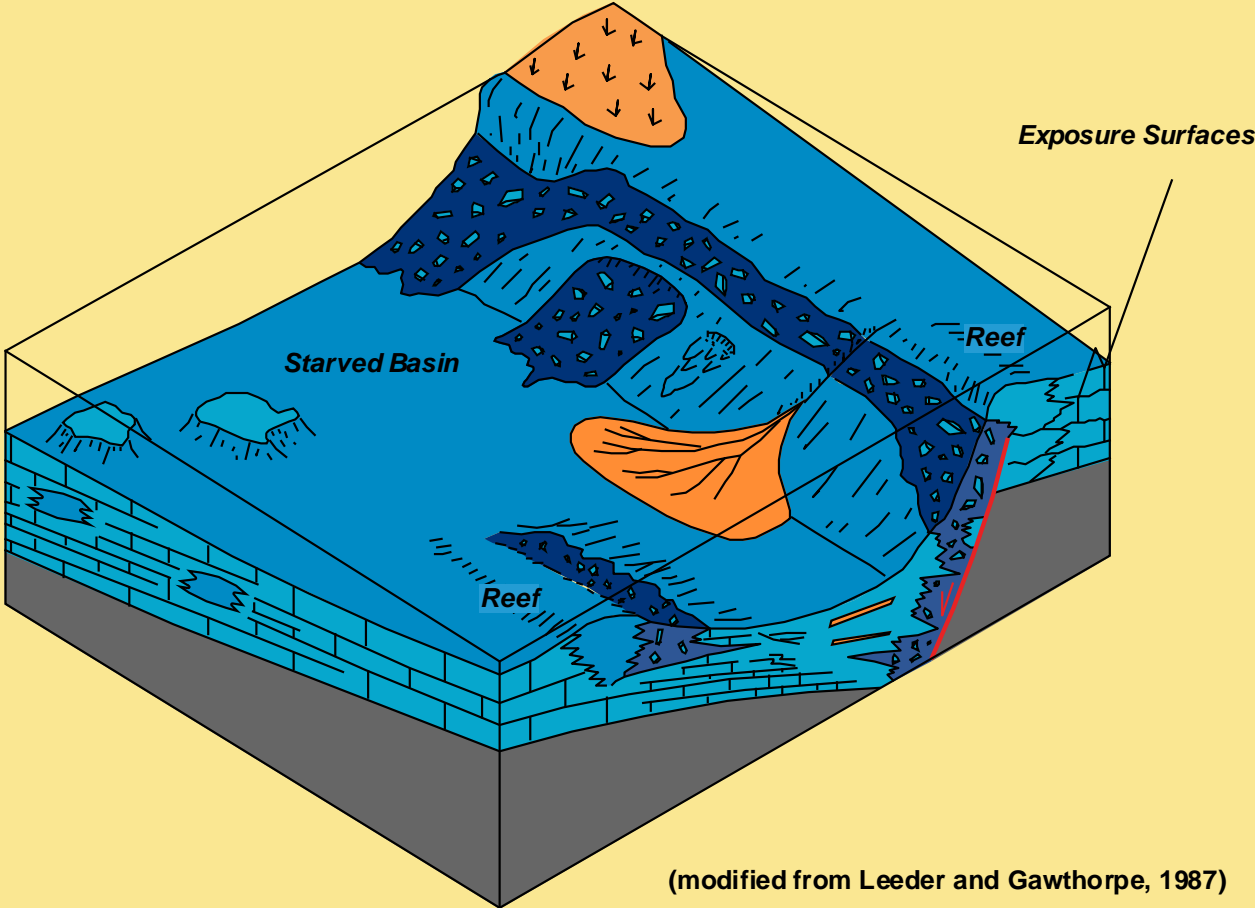
(modified from Leeder and Gawthorpe, 1987)

COASTAL-MARINE GULF HALF-GRABEN MODEL

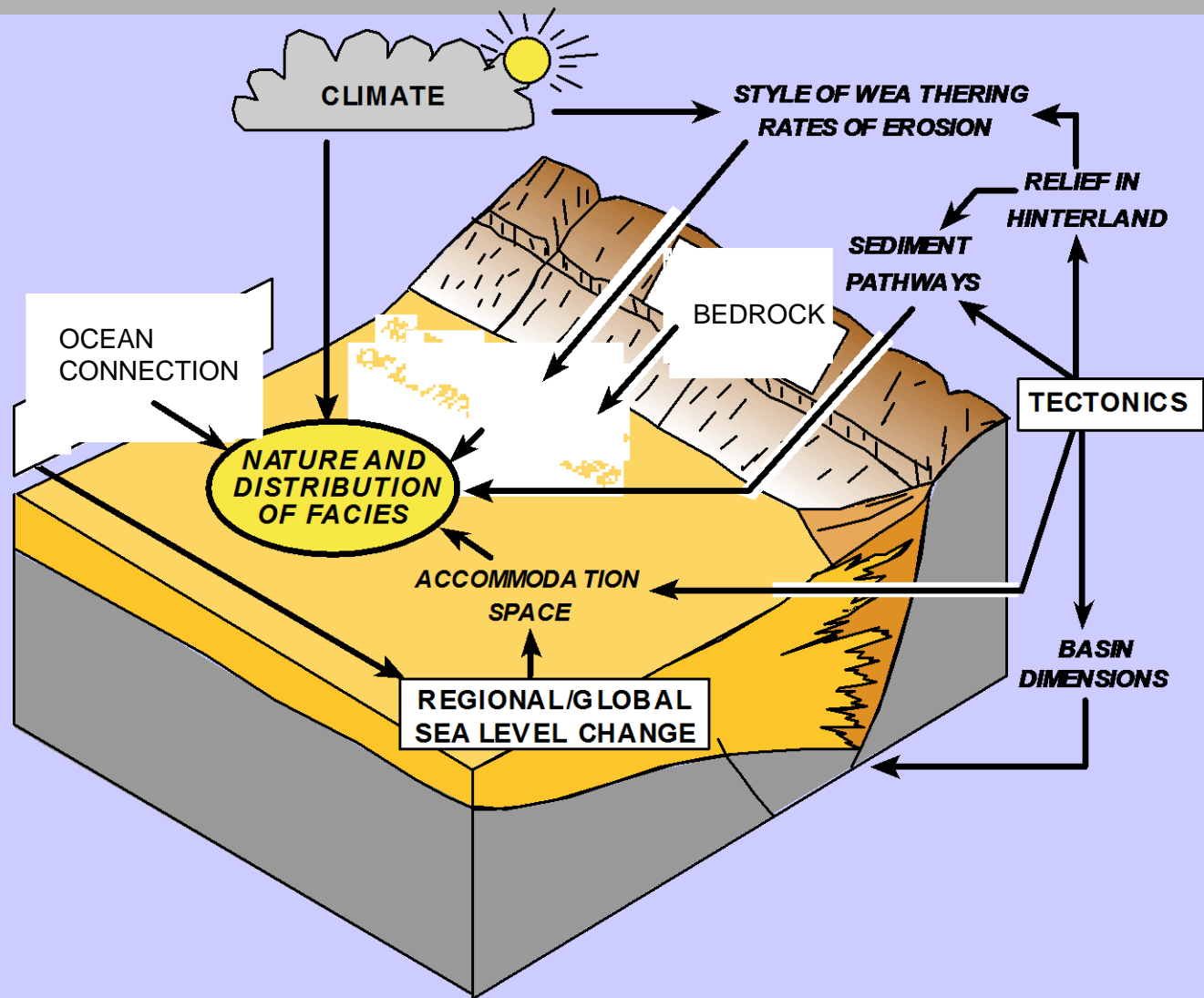


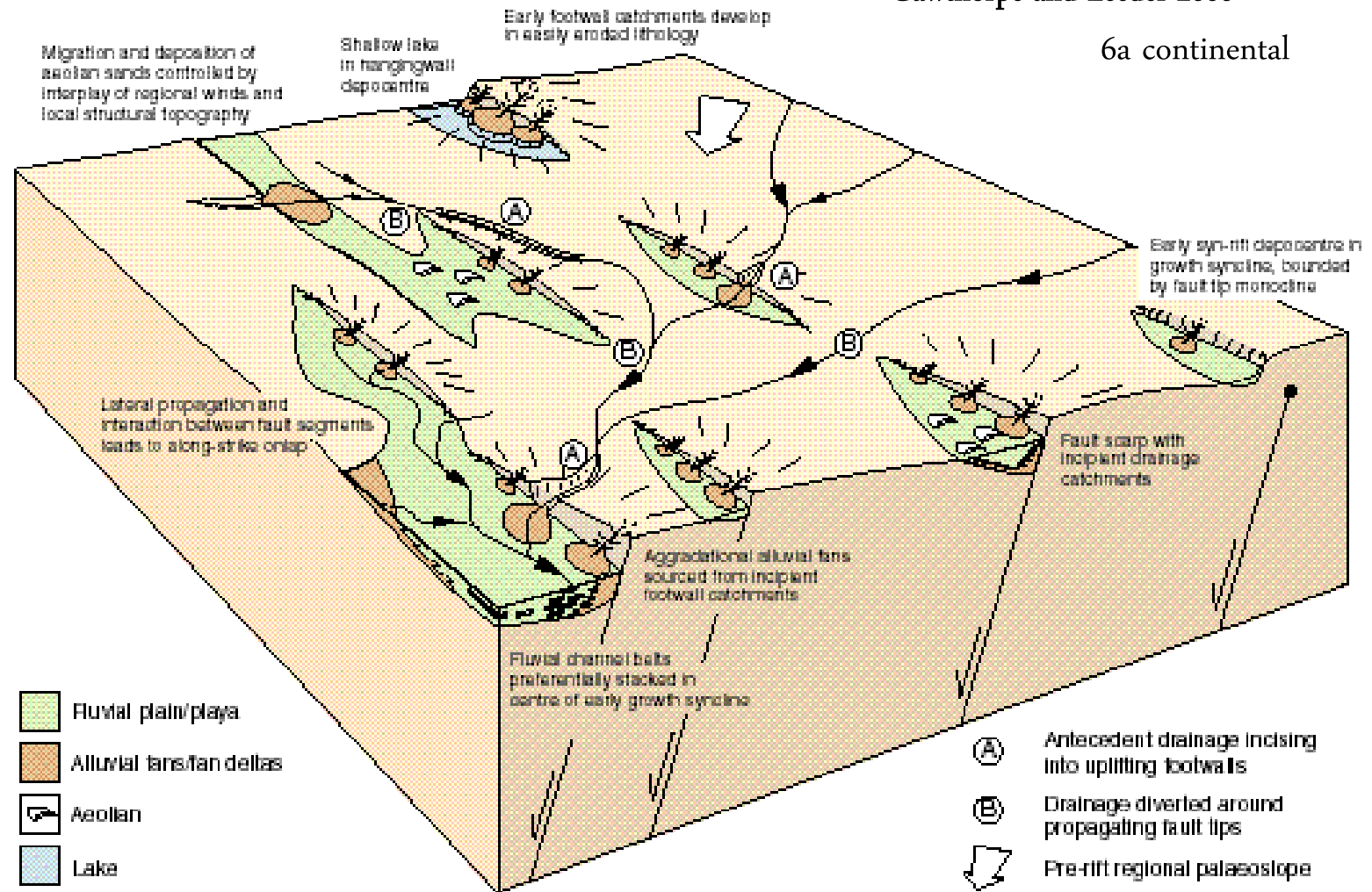
(modified from Leeder and Gawthorpe, 1987)

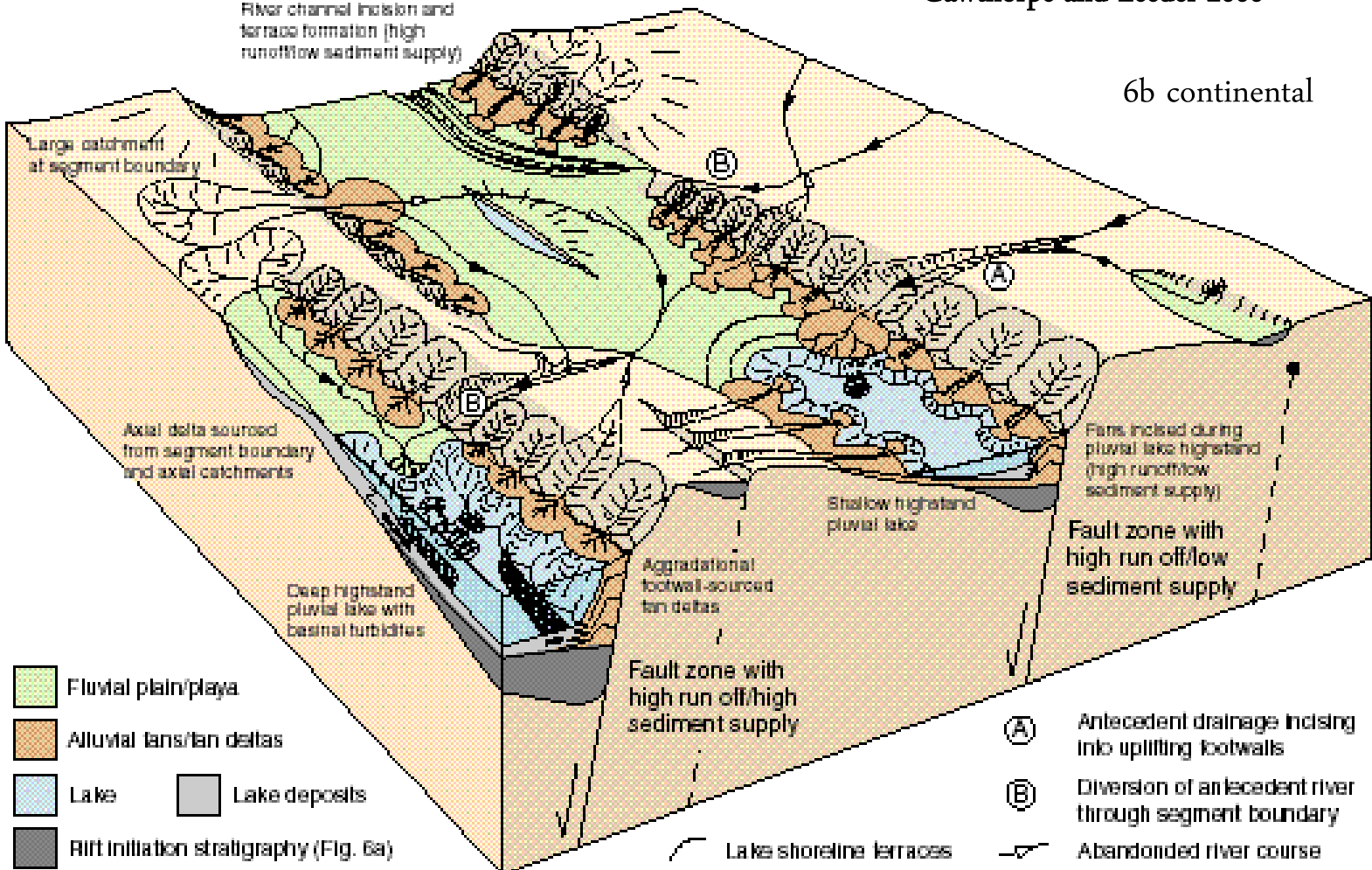
CARBONATE COASTAL/SHELF HALF-GRABEN MODEL

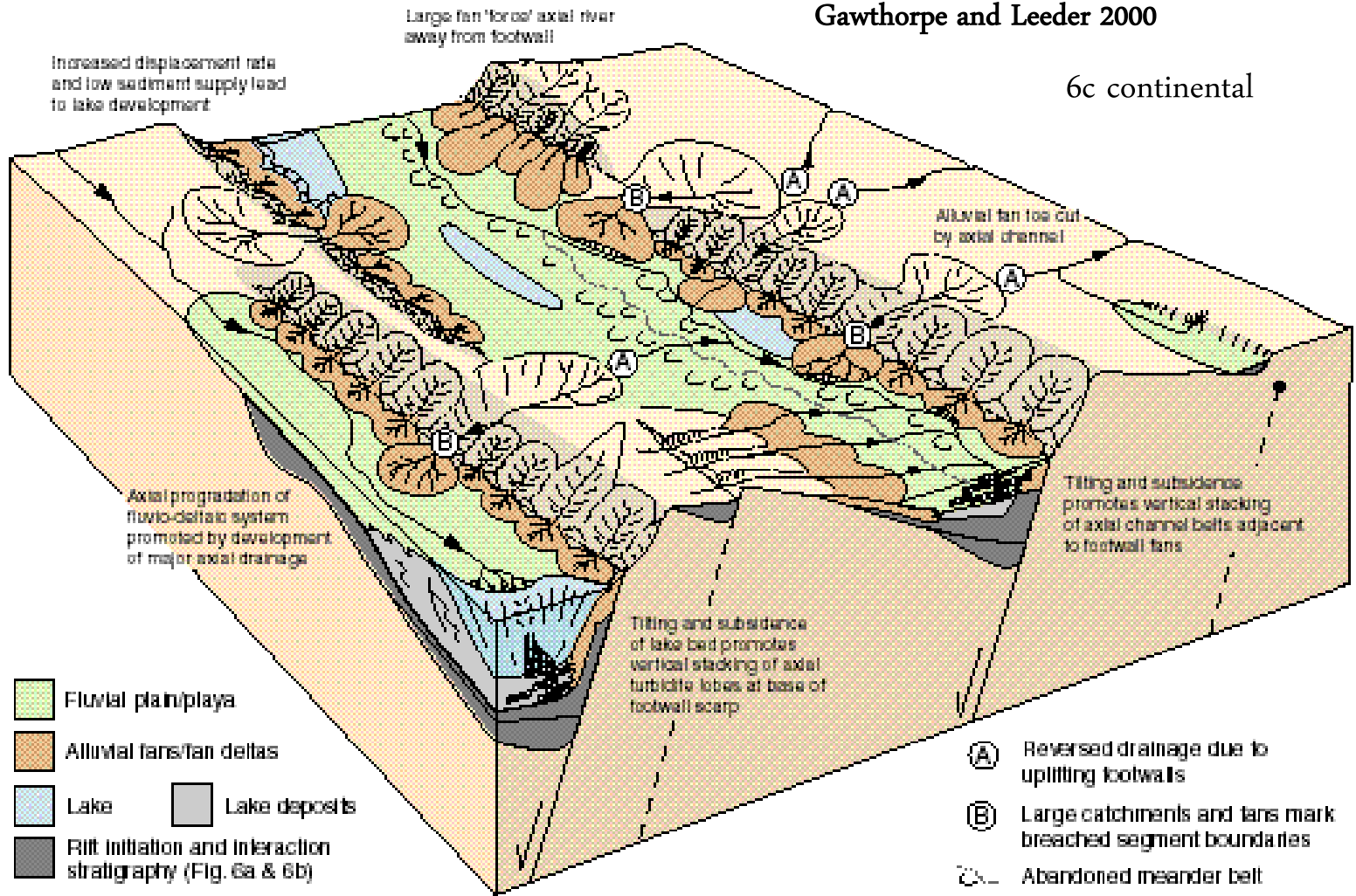


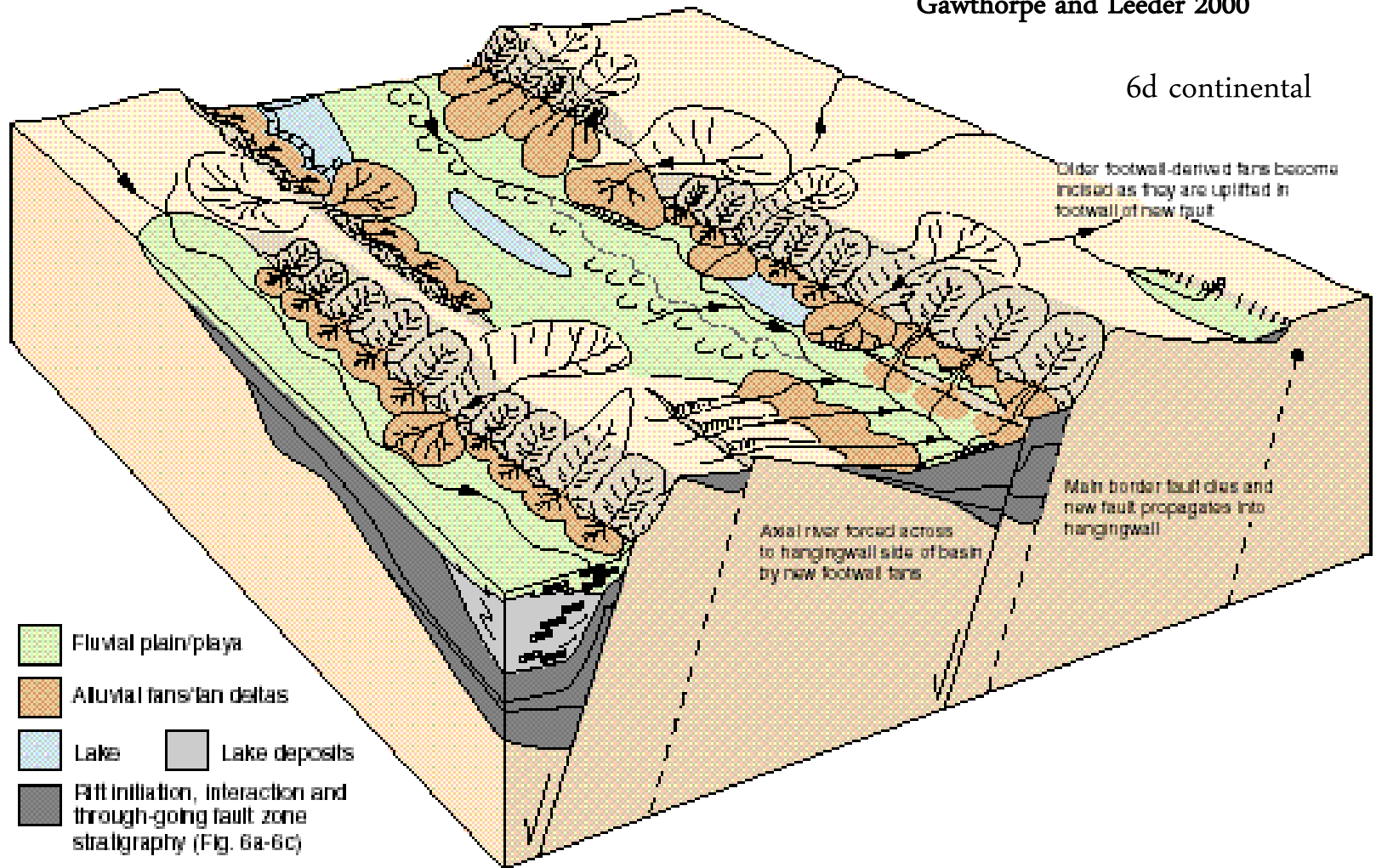
(modified from Leeder and Gawthorpe, 1987)



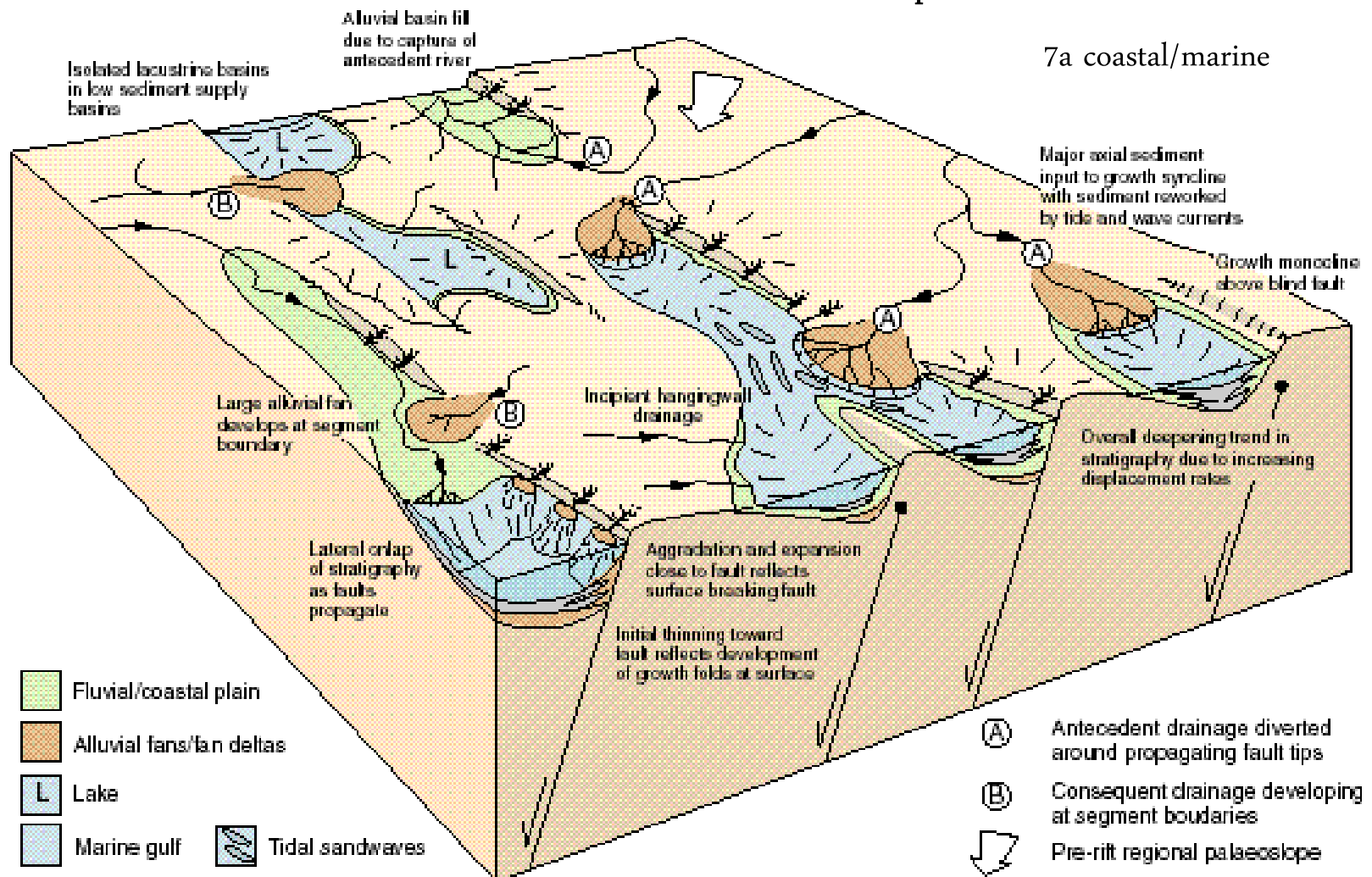




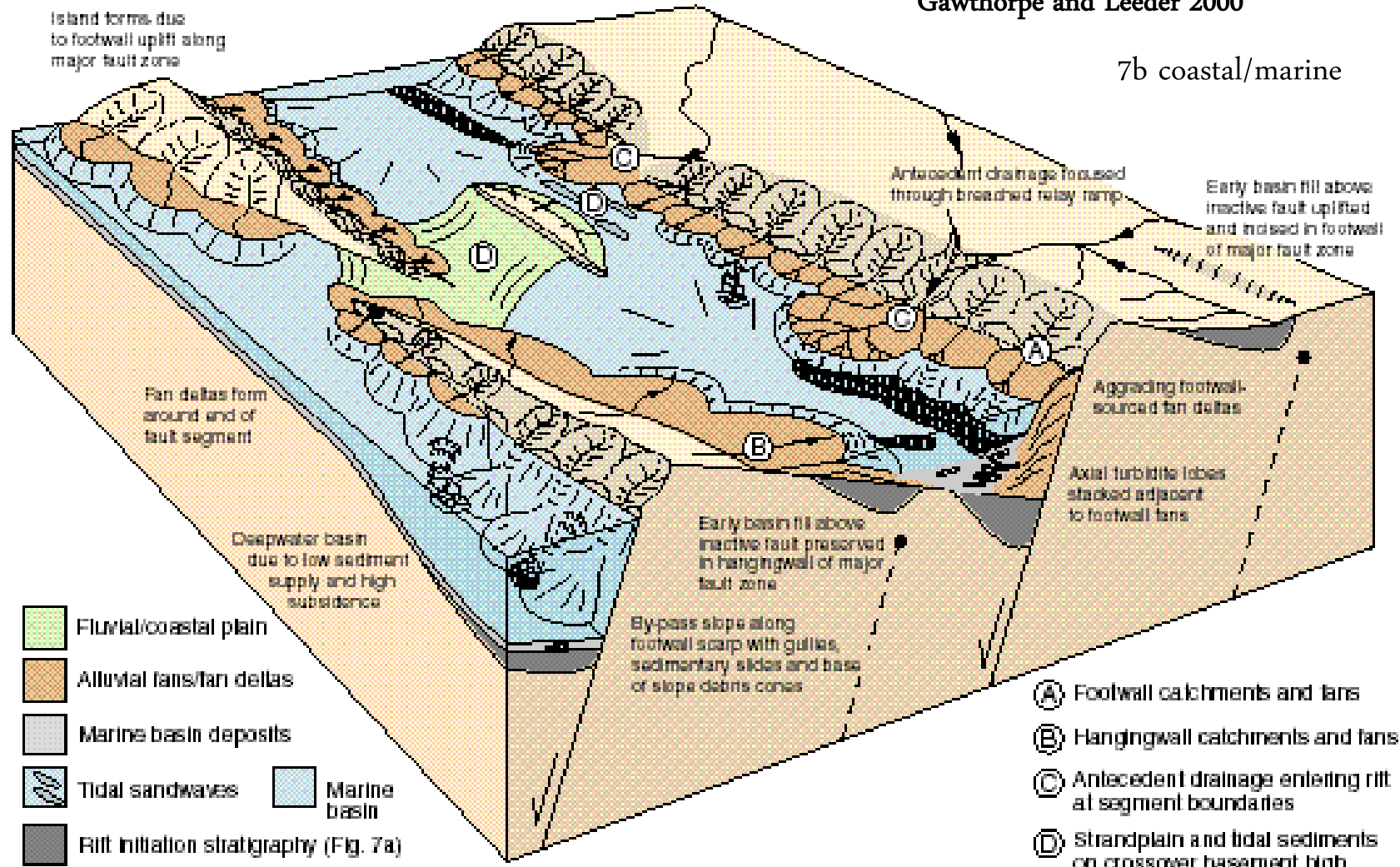




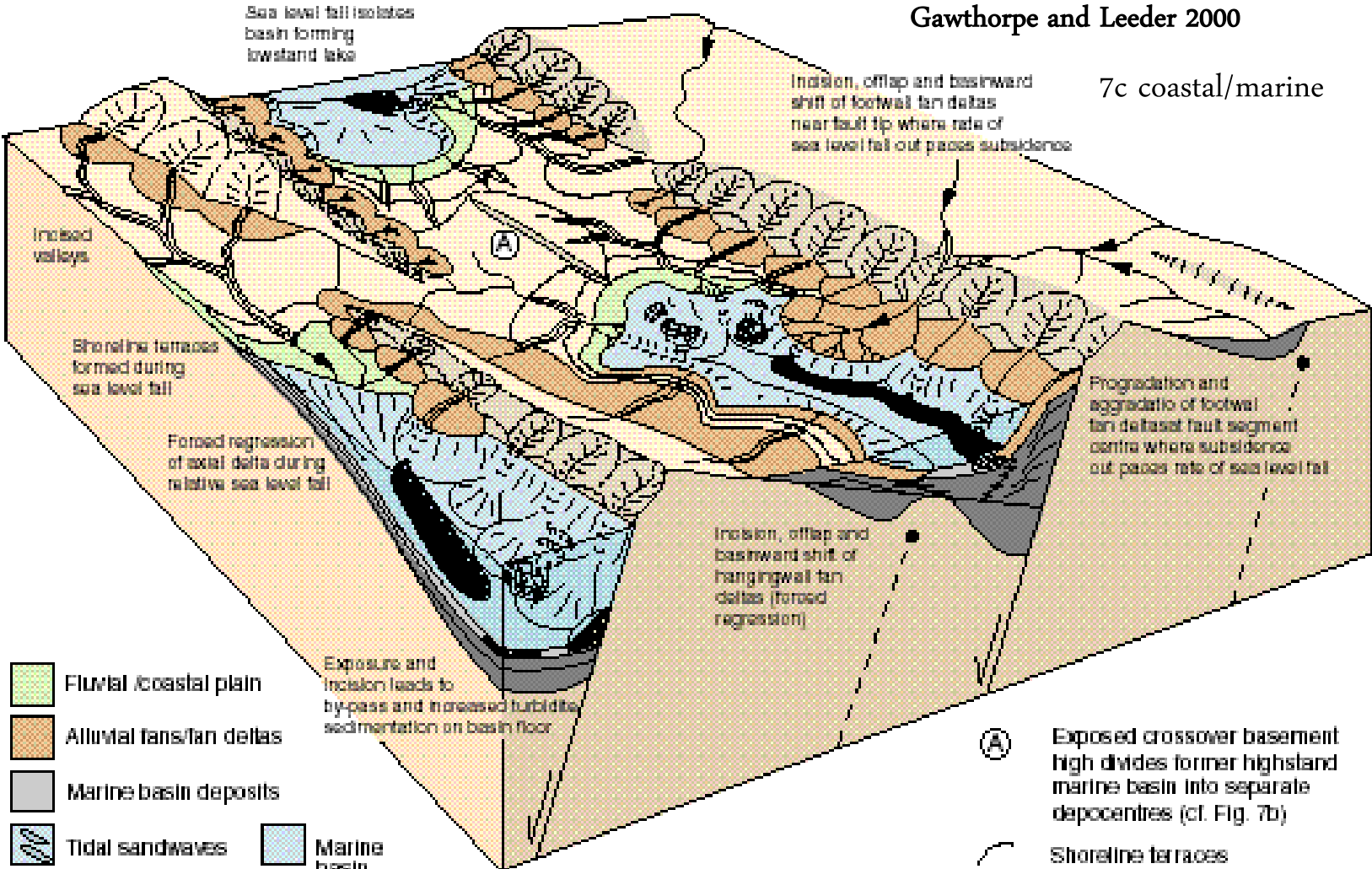
7a coastal/marine



7b coastal/marine



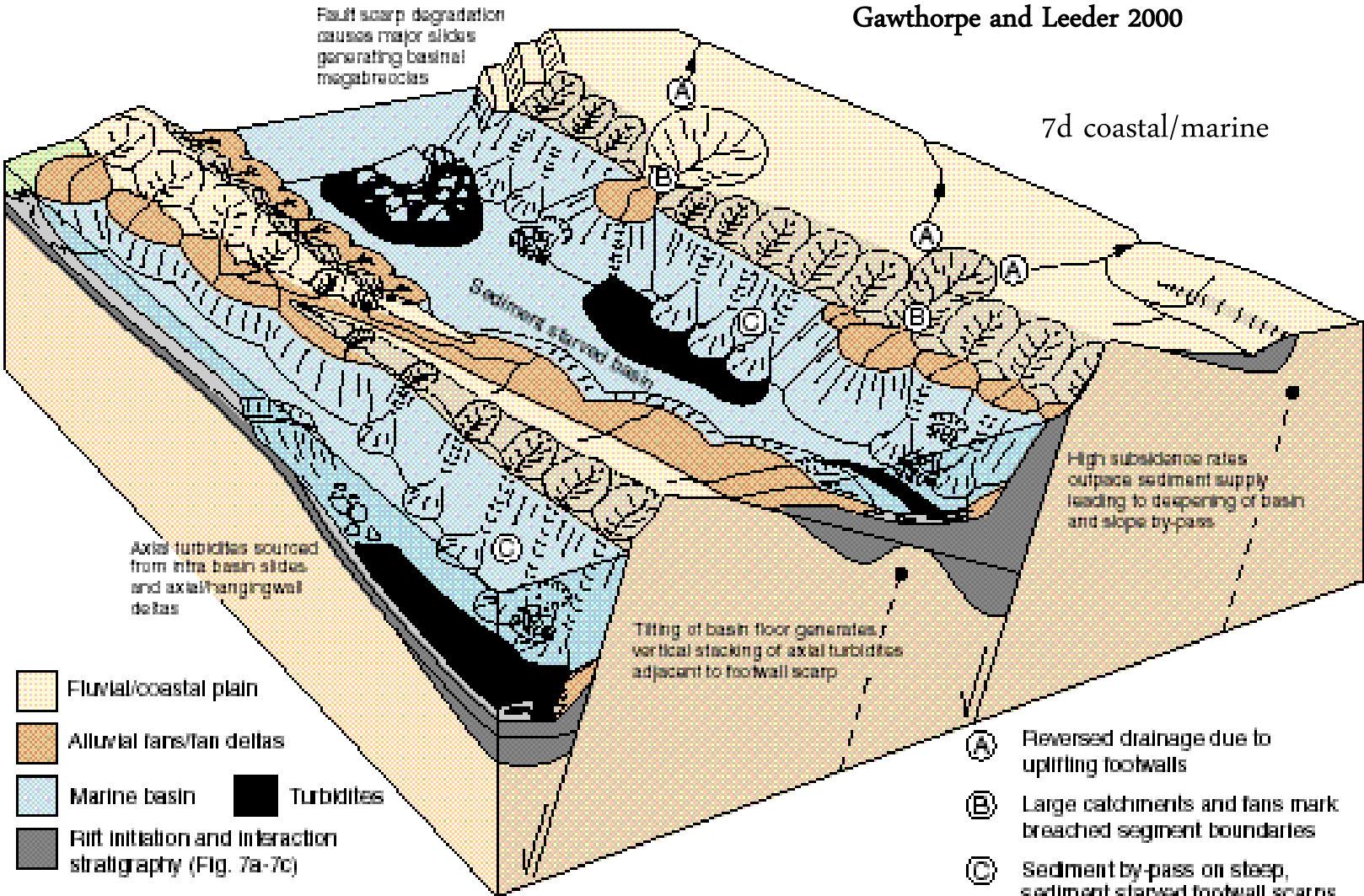
7c coastal/marine



- Fluvial /coastal plain
- Alluvial fans/fan deltas
- Marine basin deposits
- Tidal sandwaves
- Marine basin
- Rift initiation and interaction (highstand) stratigraphy (Fig. 7a, b)

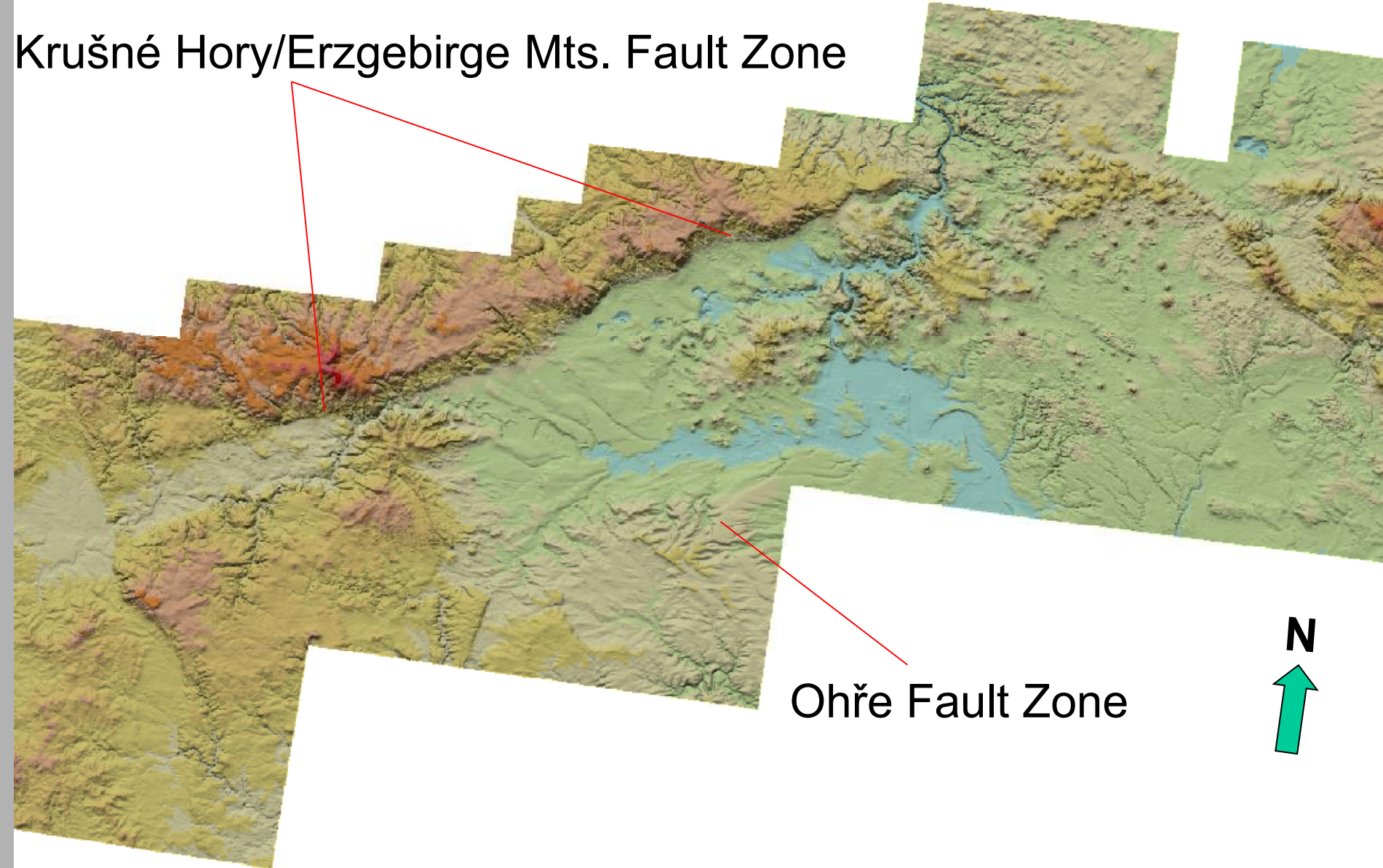
- A Exposed crossover basement high divides former highstand marine basin into separate depocentres (cf. Fig. 7b)
- Shoreline terraces

7d coastal/marine



SW-NE – trending structures prominent in Recent topography

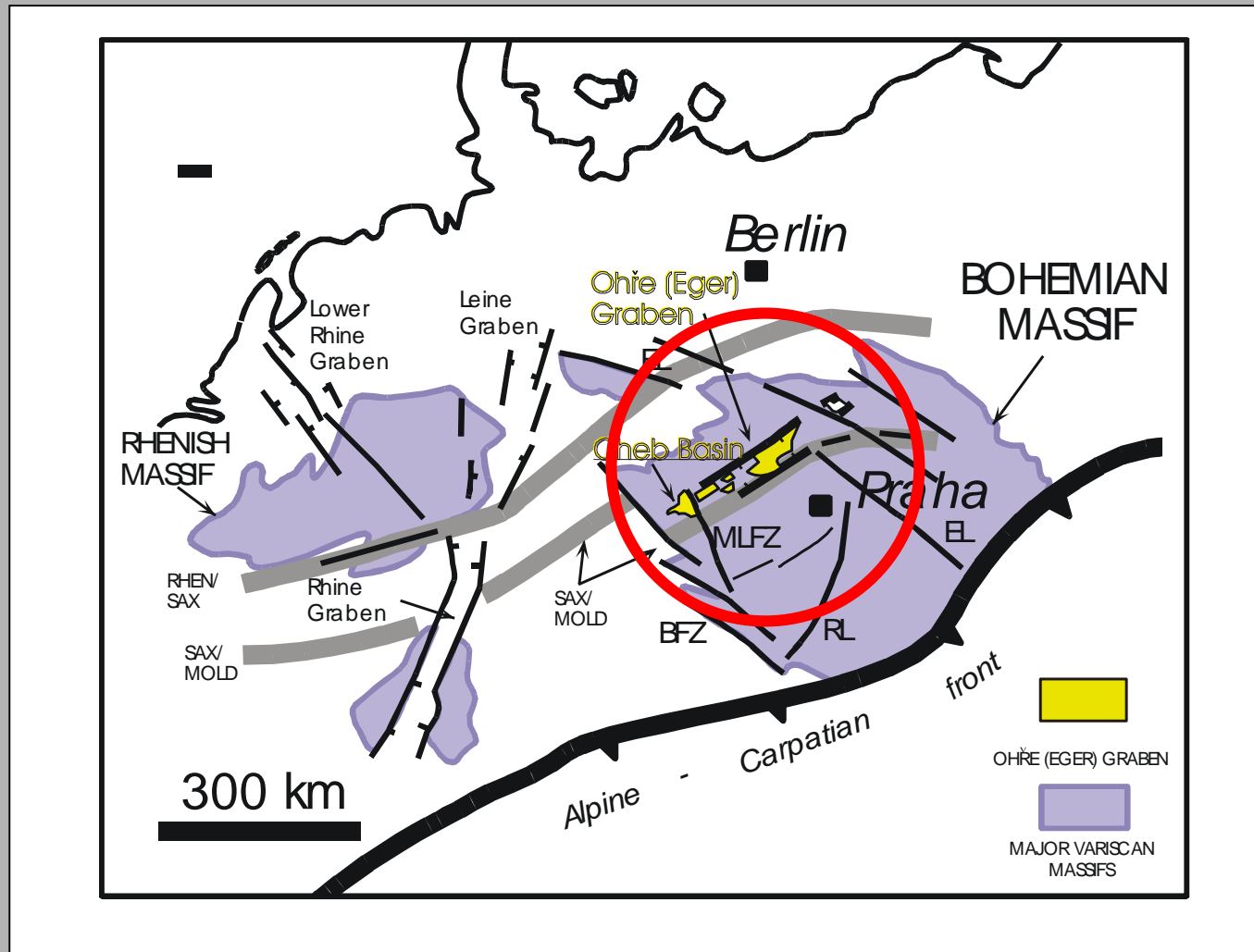
Krušné Hory/Erzgebirge Mts. Fault Zone



Ohře Fault Zone

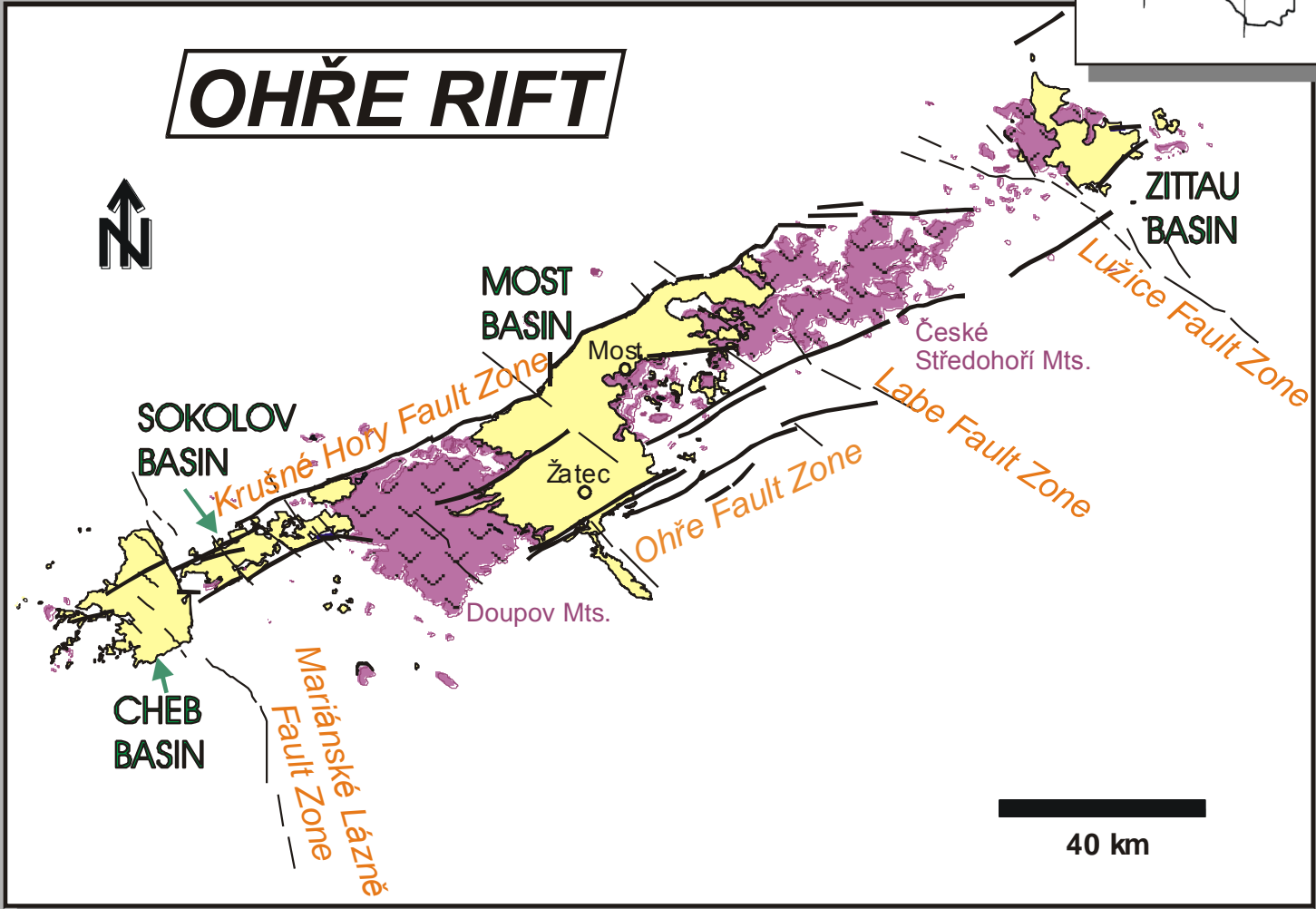
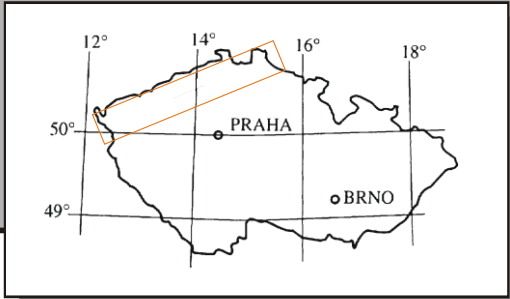


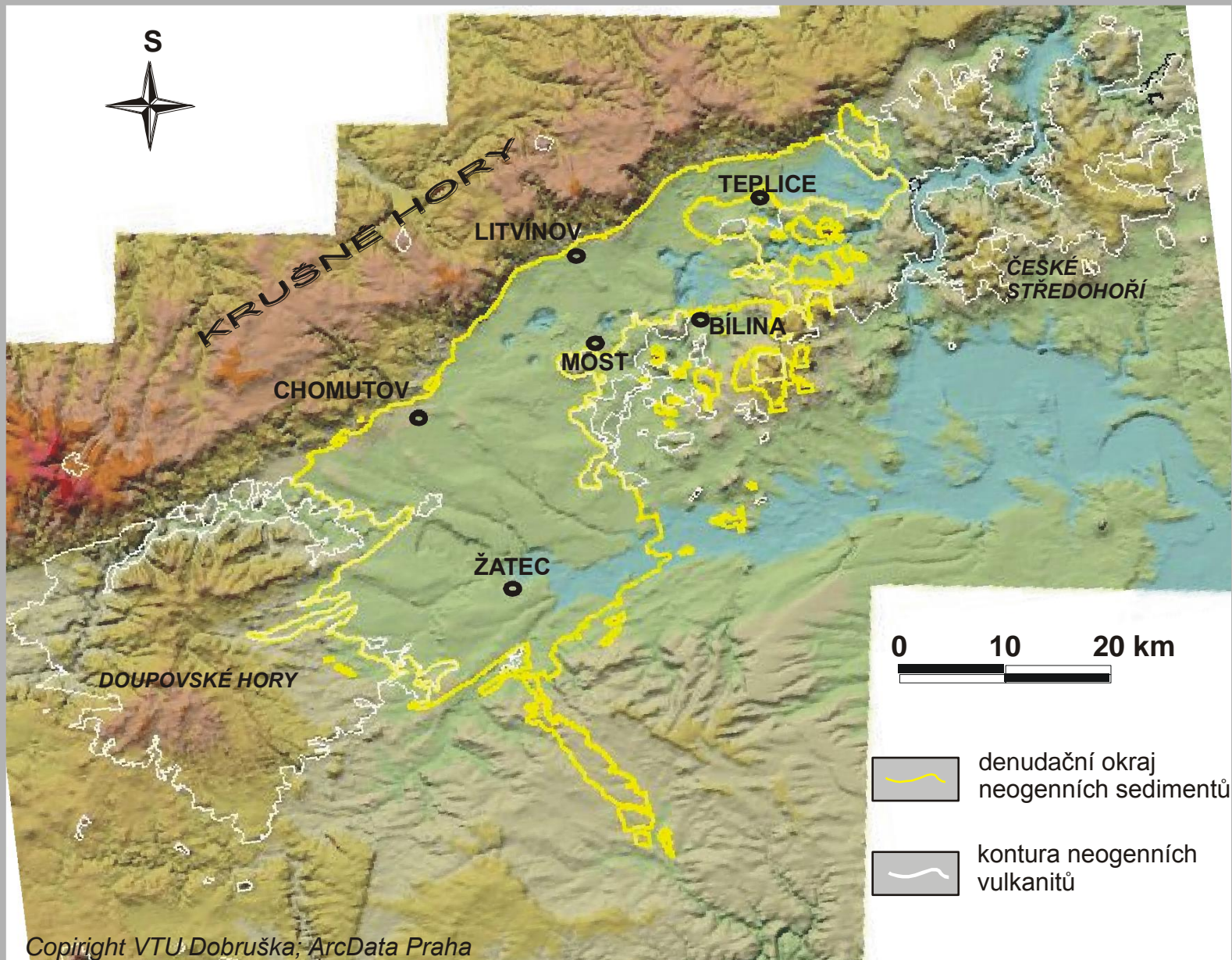
Eger Rift - part of the European Cenozoic rift system:



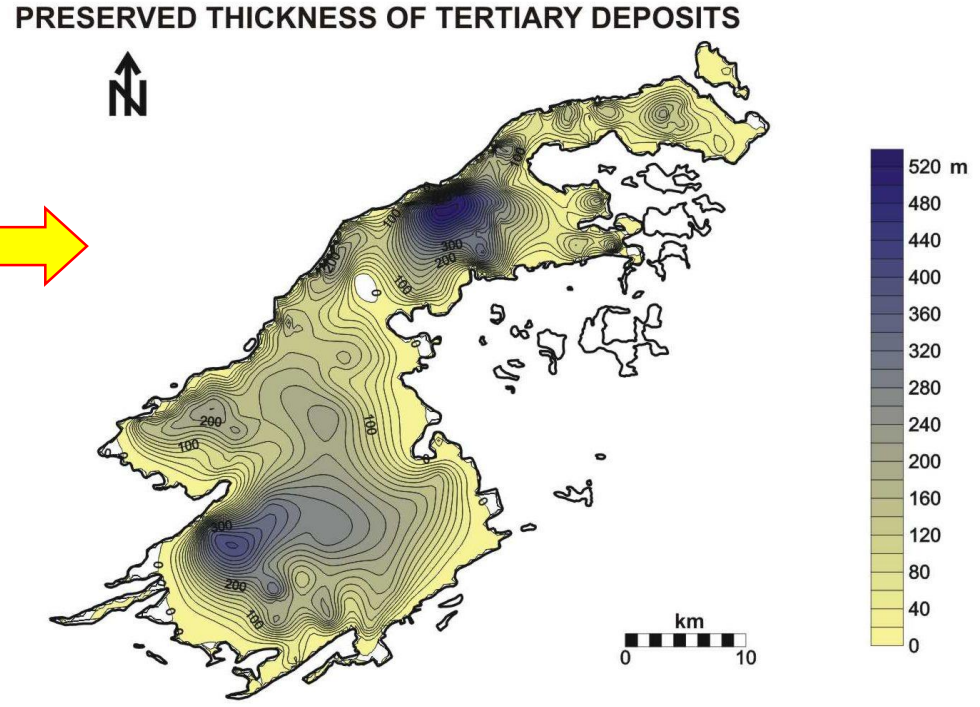
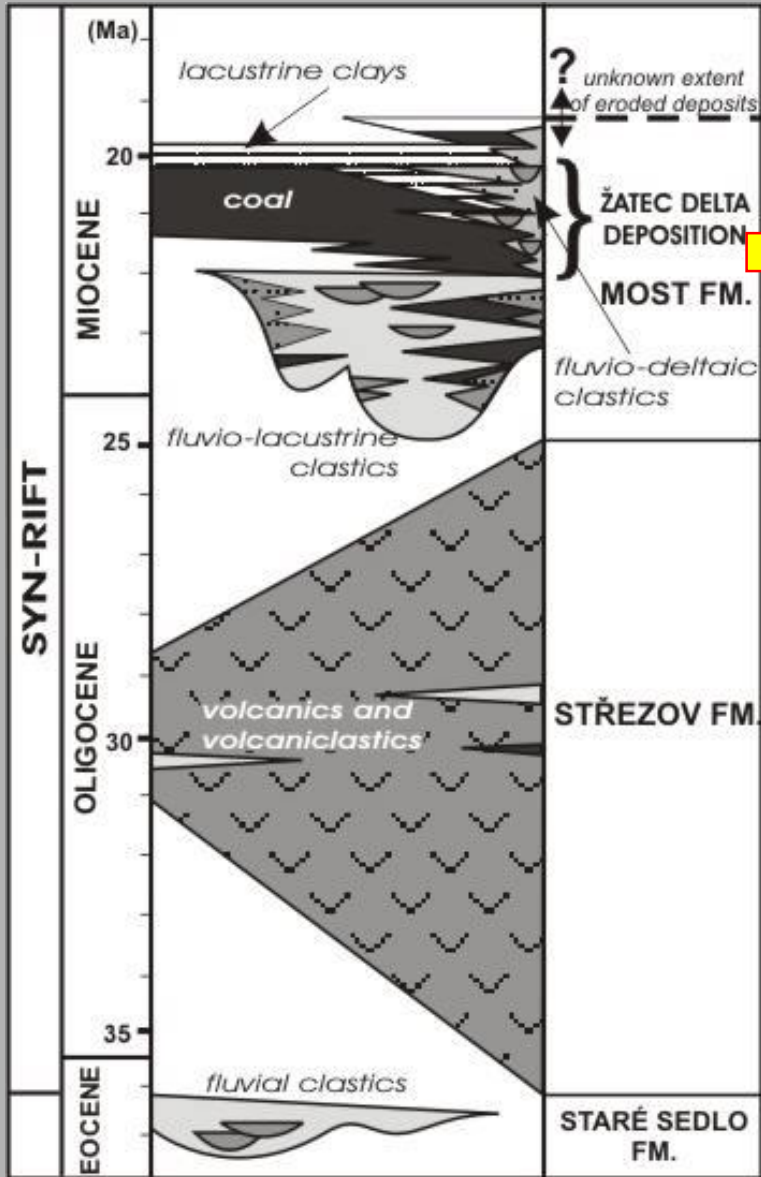
Eger Rift axis parallel to the Saxothuringian/Teplá-Barrandian suture

OHŘE RIFT



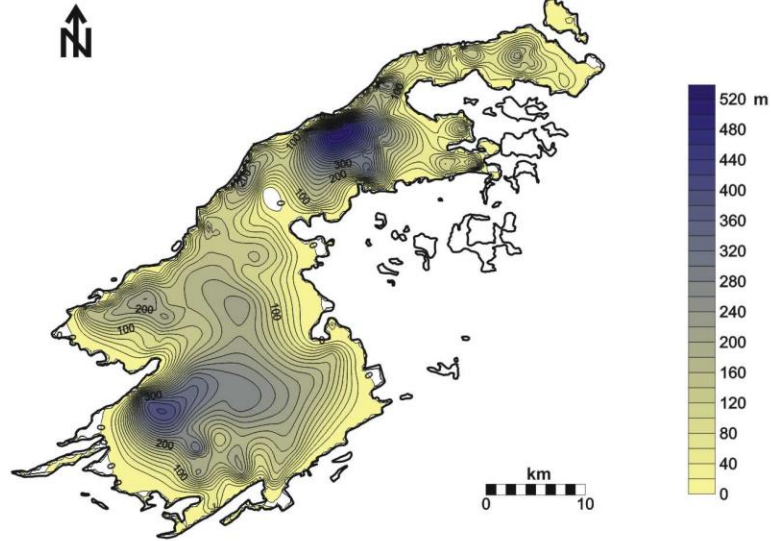


THE MOST BASIN

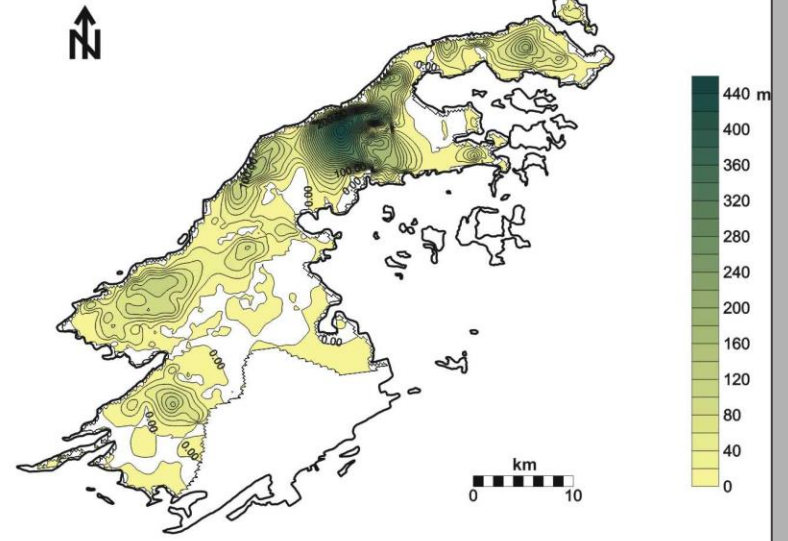


ISOPACH MAPS OF BASIN FILL - MOST BASIN

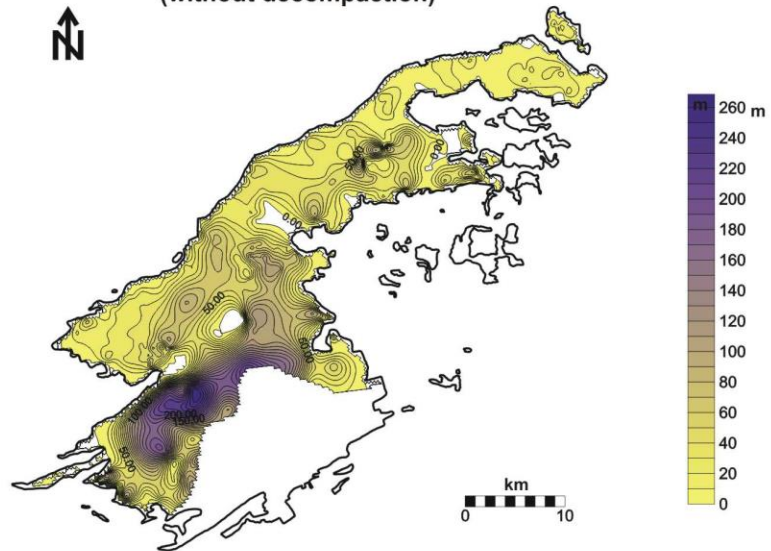
PRESERVED THICKNESS OF TERTIARY DEPOSITS



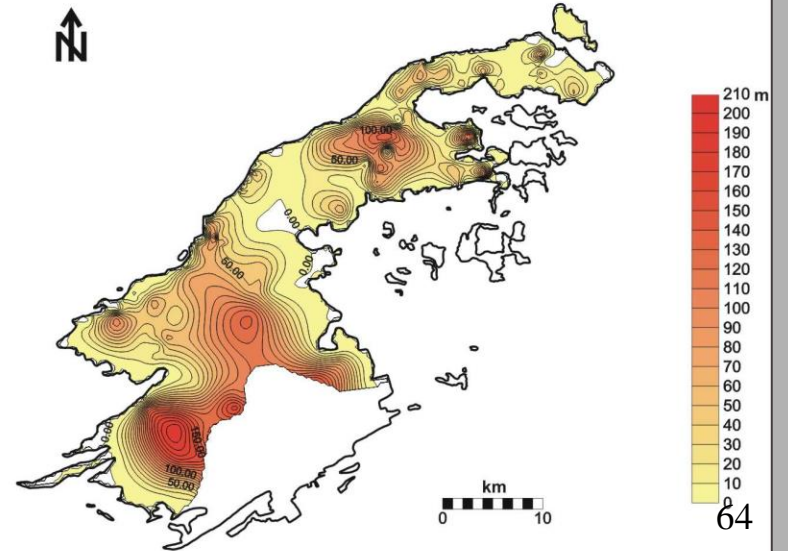
CLASTICS OVERLYING THE MAIN SEAM



THE MAIN SEAM AND COEVAL CLASTICS (without decompaction)

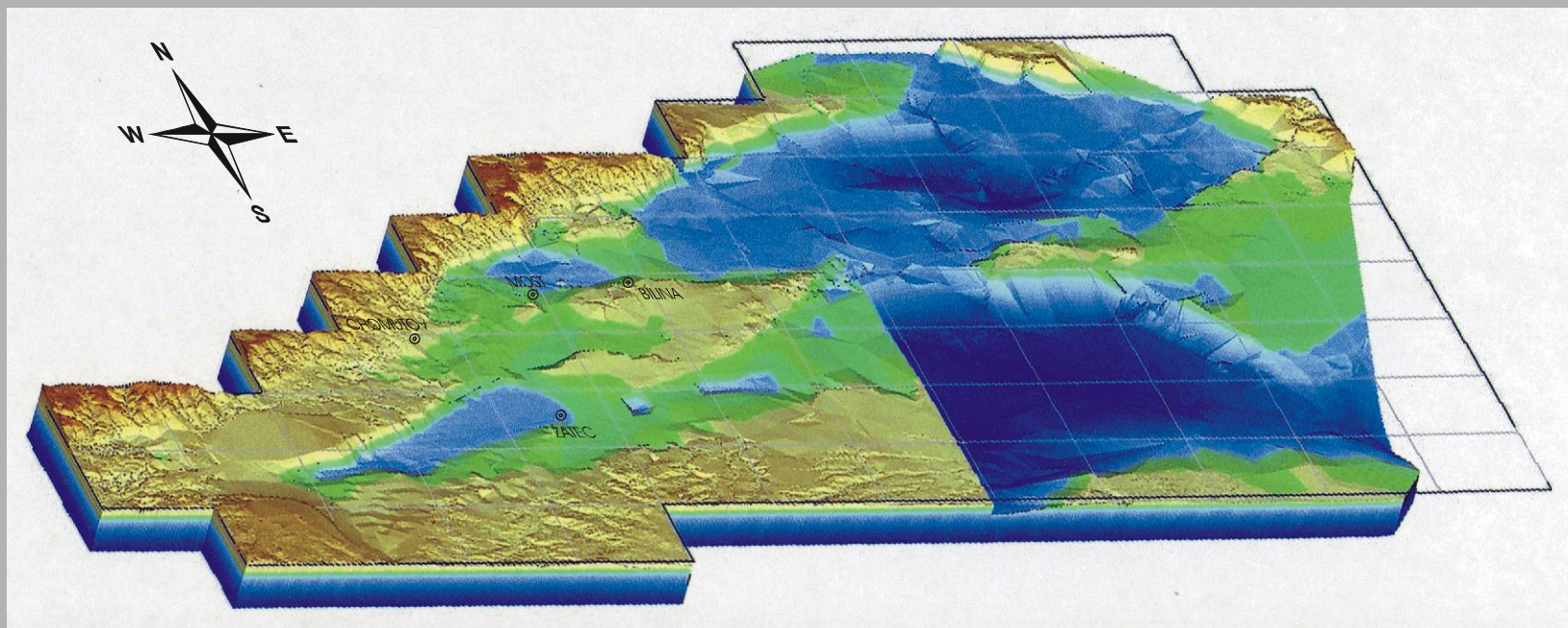


CLASTICS /VOLCANICS UNDERLYING THE MAIN SEAM

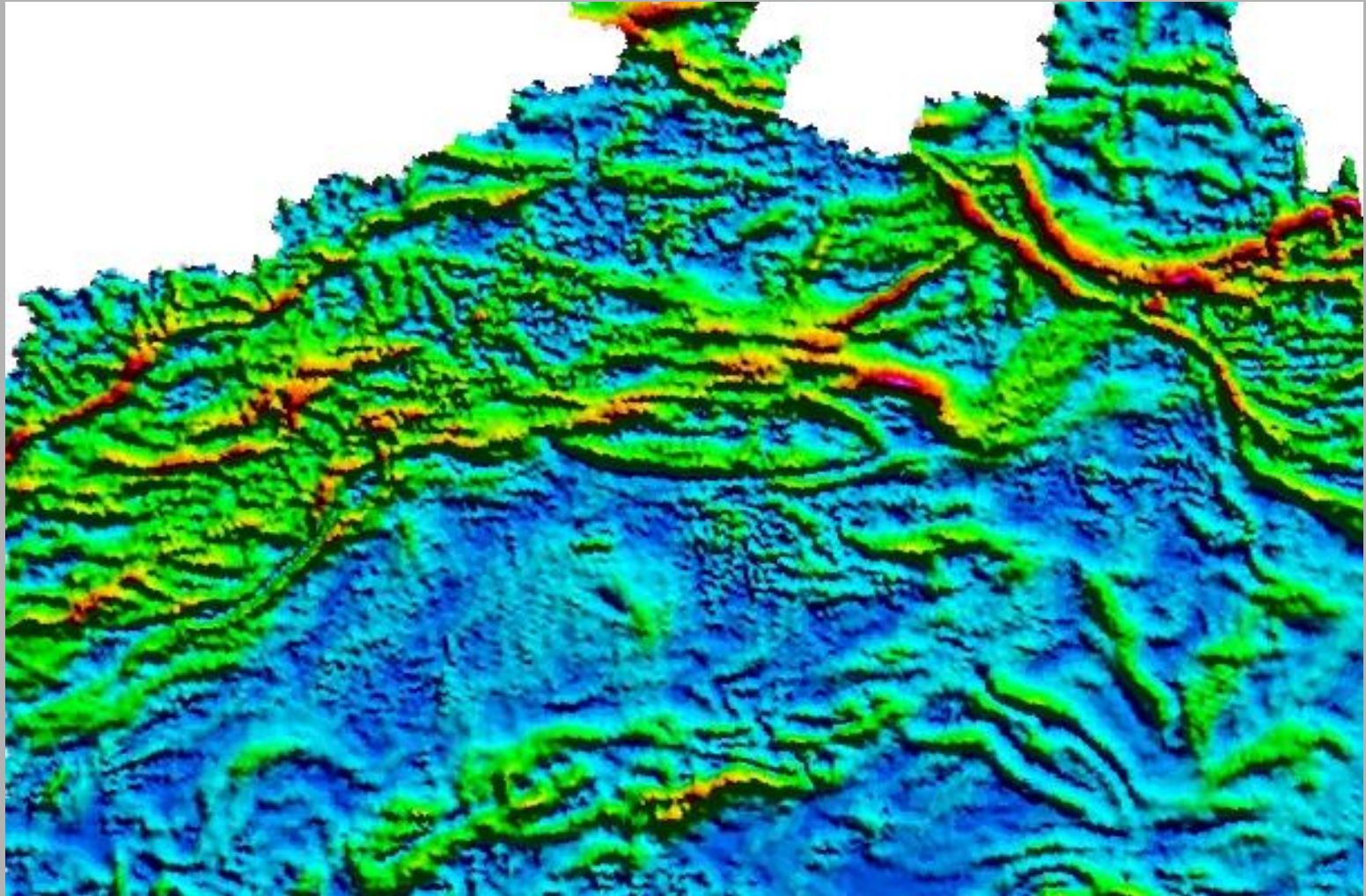


DMR KRYSTALINIKA V PODLOŽÍ MOSTECKÉ PÁNVE

(podle Mlčocha et al. 2002)

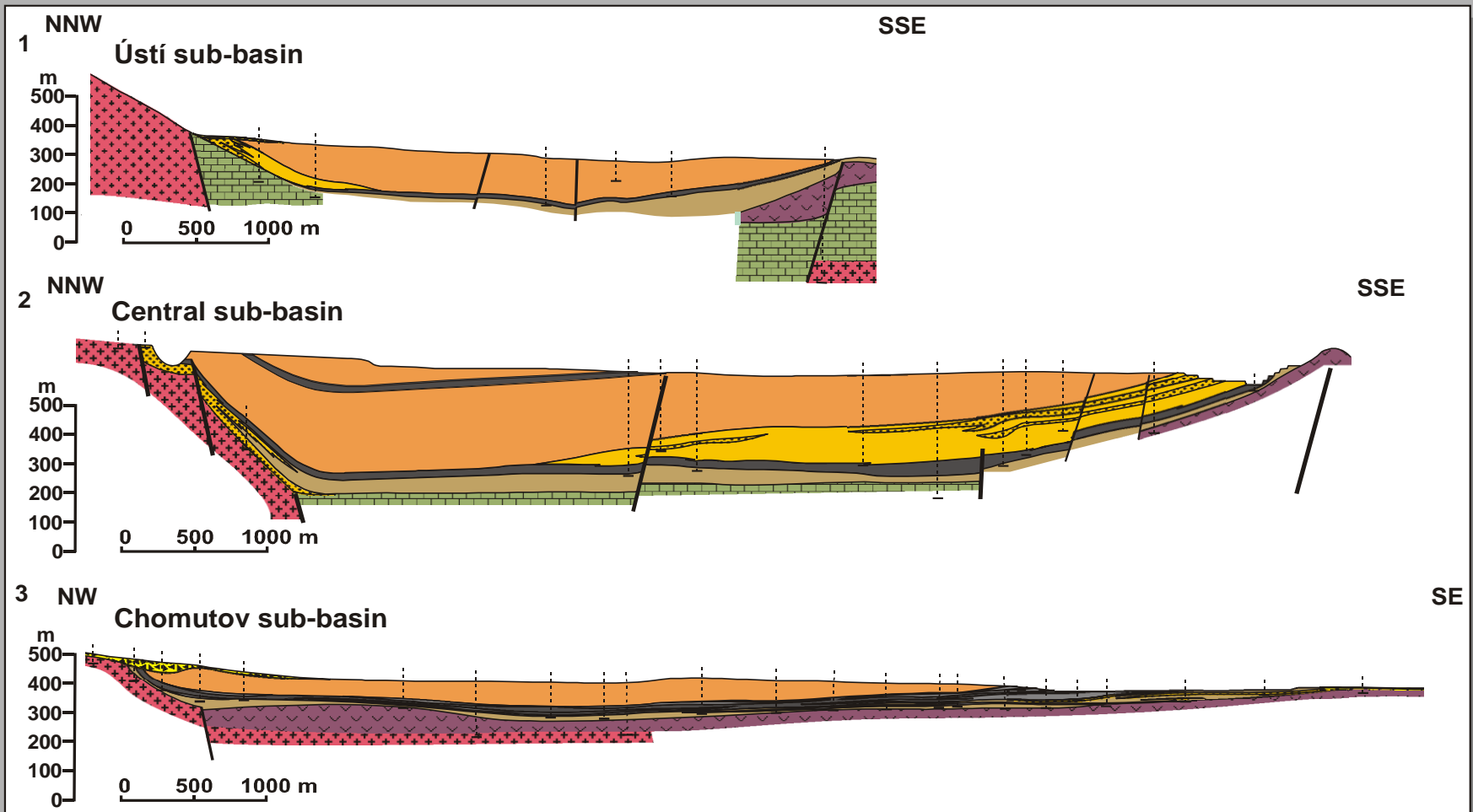


northern part of the Eger Rift

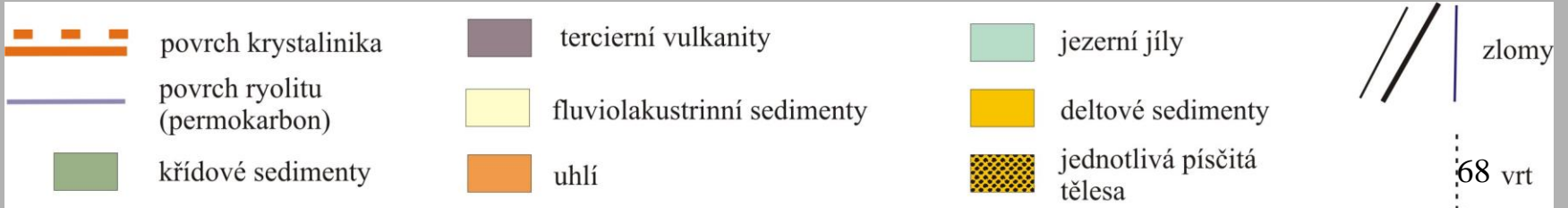
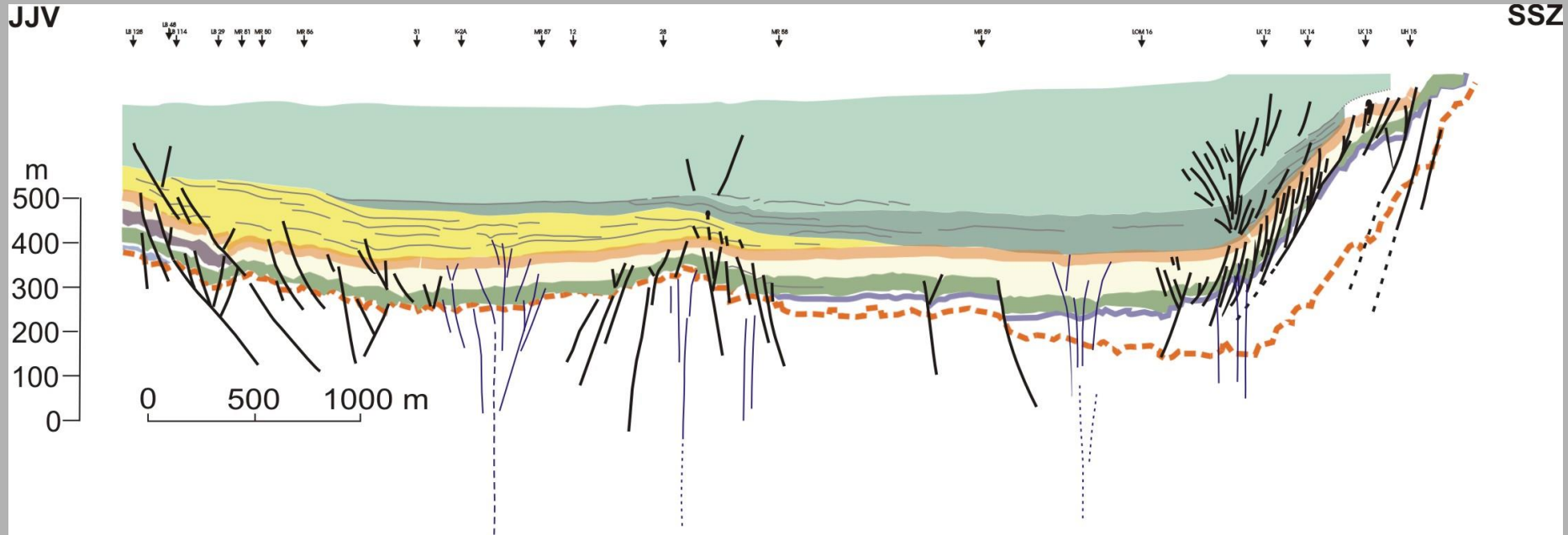
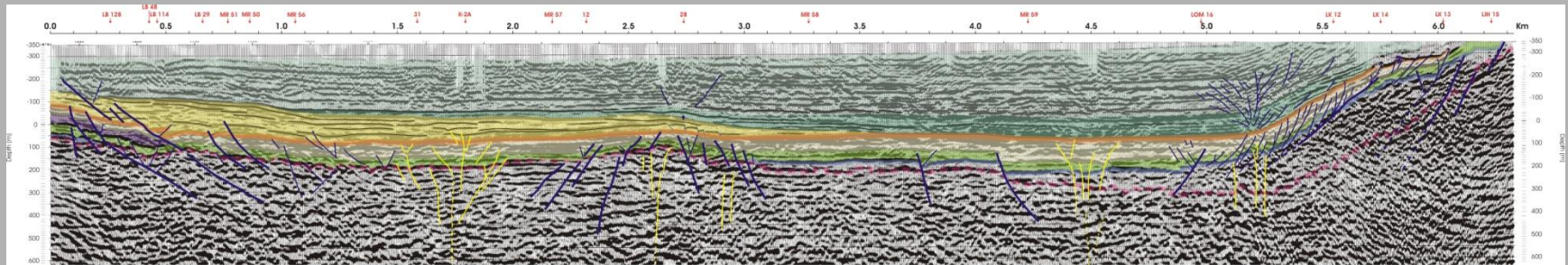


E-W structures in horizontal gravity gradients, illumination from the NNE

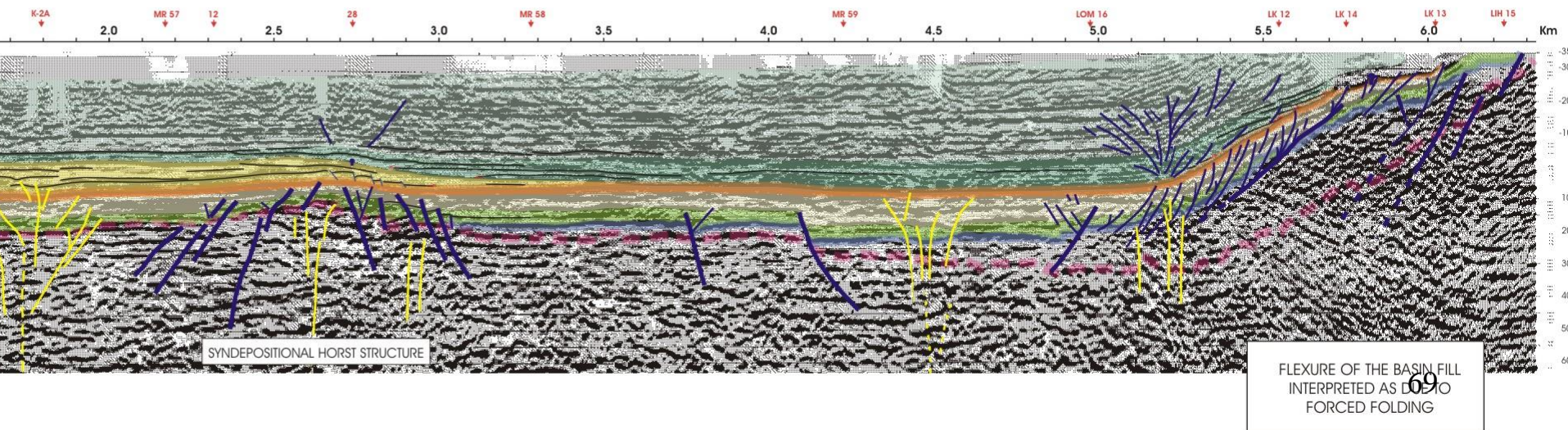
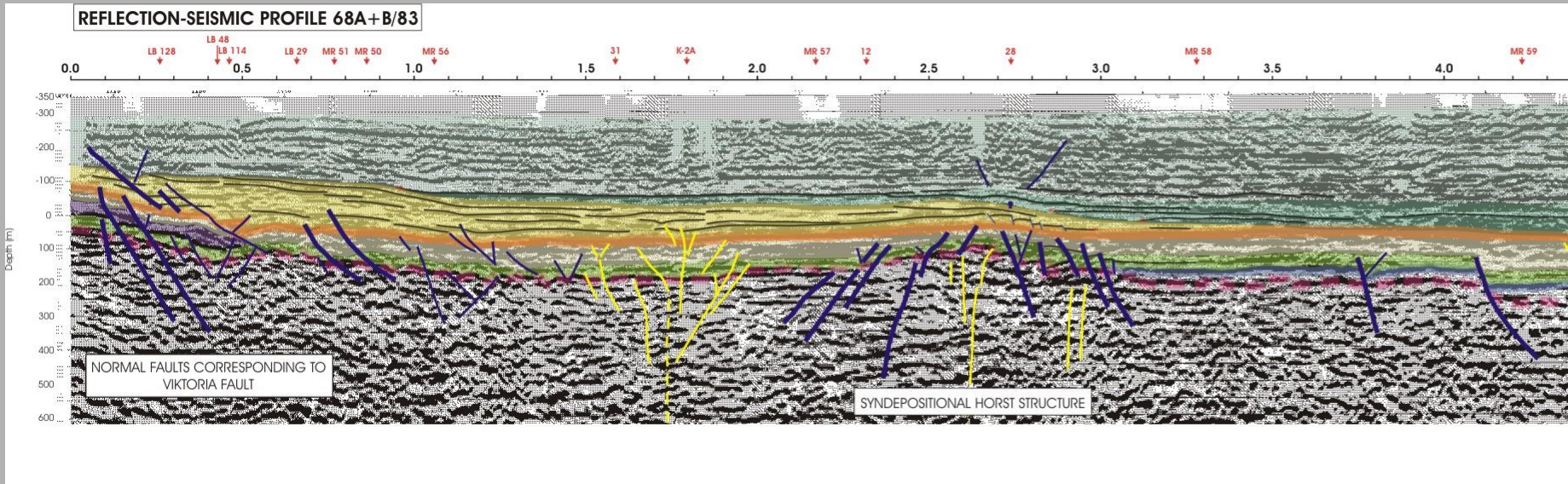
1. GEOMETRIE DEPOCENTER – GEOMETRIE VÝPLNĚ



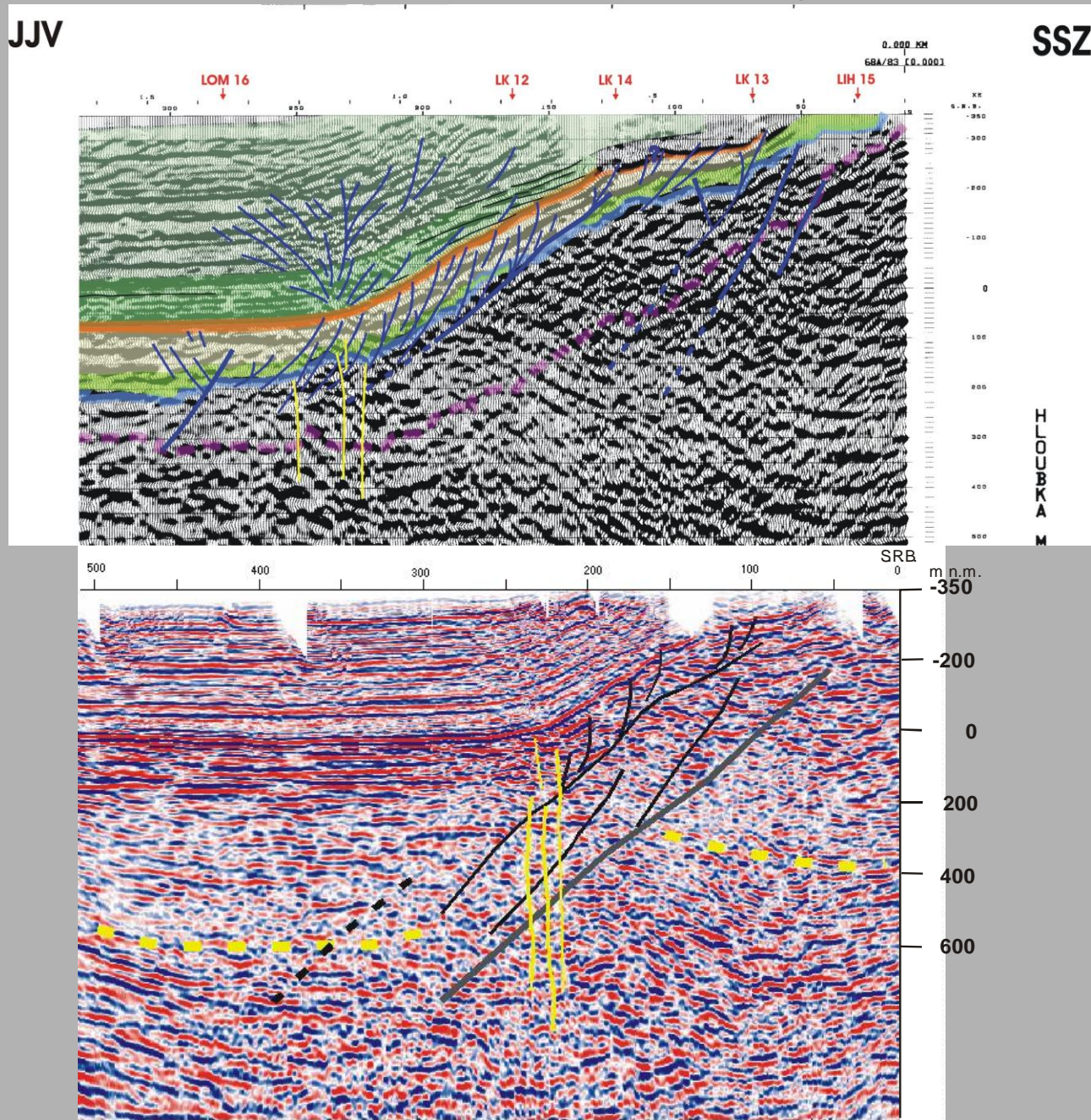
68A+B/83

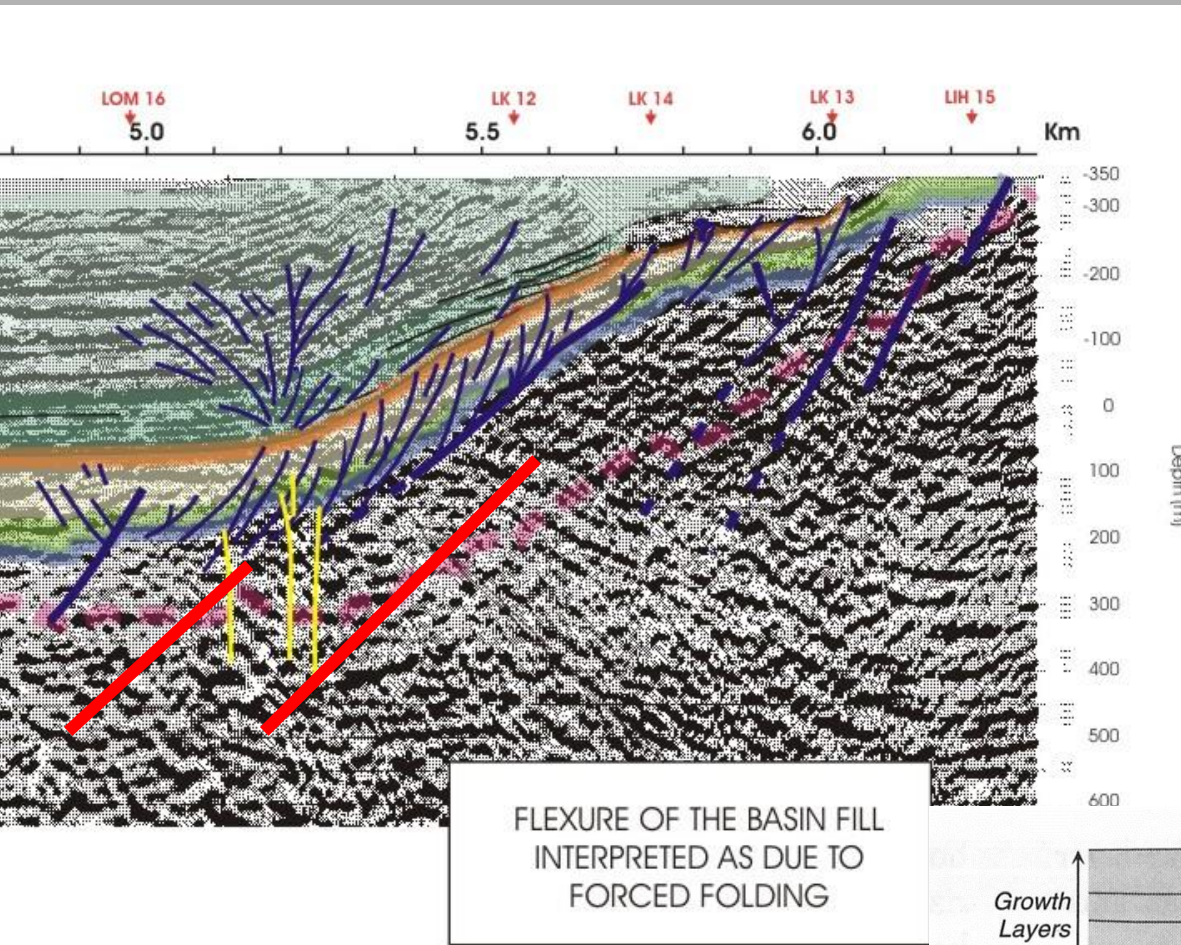


History of the Most Basin summarized in the seismic reflection line 68/83



Tektonická interpretace krošňohorského okraje mostecké pánve

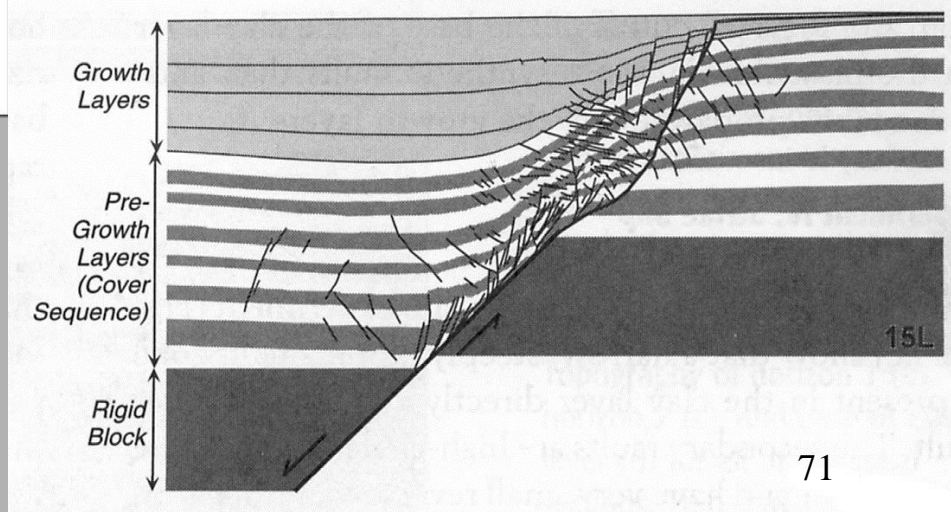




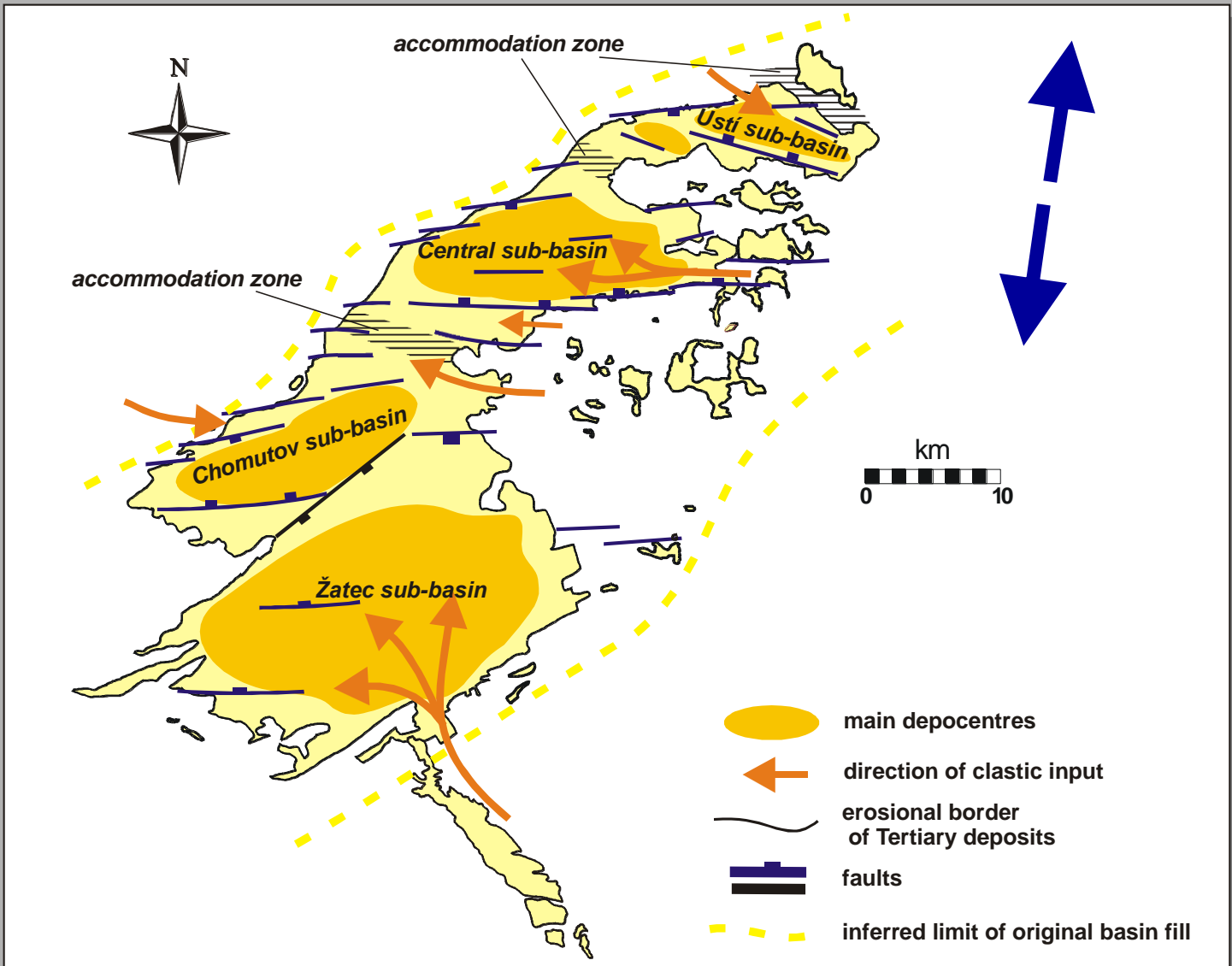
Krušné Hory/Erzgebirge
Mts. uplift:
Plio-Pleistocene,
deformation of part
of basin fills

vertical displacement
over 1000 m

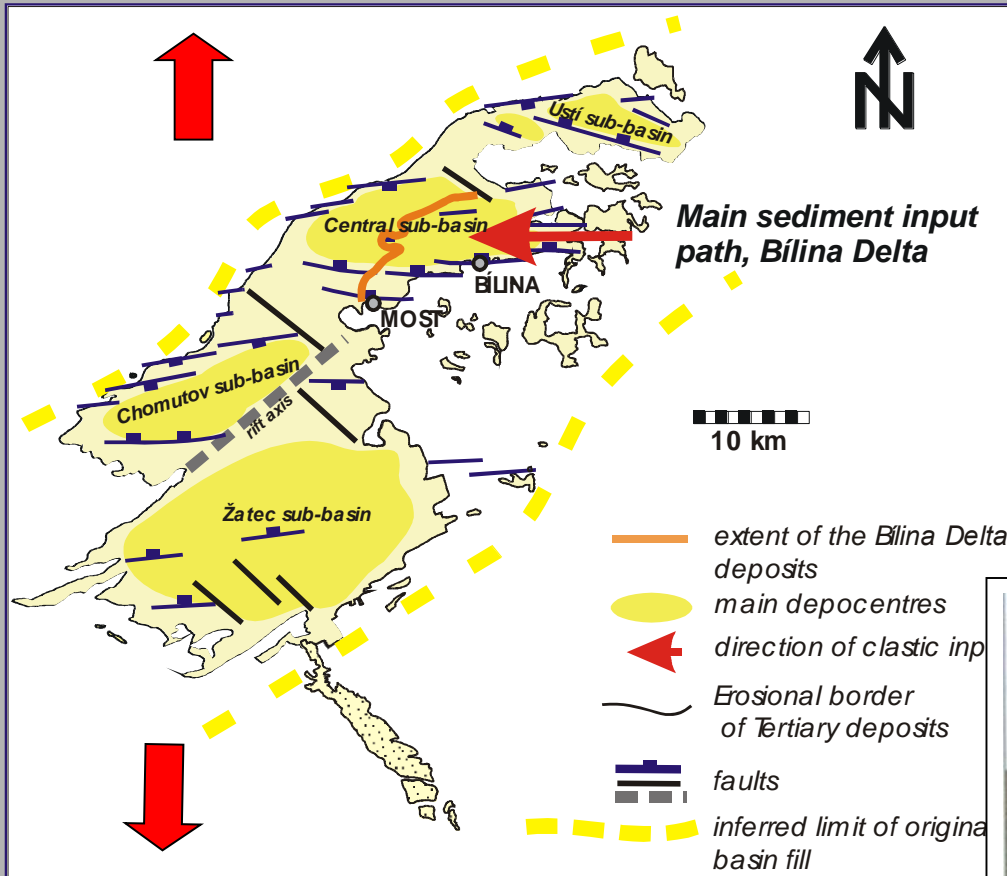
analogue model from
Schlische et al., 2002



TEKTONICKÝ MODEL MOSTECKÉ PÁNVE



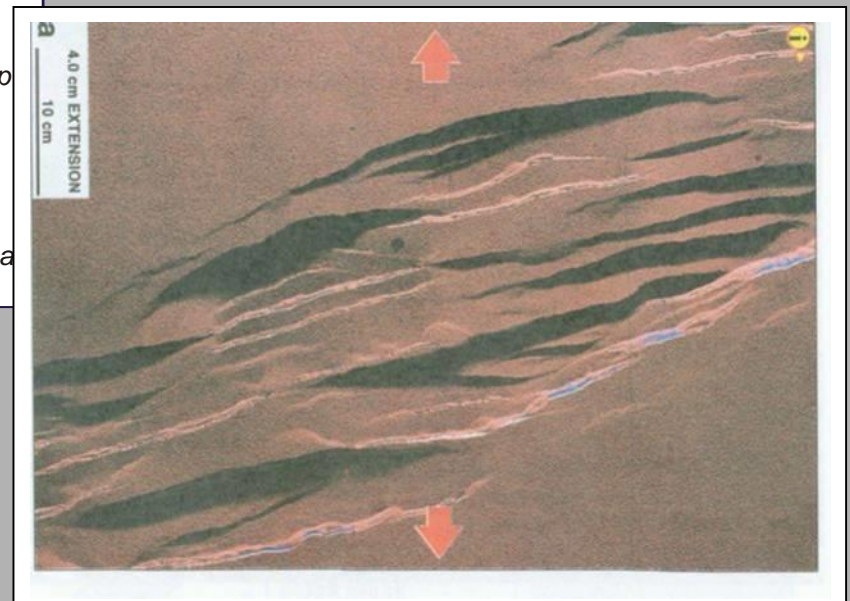
FORMATION OF MOST BASIN: OBLIQUE EXTENSION



short E-W normal fault segments
 en-échelon arrangement
 partly overlapping depocenters
 separated by accommodation
 /transfer zones

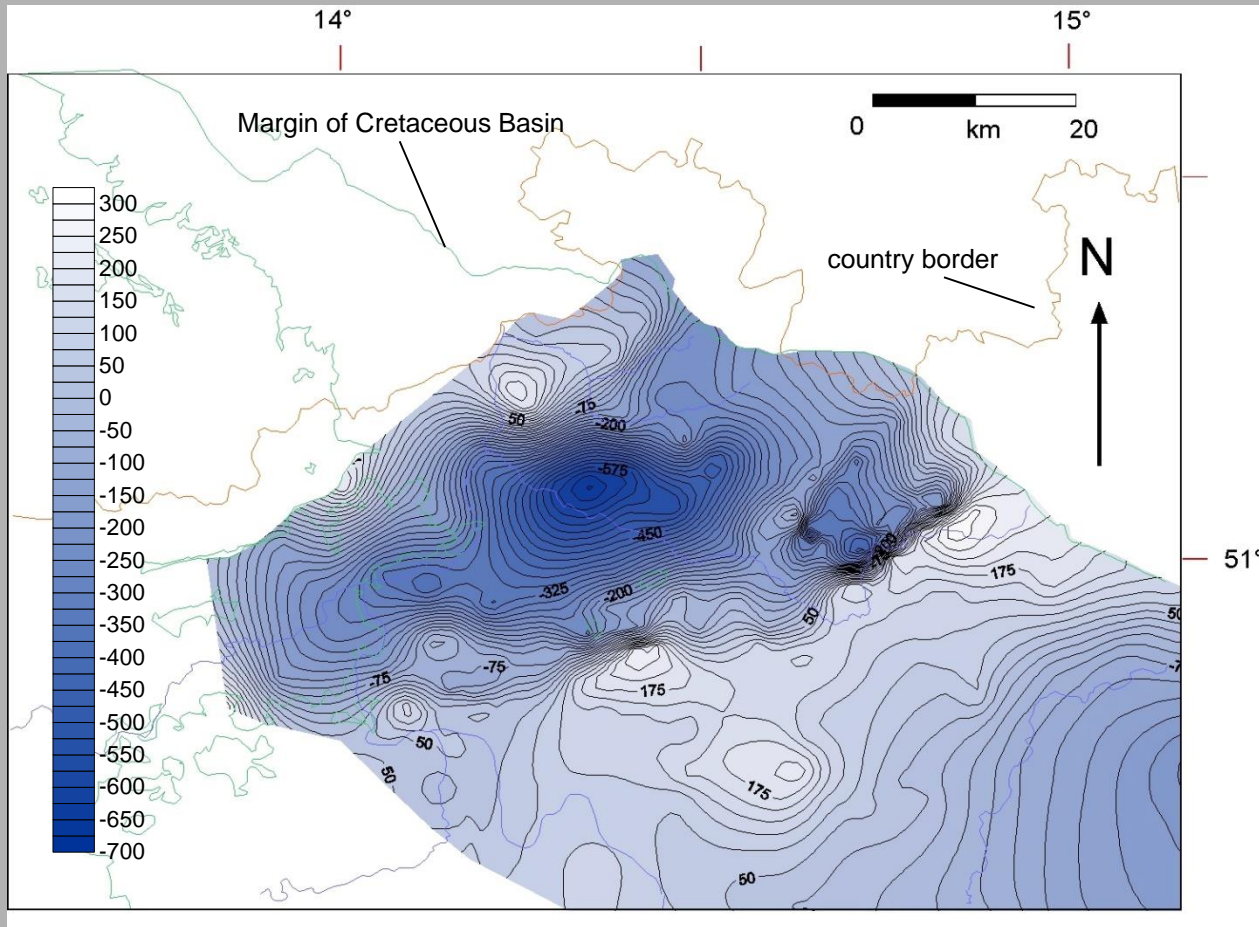
approx. N-S extension in
 Mid-Oligocene to early Miocene

(also: volcanic body geometries,
 Adamovič & Coubal, 1999)



comparison: analogue model from McClay, 2002

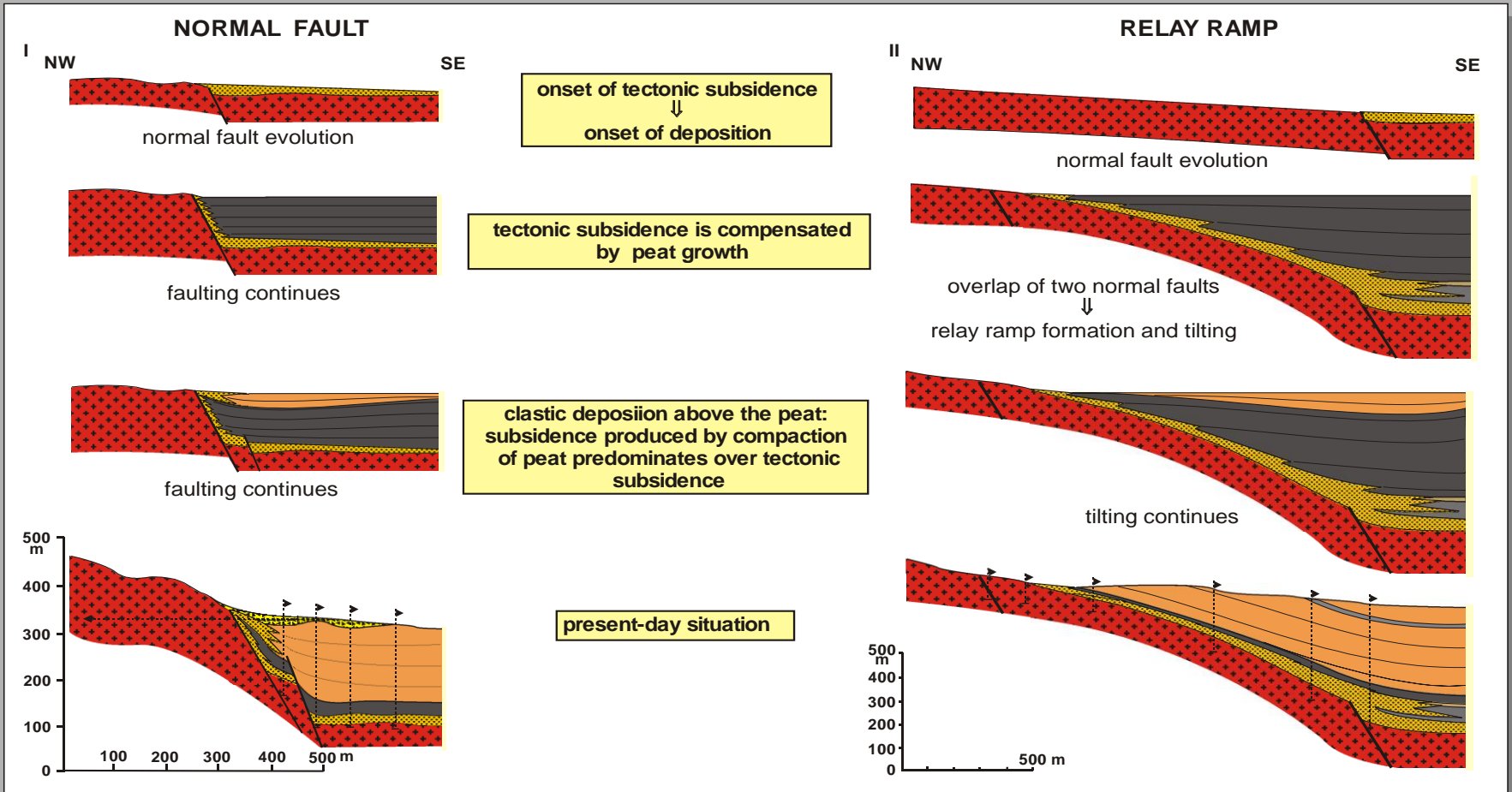
Northern Eger Rift – graben geometry



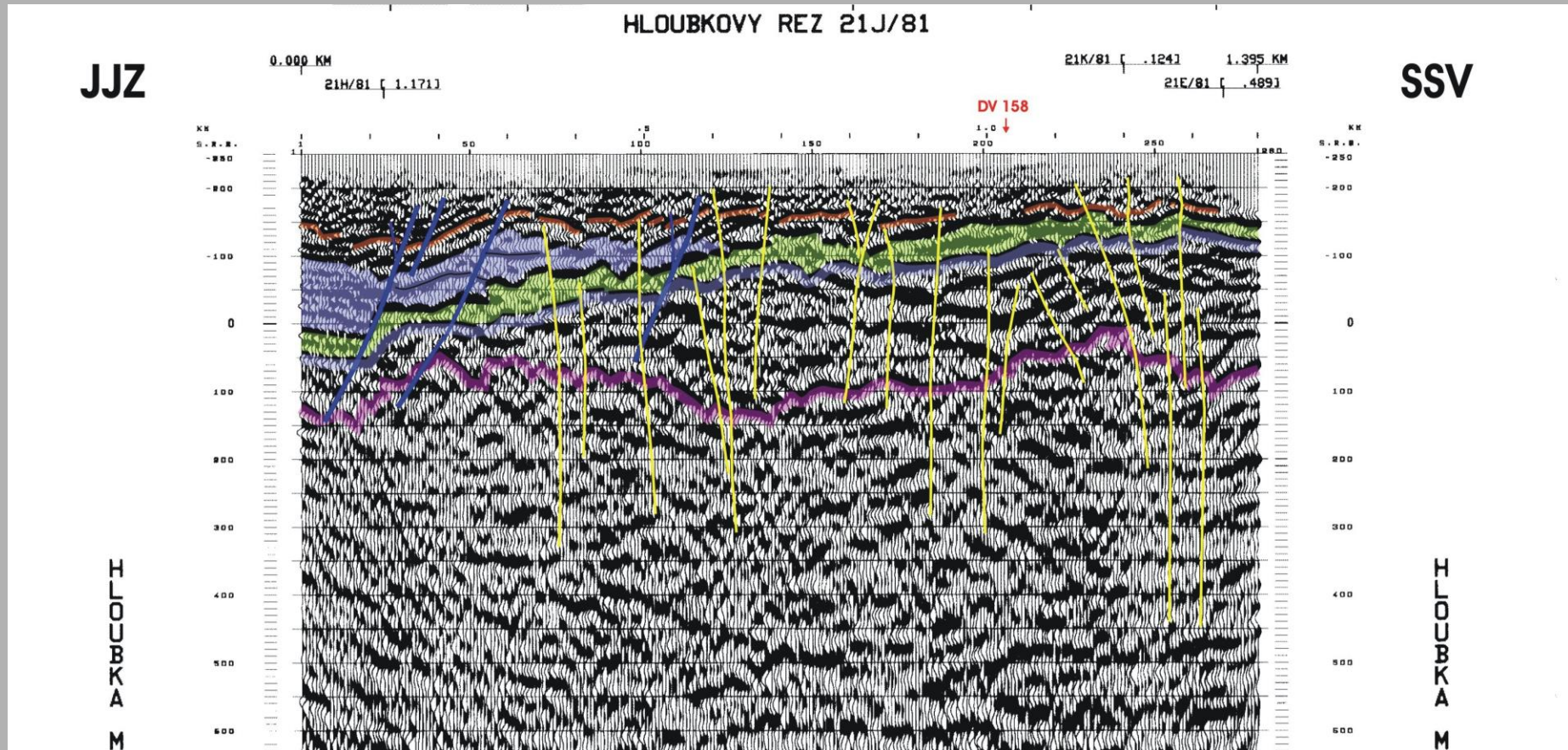
no preserved
Cenozoic basins,
but subsided blocks
of pre-rift strata
(Cretaceous)
mark E-W faults
and graben
geometries

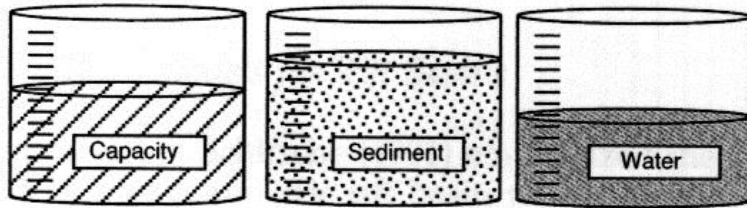
-655 to 295 m elevation, base of Turonian
deposits (mid/Cretaceous, quasi-horizontal
marker surface); based on c. 500 boreholes

1. VLIV KOMPAKCE RAŠELINY NA POVRCHOVÝ PROJEV ZLOMOVÉHO SYSTÉMU



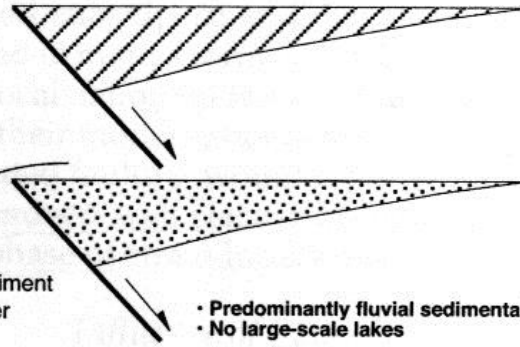
JZ okraj transferové zóny zastižené profilem 21/81



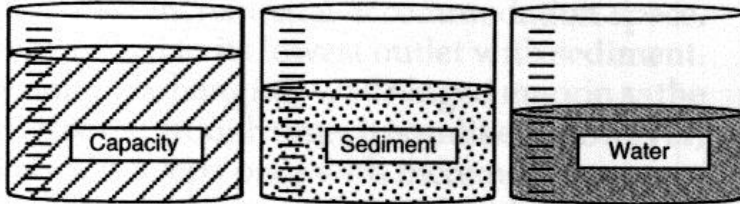


A

Excess sediment
and all water

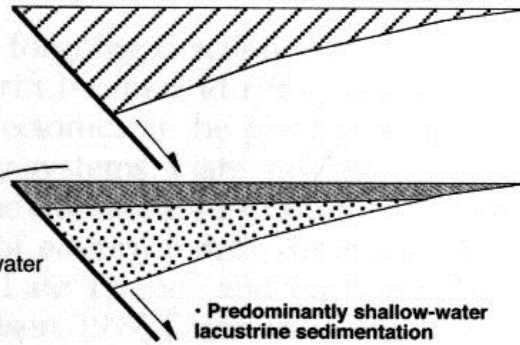


- Predominantly fluvial sedimentation
- No large-scale lakes

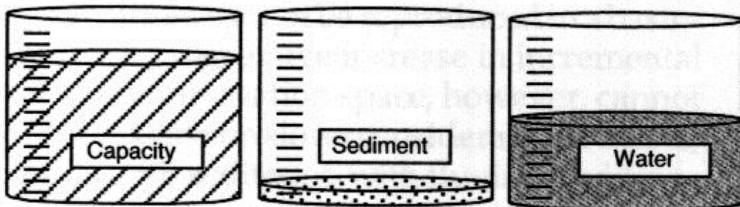


B

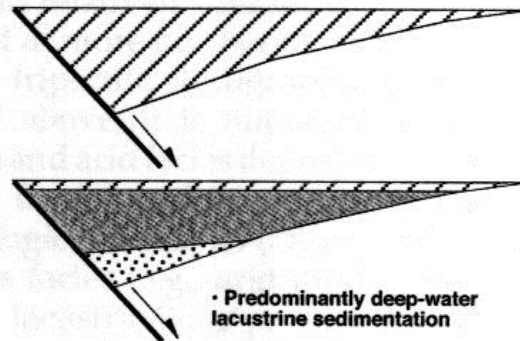
Excess water



- Predominantly shallow-water
lacustrine sedimentation



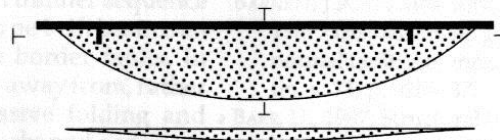
C



- Predominantly deep-water
lacustrine sedimentation

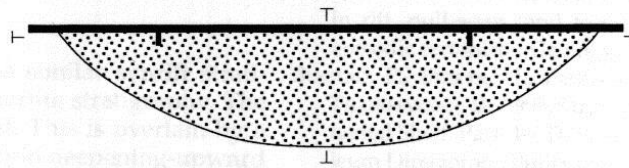
Stage 1

Capacity < sediment supply
Fluvial sedimentation



Stage 2

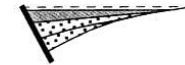
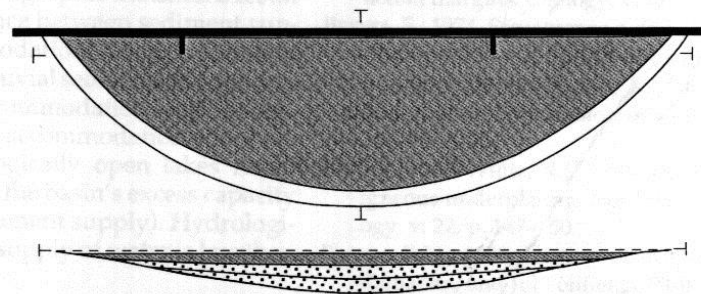
Capacity = sediment supply
Fluvial-lacustrine transition



Stage 3

Capacity > sediment supply
Water volume > excess capacity

Shallow-water lacustrine sedimentation

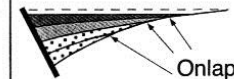
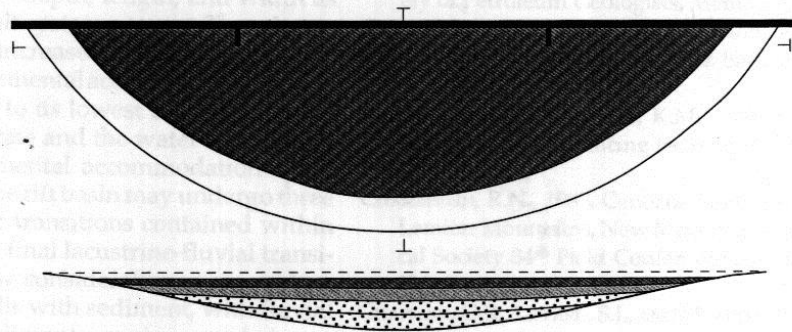


Stage 4

Capacity >> sediment supply

Water volume = excess capacity

Deep-water lacustrine sedimentation

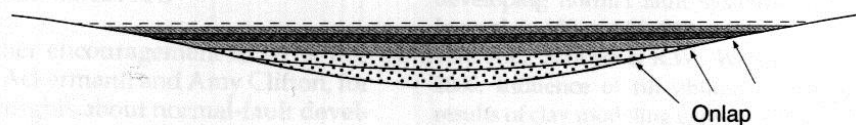


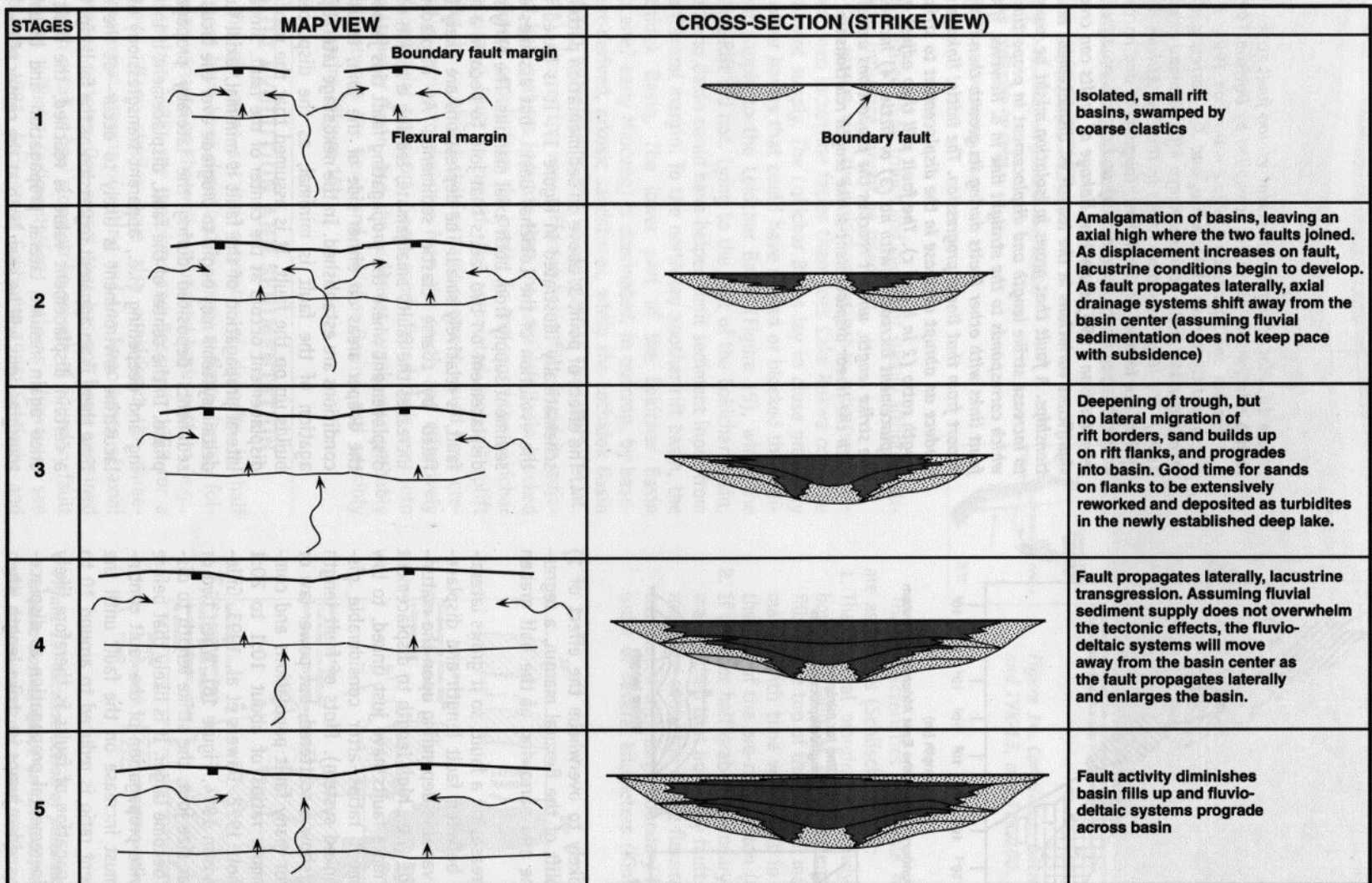
Stage 5

Capacity > sediment supply

Water volume < excess capacity

Shallow-water lacustrine sedimentation





fluvial system 0 20 km

Figure 17. Idealized strike section evolution of a half graben, emphasizing how the lateral propagation and evolution of a boundary fault zone can influence sedimentation patterns.

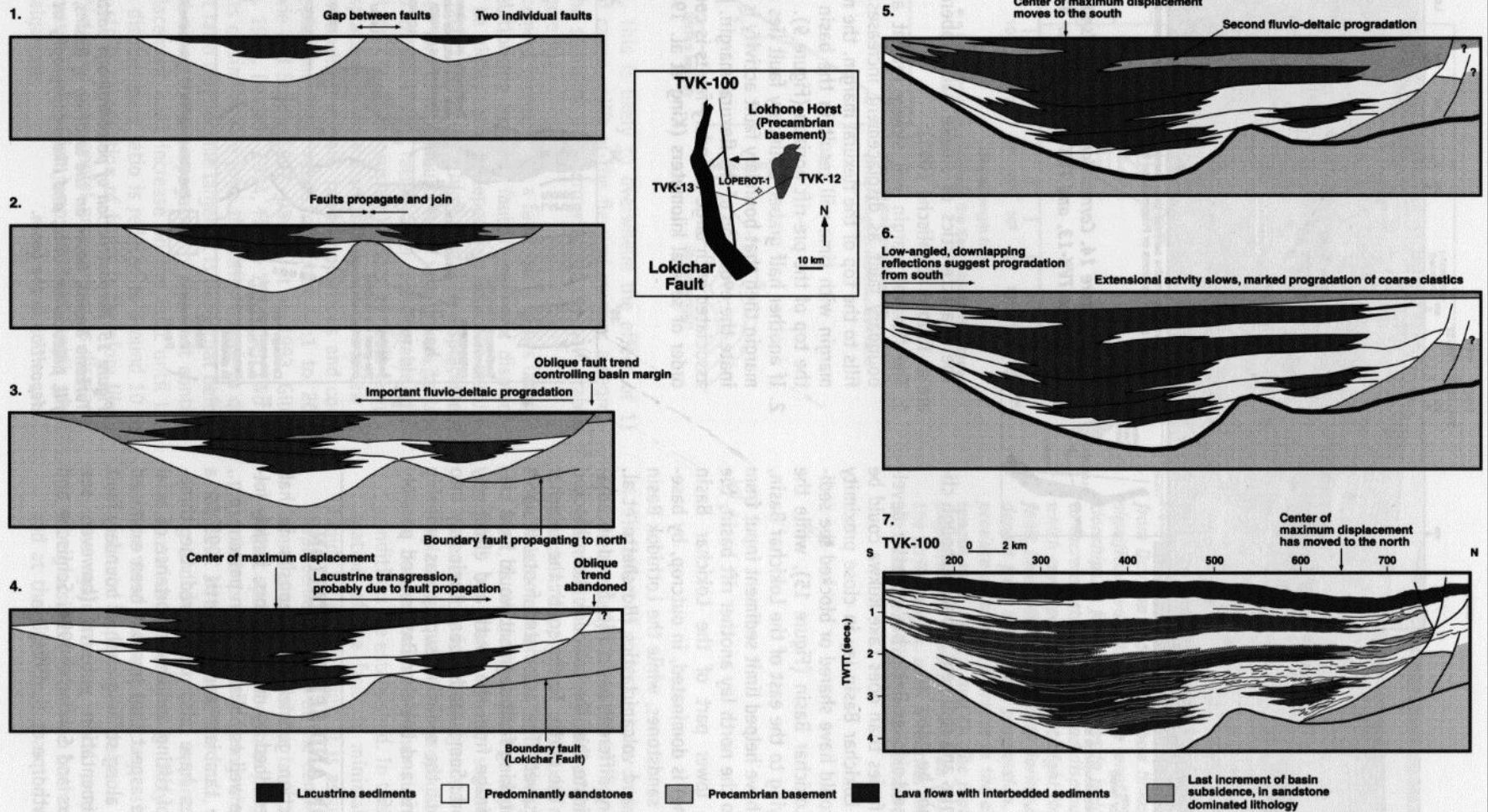


Figure 13. Evolution of seismic line TVK-100, Lokichar Basin, Kenya, based on Figure 12.